

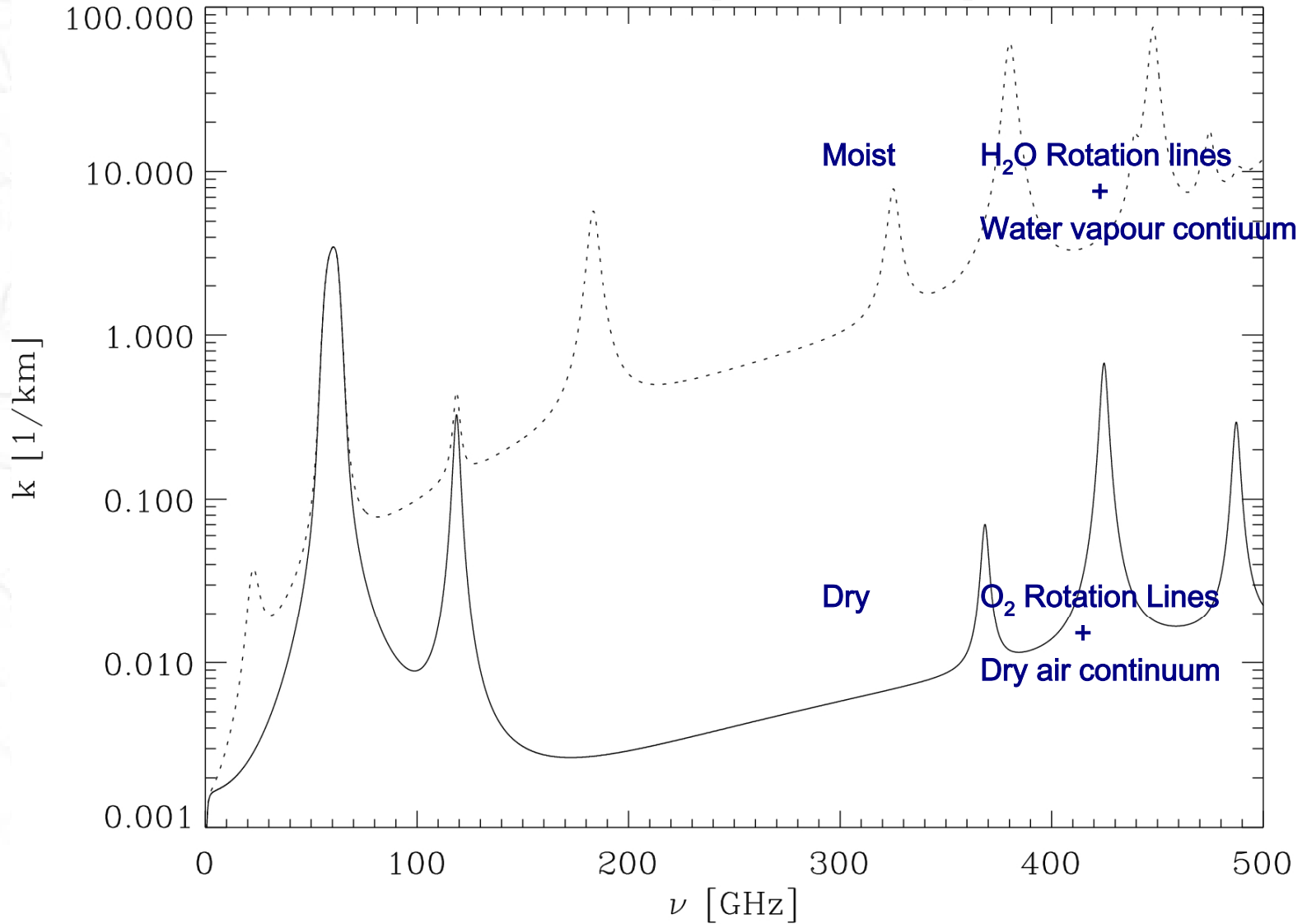
Sources of Biases in Microwave Radiative Transfer Modelling

Peter Bauer, Sabatino Di Michele, ECMWF
William Bell, Stephen English, The Met Office
Christian Mätzler, University of Bern
Ralf Bennartz, University of Wisconsin, Madison

