

ECMWF 2004 SEMINAR

**Recent Developments in
Numerical Methods for
Atmospheric and Ocean
Modelling**



6 to 10 September 2004

Max-Planck-Institut für Meteorologie
Max Planck Institute for Meteorology



The ICON project: development of a unified model using triangular geodesic grids



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Max Planck Institute for Meteorology





ICON: ICOsahedral grid, **N**onhydrostatic unified (**NWP**+ **climate**+**chemistry**) model

- **ICON** development team:

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- **Discussions and/or joint work: N.Botta, F. Giraldo, J.Klemp, R.Klein, D.LeRoux, D.Randall, T.Ringler, H.Tomita**





Outline

- Overview of the **ICON** development project and of the project goals
- Model **equations** and **discretization** approach
- Preliminary results of a **shallow water** model
- **Vertical** discretization
- Outlook on **future** work





Desired features for a new model

- **Unique framework** for **large/small** scale, **lower/upper** atmospheric dynamics
- **Consistency** between **conservative** discrete tracer advection and continuity equation
- **Mass conservative local grid refinement** approach without spurious interface effects: building block for a **multiscale** model





Concept of discretization approach

- Achieve the same **accuracy** and **efficiency** of advanced **NWP** models...
- ...but preserve some **discrete** equivalents of **global** invariants relevant to geophysical flow...
- ...and narrow the **gap** with **Computational Fluid Dynamics (CFD)** models.





Nonhydrostatic, compressible flow

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0$$

$$\frac{\partial \mathbf{u}}{\partial t} + \boldsymbol{\eta} \times \mathbf{u} = -\nabla K - \frac{1}{\rho} \nabla p - \nabla \Phi$$

$$\frac{\partial(\rho \varepsilon)}{\partial t} + \nabla \cdot [(\rho \varepsilon + p) \mathbf{u}] = -\nabla \cdot \mathbf{R}$$





Shallow water flow

$$\frac{\partial h}{\partial t} + \nabla \cdot (H\mathbf{u}) = 0$$

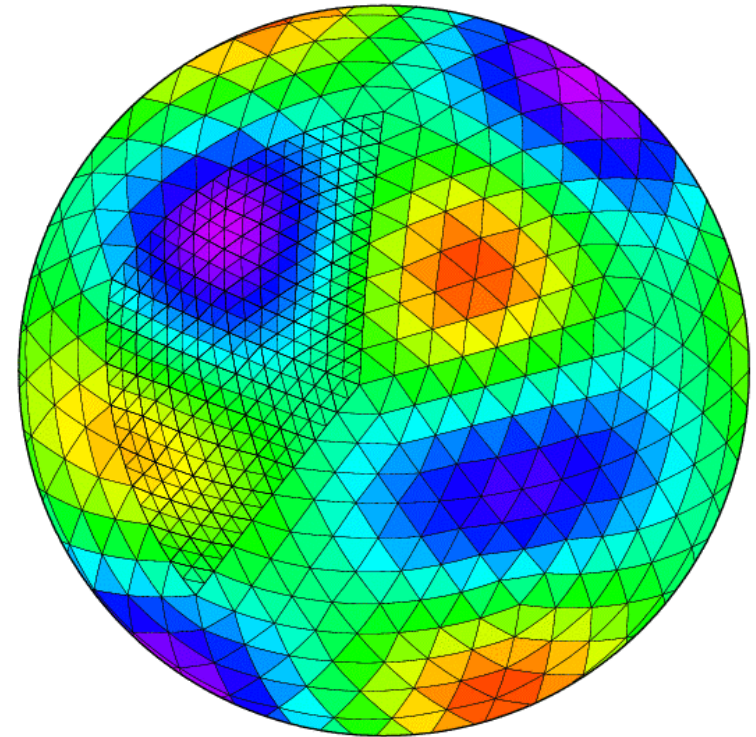
$$\frac{\partial \mathbf{u}}{\partial t} + (\zeta + f)\mathbf{k} \times \mathbf{u} + \nabla(gh + K) = 0$$

$$\frac{\partial(cH)}{\partial t} + \nabla \cdot (cH\mathbf{u}) = 0$$



Geodesic icosahedral grids

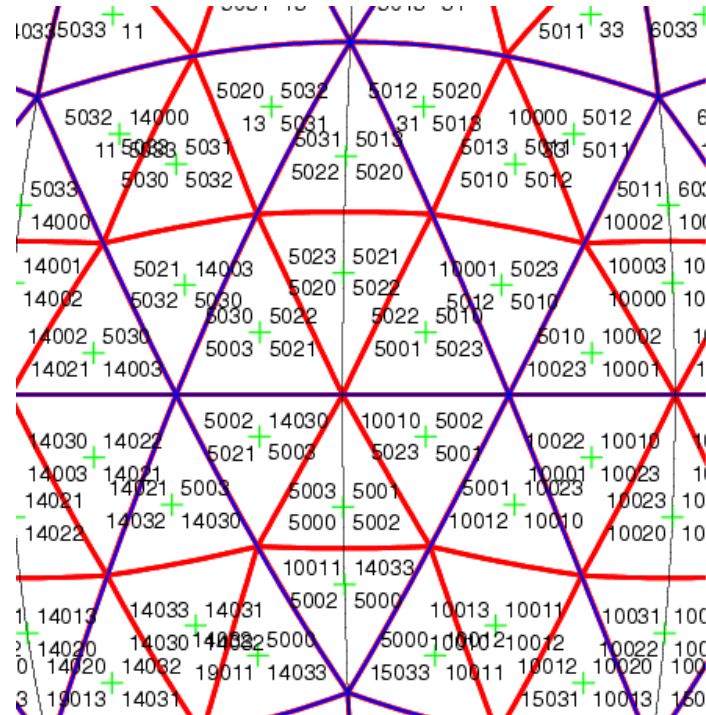
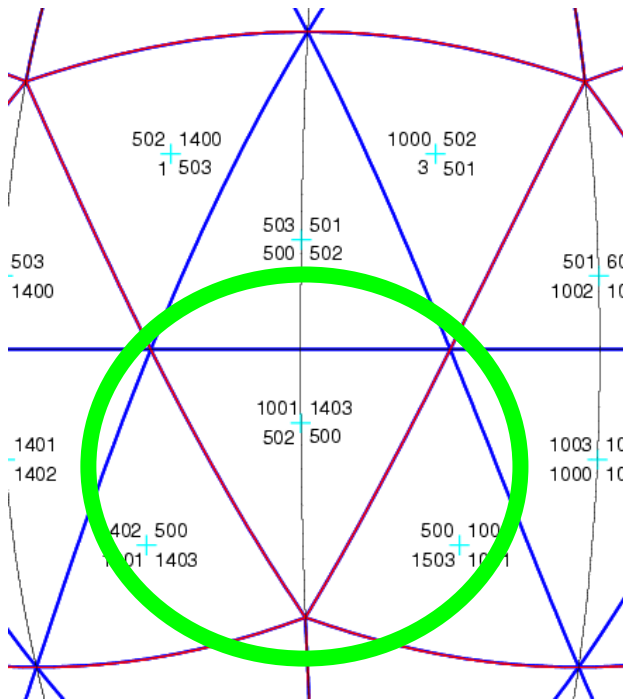
- Solve the **pole** problem
- Special case of **Delaunay** triangulation
- **Local** grid refinement
- **Multiscale** modelling