μ - Spectrum



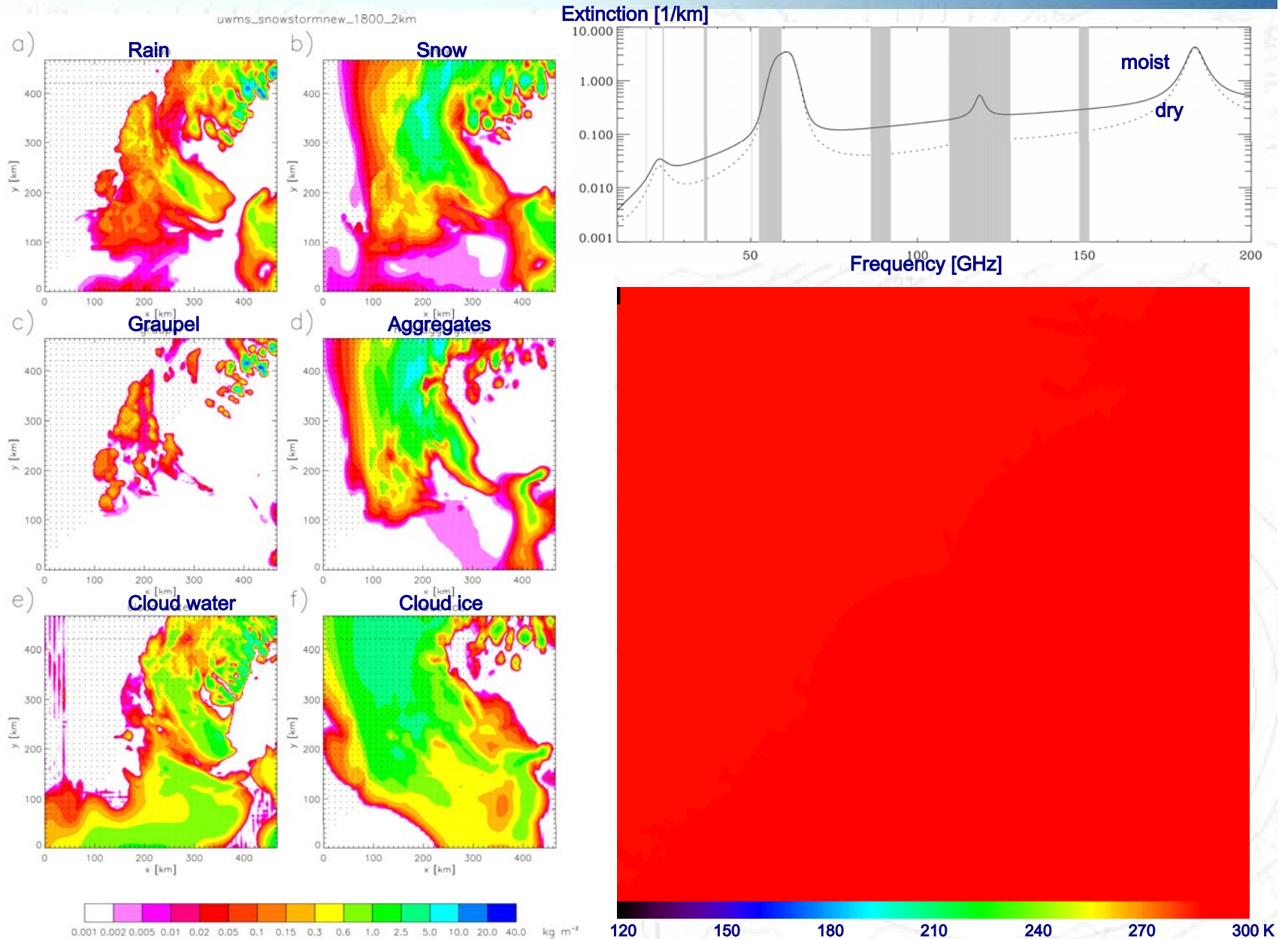
T = 288 K, p = 1000 hpa



Combined cloud-radiative transfer modelling



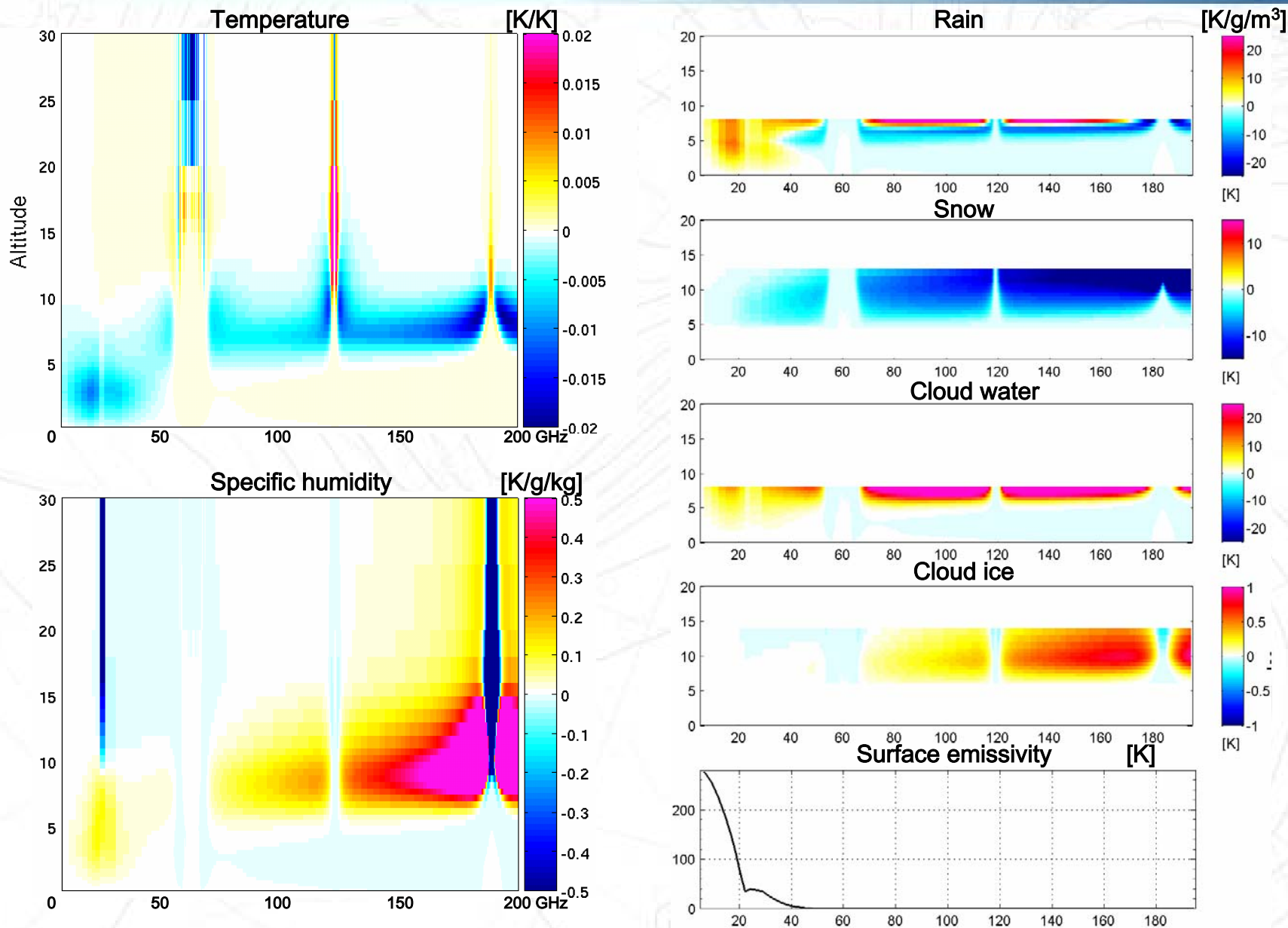
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Microwave H_{TB} : Single profile over ocean