Data structures for grid representation

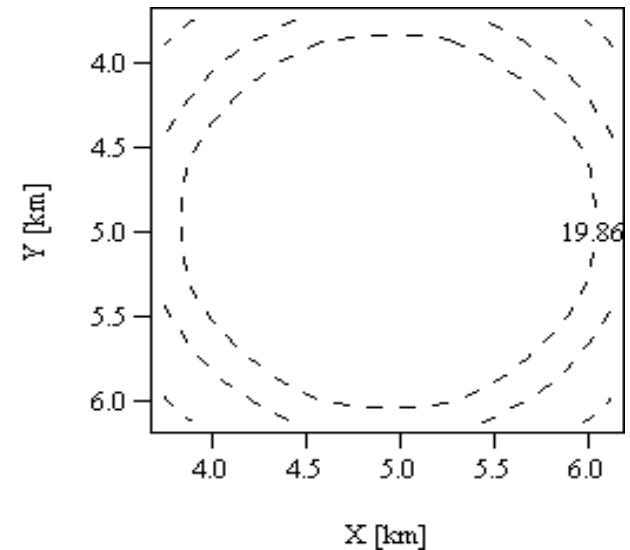
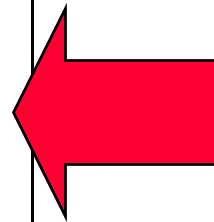
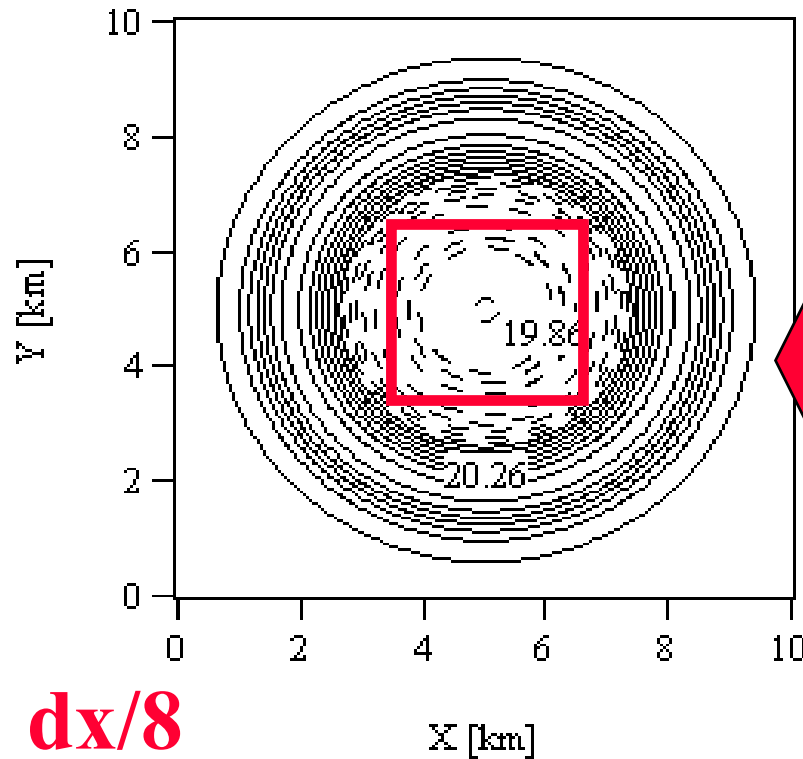
Indirect addressing that preserves **data locality**



Parallelization: horizontal data decomposition



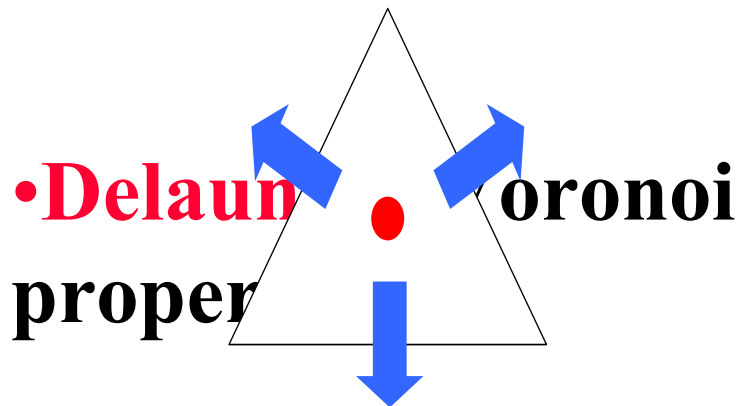
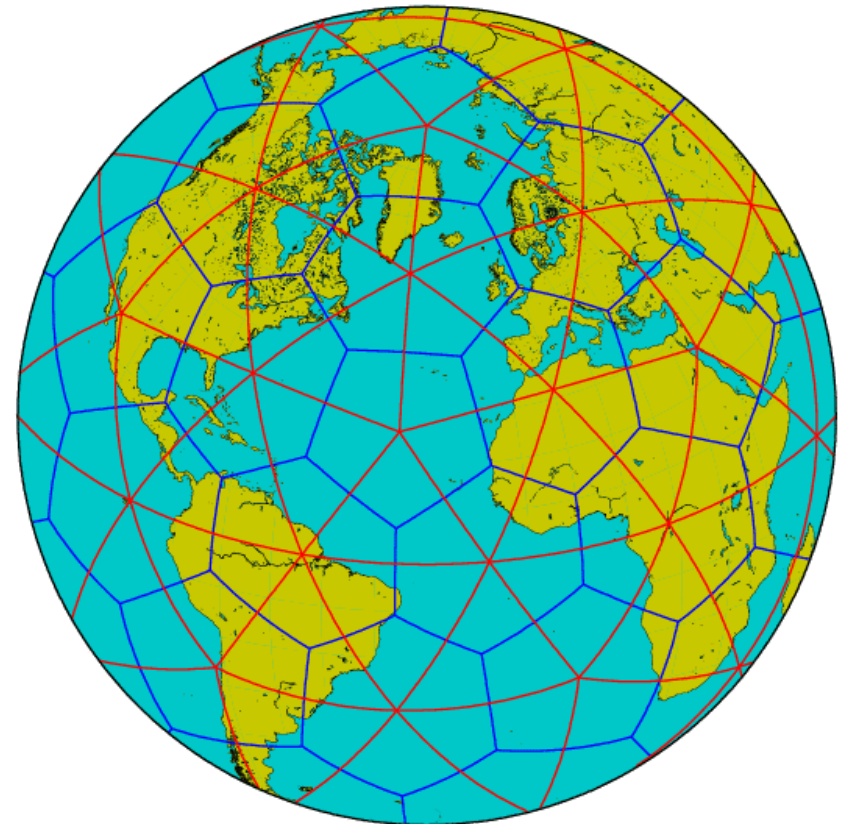
Consistent fluxes at coarse/fine interface



Edwards JCP 1996, Bornemann and Deufhard Num.Math.1996,
B. and Rosatti, IJNMF 2002

Spatial discretization

- Finite volume discretization with **triangular** control volumes:
triangular **C** grid





Spatial discretization, properties

- Vorticity at triangle **vertices**: discrete **Helmholtz** decomposition (Nicolaidis 1992)
- **No spurious vorticity production**
- **Raviart Thomas** reconstruction of velocity, **average** onto edge for tangential component

$$\mathbf{u}(\mathbf{x}) = \mathbf{u}_0 + \alpha \mathbf{x}$$





Discrete shallow water system

$$\frac{\partial h_i}{\partial t} = - \sum_{l \in C(i)} u_l H_l \sigma_{i,l}$$

$$\frac{\partial u_l}{\partial t} = -(\zeta + f)_l v_l - \delta_v (K + gh)_l$$

$$\frac{\partial (c_i H_i)}{\partial t} = - \sum_{l \in C(i)} c_l u_l H_l \sigma_{i,l}$$



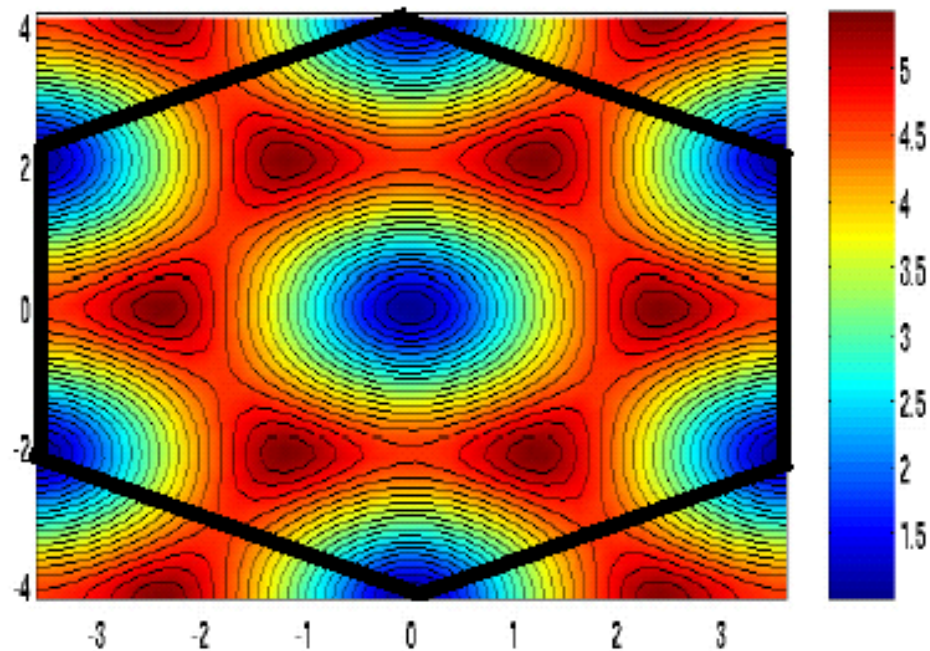


Discrete wave dispersion analysis

- **Stationary** geostrophic solution, no spurious pressure modes
- Two **physical** gravity wave modes
- Two **spurious** gravity wave modes: frequencies always **higher** than physical ones



Dispersion plot, physical mode



- **Less good** wavenumber space than quad C
- **Zero group velocity** at high wavenumbers



Discrete global invariants

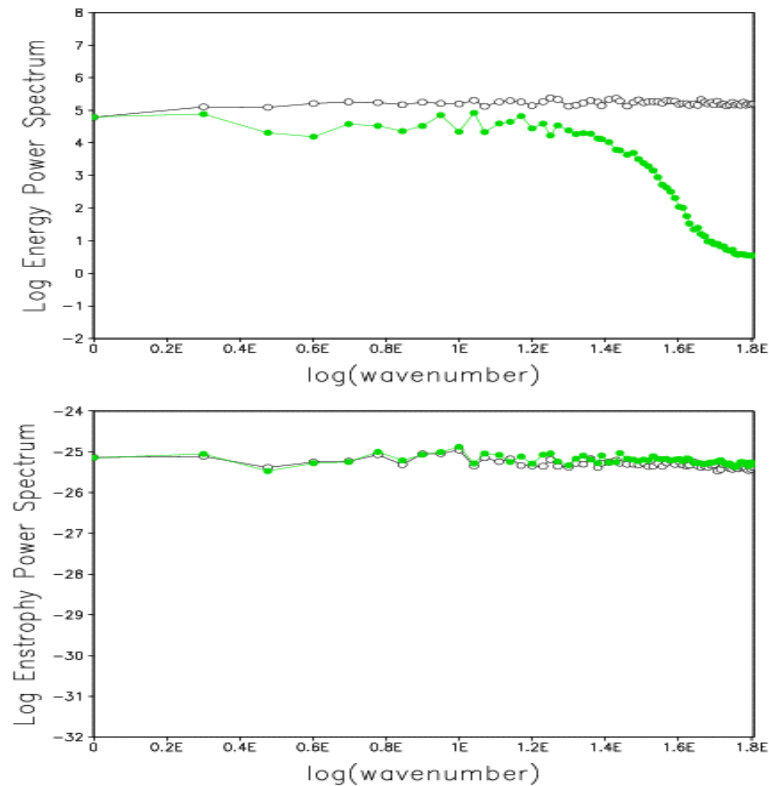
- Mass conservation, **consistent** discretizations of continuity equation and tracer transport
- Potential vorticity conservation, **no spurious vorticity production**
- Potential **enstrophy** conserving variant, **energy** conserving variant: **Sadourny JAS 1975**





Random initial data, f plane

Relative vorticity
after 1000 days
integration with
random initial
data (numerical
test carried out
by Todd Ringler,
CSU)





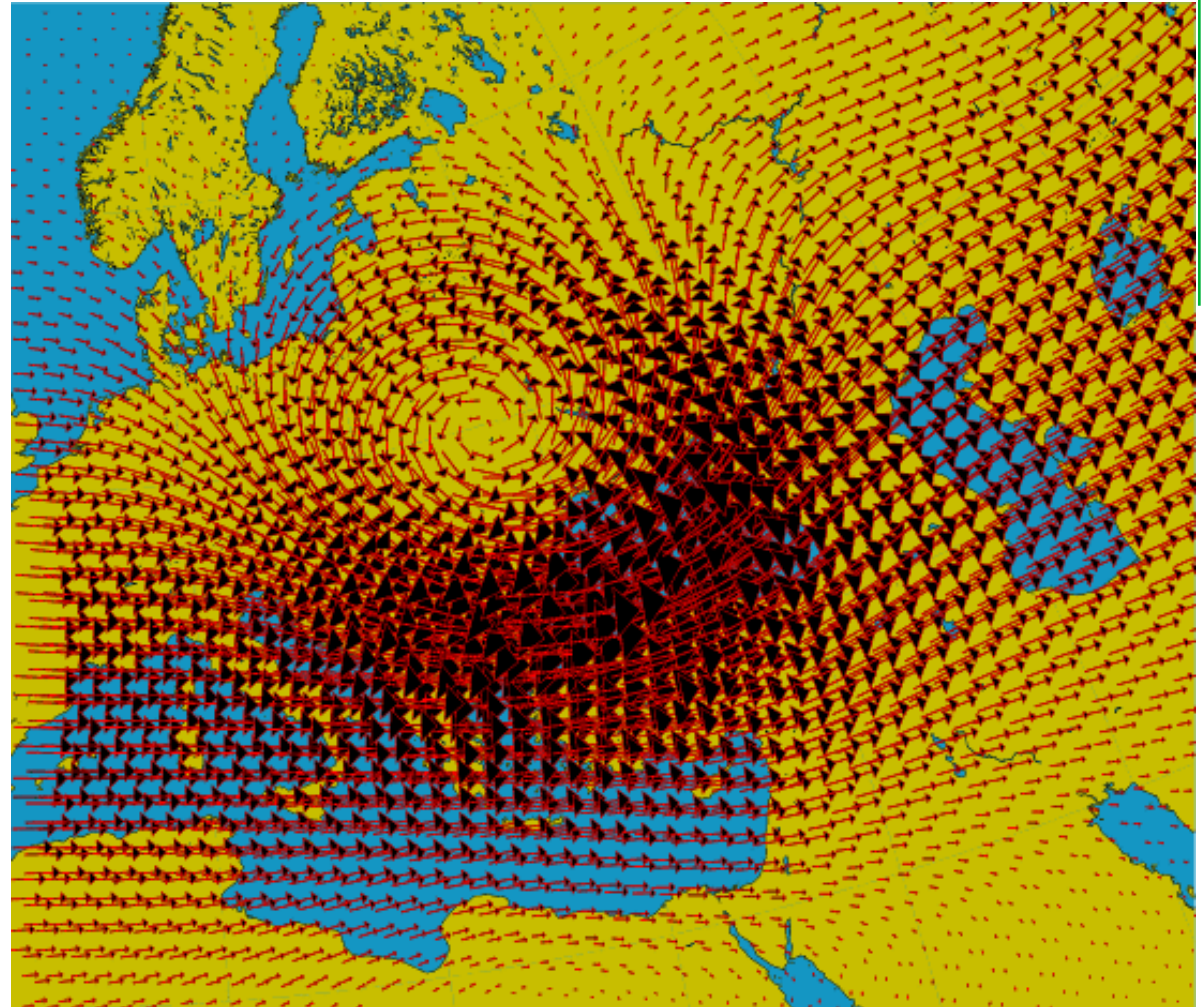
Semi-implicit time discretization

$$\begin{aligned}
 u_l^{n+1} &= u_l^n - \Delta t (\tilde{\zeta}^{n+1/2} + f)_l v_l^{n+1/2} \\
 &\quad - \Delta t \delta_v (\tilde{K}^{n+1/2} + gh^{n+1/2})_l
 \end{aligned}$$

$$h_i^{n+1} = h_i^n - \Delta t \sum_{l \in C(i)} u_l^{n+1/2} H_l^n \sigma_{i,l}$$



Idealized vortex, day 2



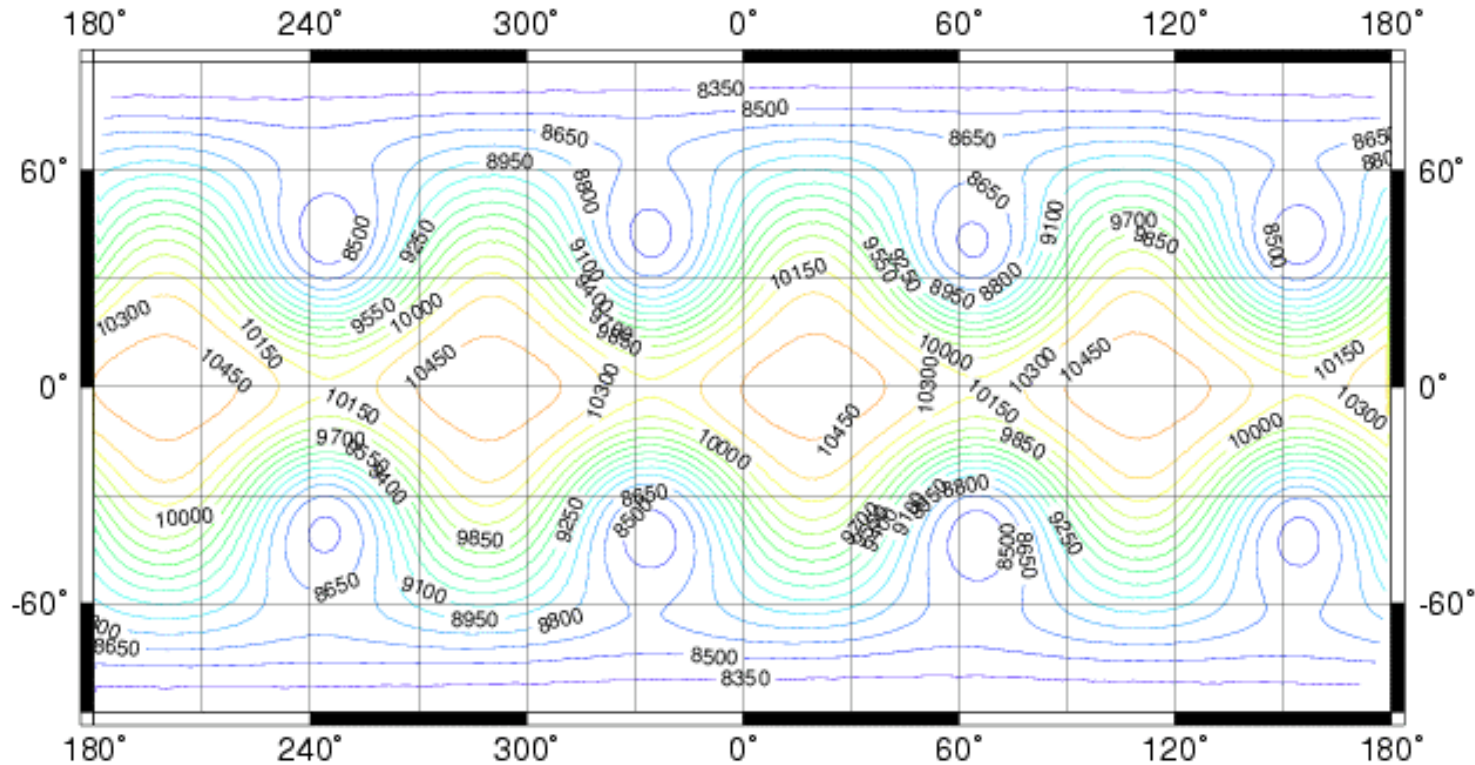
Maximum
resolution **40 km**

Maximum gravity
wave Courant
number **7**

(dt=900 s)

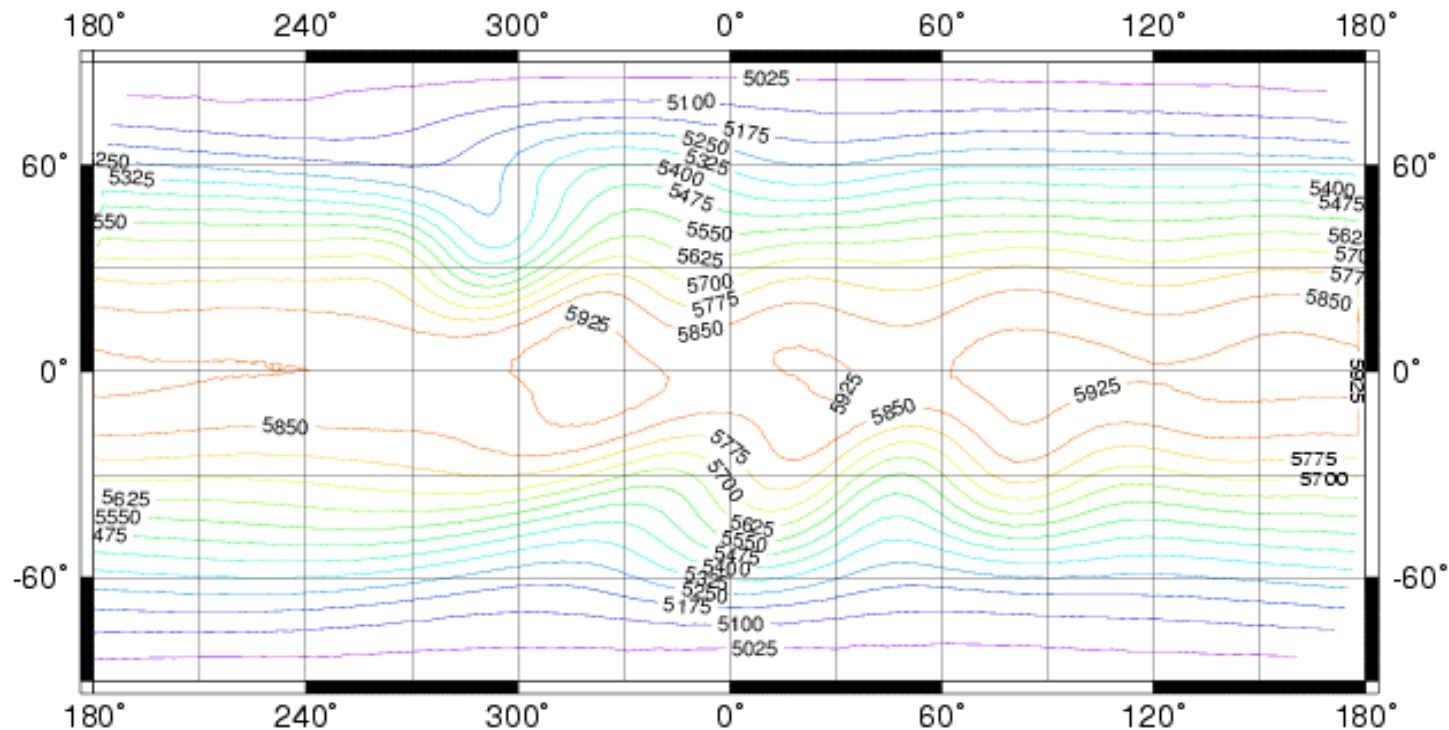


Rossby Haurwitz wave, day 10





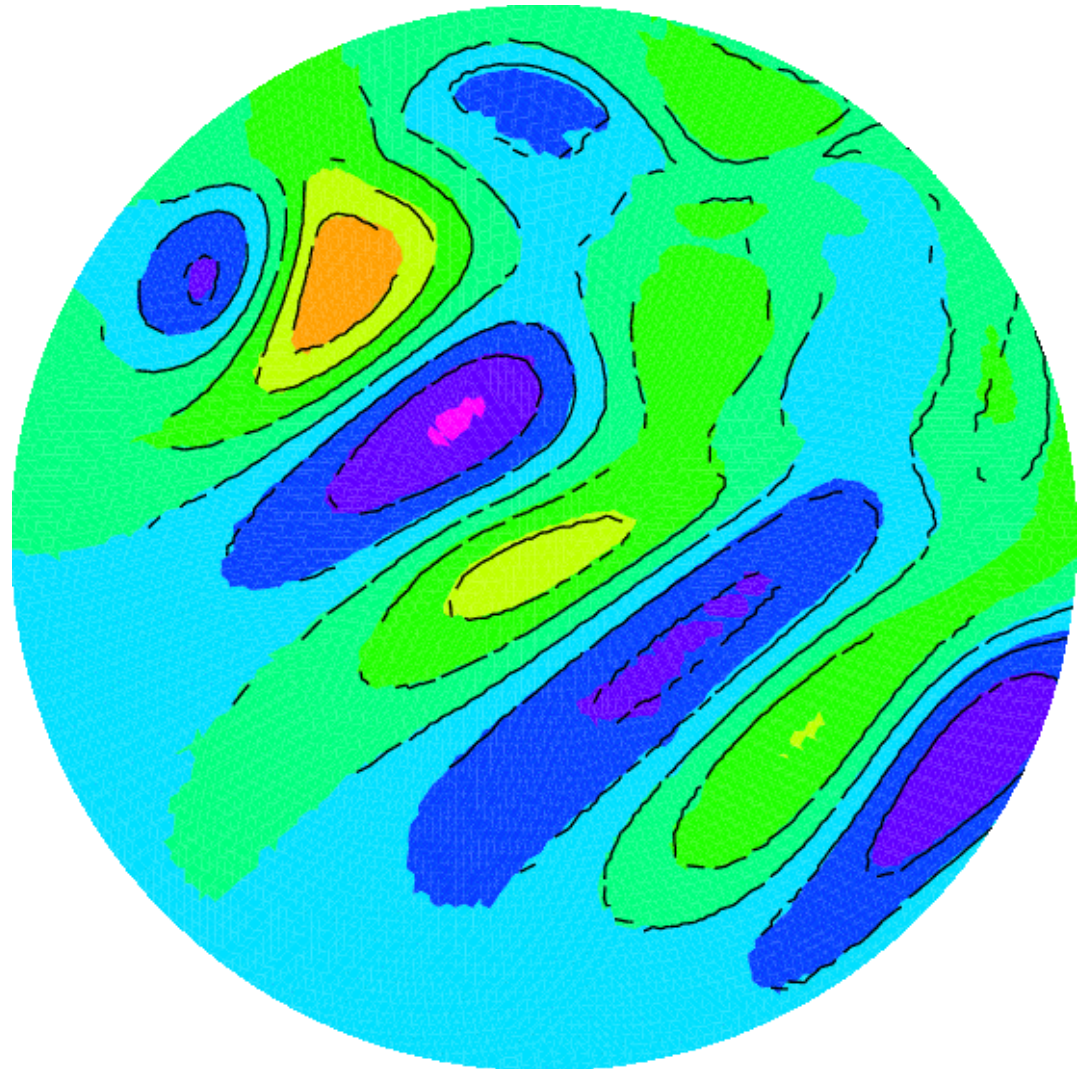
Flow over a mountain, day 10



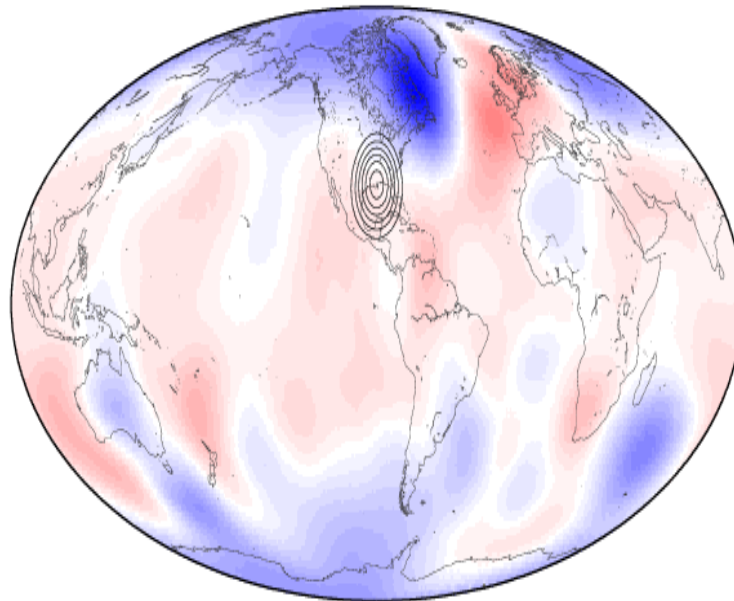
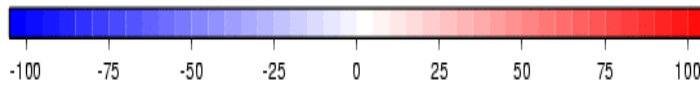
Flow over a mountain: relative vorticity, day 10

Colour shading:
ICON model
results

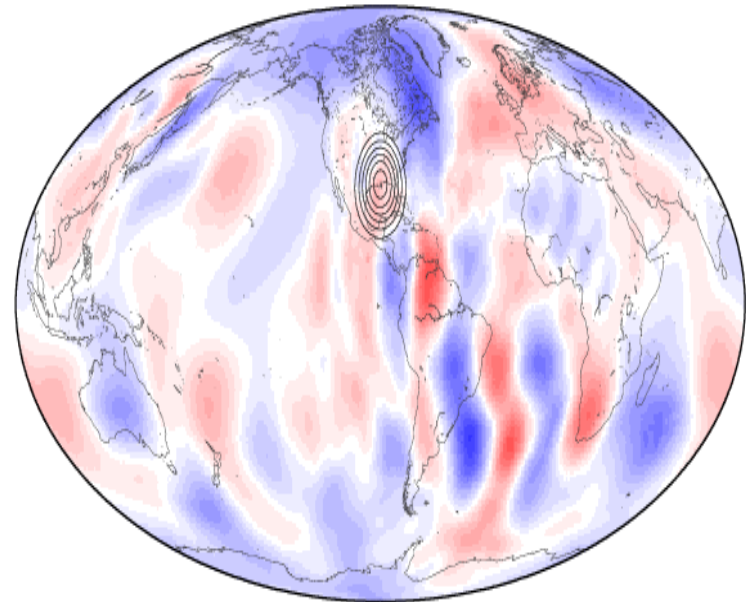
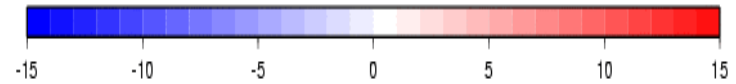
Black contours:
NCAR reference
spectral model



Height field error at day 15



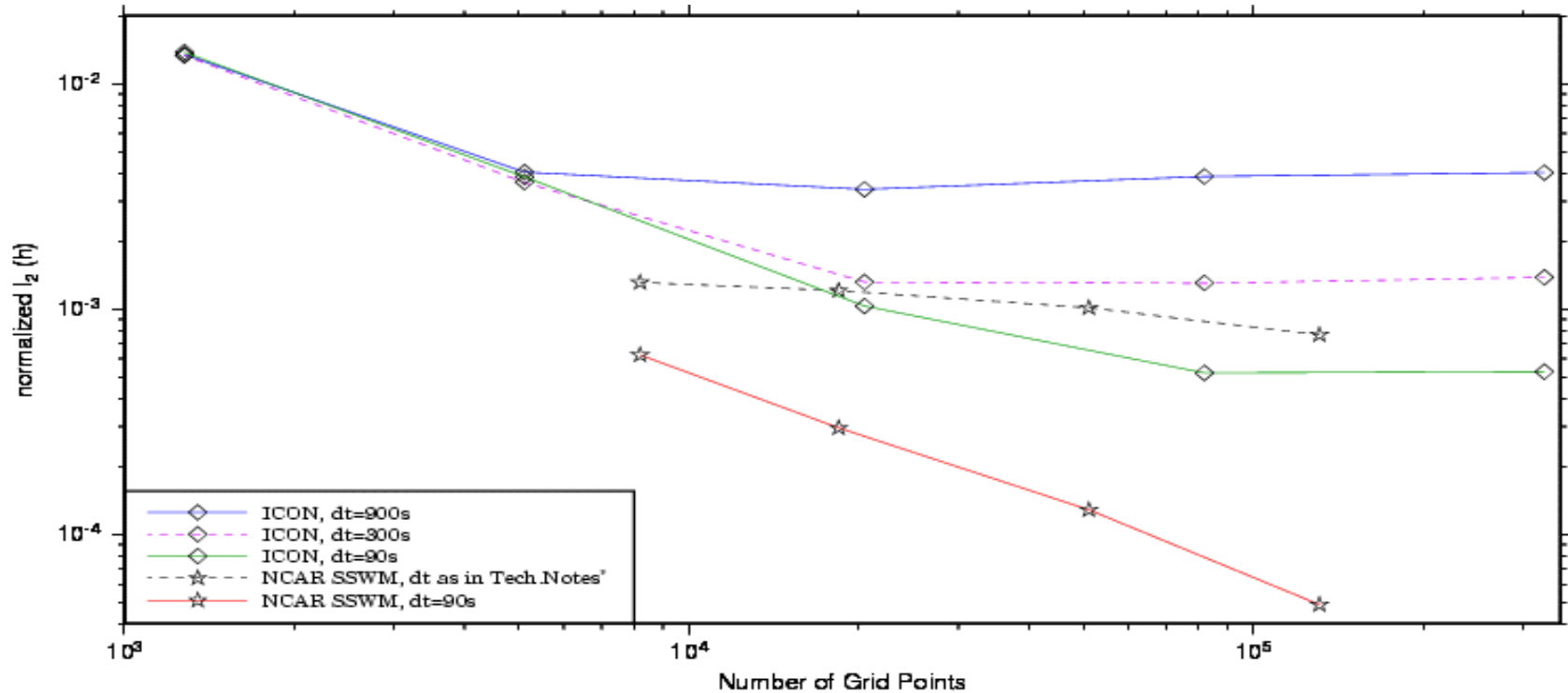
$dx \approx 120 \text{ km}, dt = 900 \text{ s}$



$dx \approx 60 \text{ km}, dt = 90 \text{ s}$



Error at day 15, convergence test



TEST CASE 5. L2 NORM AT DAY 15. NCAR SSWM T213/dt=90s as reference^{*}

NCAR SSWM: T42, T63, T106 and T170; ICON: refinement levels 4 to 8, optimized grids;

^{*}black dashed line: reference as in NCAR Tech.Notes: T213/dt=360s





“Shallowness is the greatest vice”

Oscar Wilde



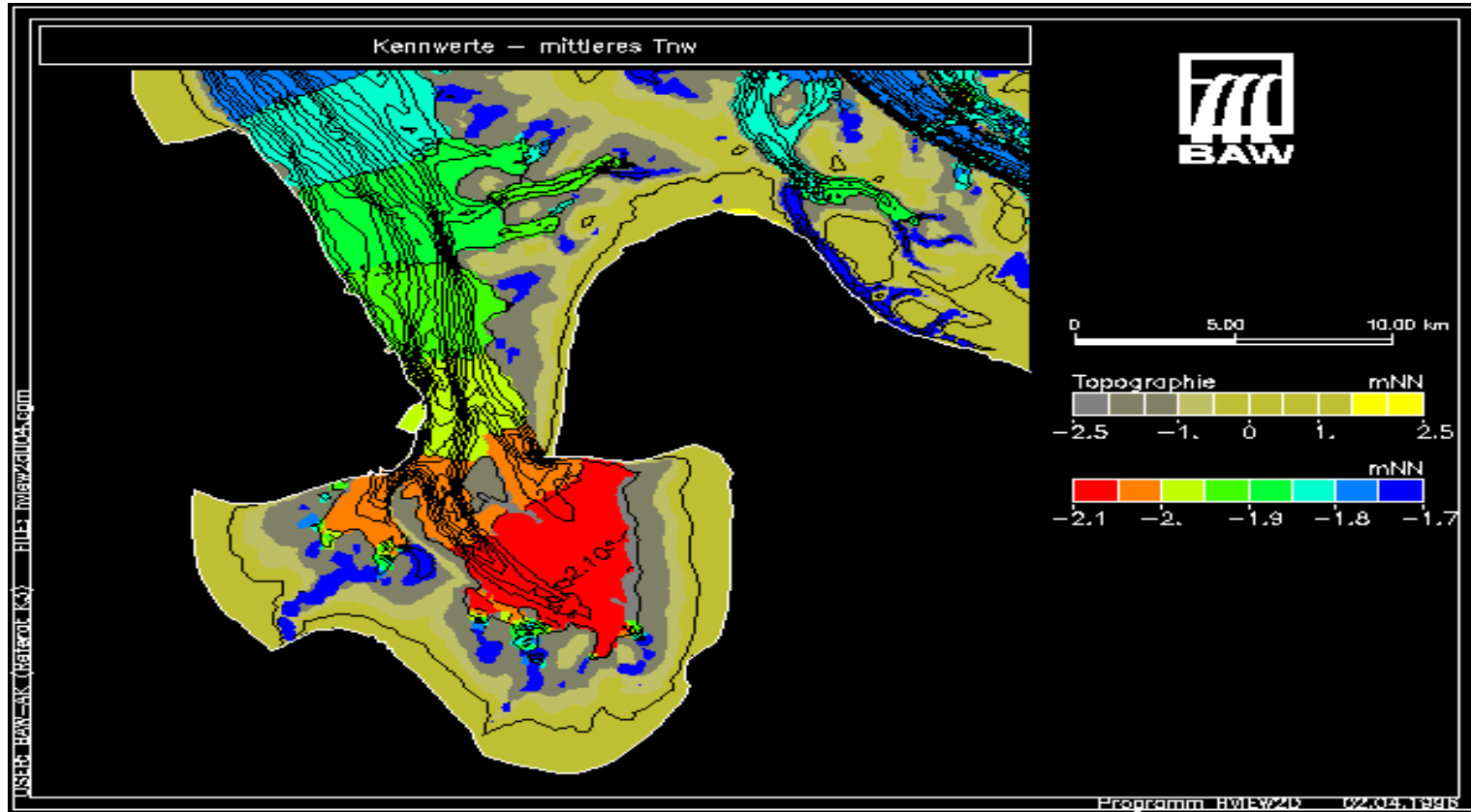


Options for vertical discretization

- **Hybrid pressure** vertical coordinate + new horizontal discretization: preliminary 3D **hydrostatic** ICON model
- **Terrain following** normalized height coordinate + new horizontal discretization: **first choice** for operational nonhydrostatic model
- **Non normalized** height coordinate: **cut cells** (B., JCP 2000, Rosatti and B., Proc. ICFD, 2004)



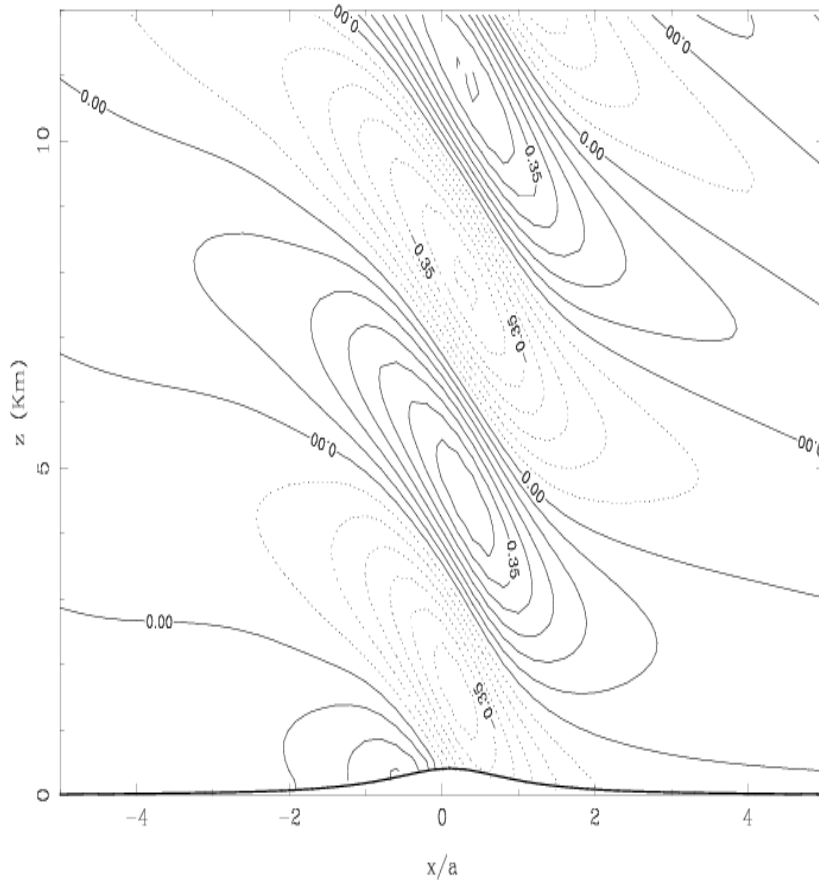
Nonhydrostatic coastal modelling



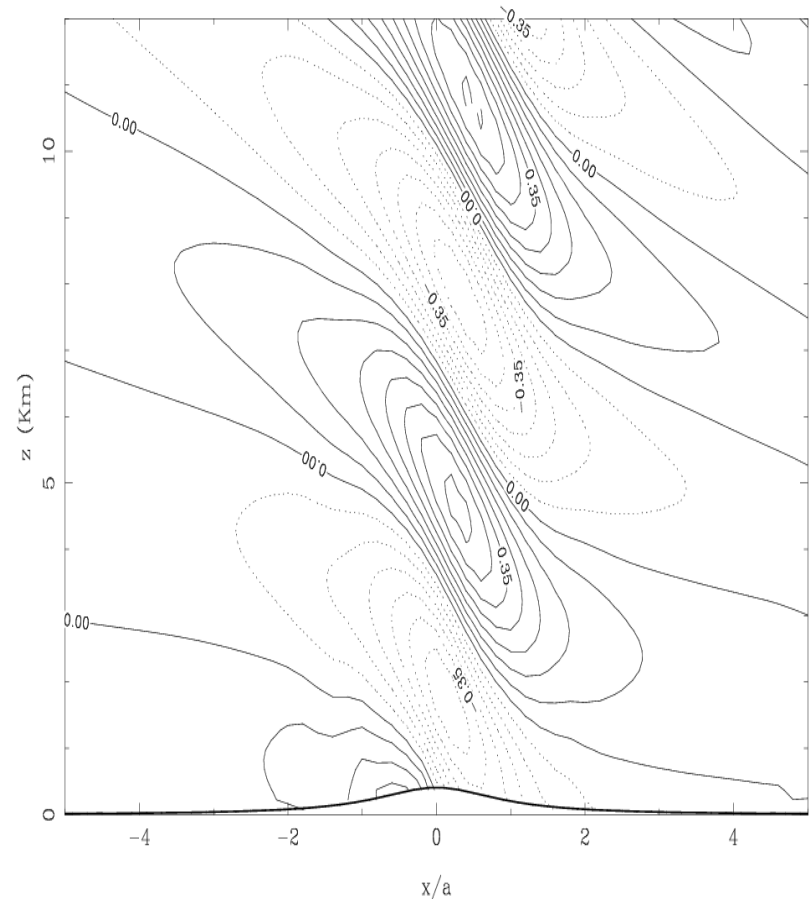
- Results: G.Lang, Bundesanstalt für Wasserbau, Germany
- Numerical model: Casulli and Walters, IJNMF, 2000



Cut cells + RBF interpolation



Terrain following model (LM)



Cut cell nonhydrostatic dynamical core (ARPA Bologna)





Computational advantages of cut cells

| | CPU time for 1 hour | CPU time solver | COMM time solver |
|-----------|---------------------|-----------------|------------------|
| S | 88.95 s | 45.03 s | 11.95 s |
| E | | | |
| SI | 56.40 s | 26.16 s | 5.12 s |
| Z | | | |

| | Residual 1% of initial value | Residual 0.1% of initial value | Residual 0.01% of initial value |
|----------|-------------------------------------|---------------------------------------|--|
| S | 6 iter | 21 iter | 50 iter |
| I | | | |
| S | 8 iter | 17 iter | 21 iter |
| I | | | |
| Z | | | |

Simulations run by D.Cesari (ARPA Bologna)





Future work

- Shallow water model on **locally refined** grids: optimized data structure and parallelization
- **Hydrostatic**, 3D model on **locally refined** grids
- Coupling to **existing MPI-M/DWD** physics packages, impact of **spurious modes** on simulations with full physics
- **Sensitivity** of results to local refinement