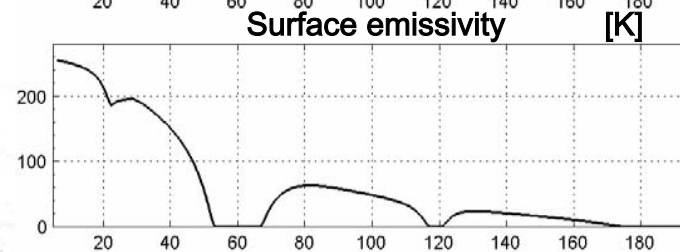
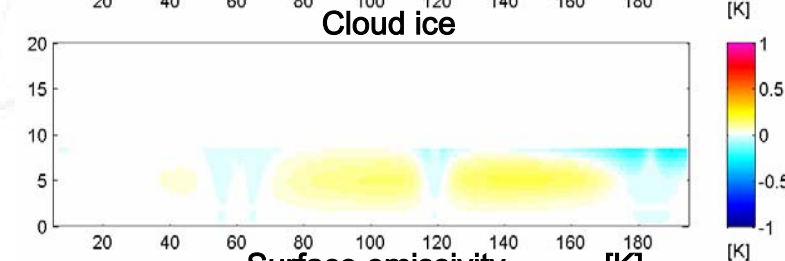
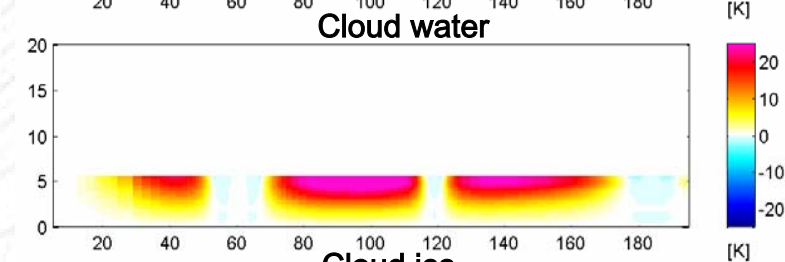
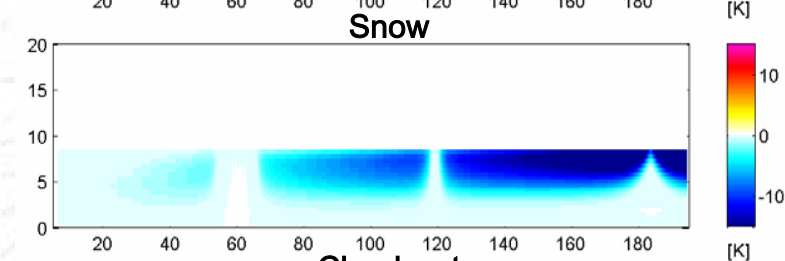
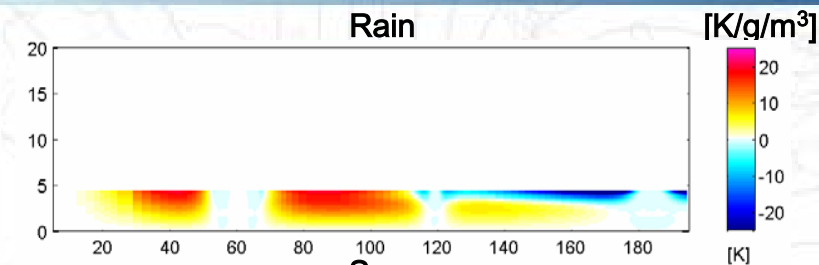
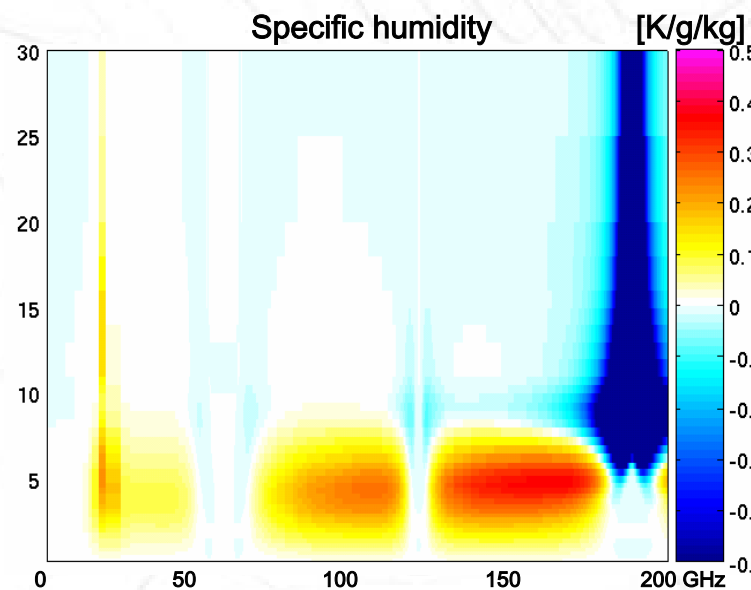
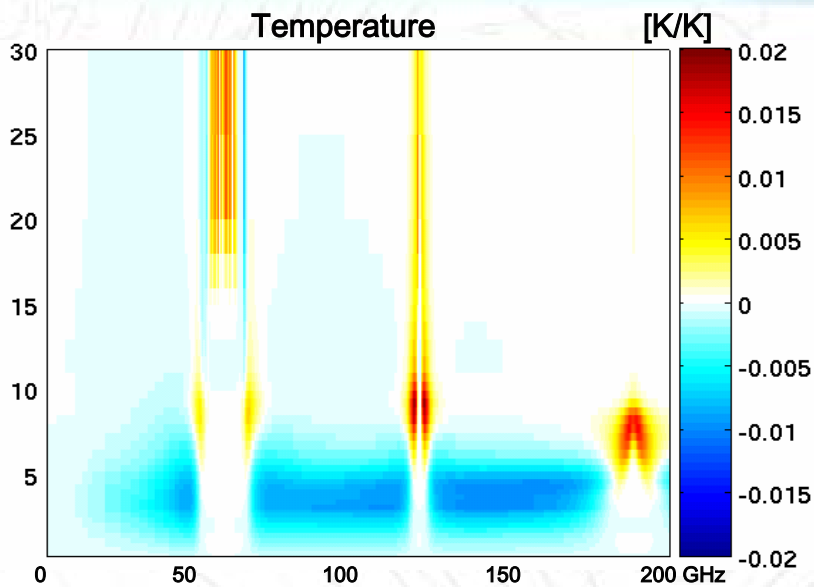
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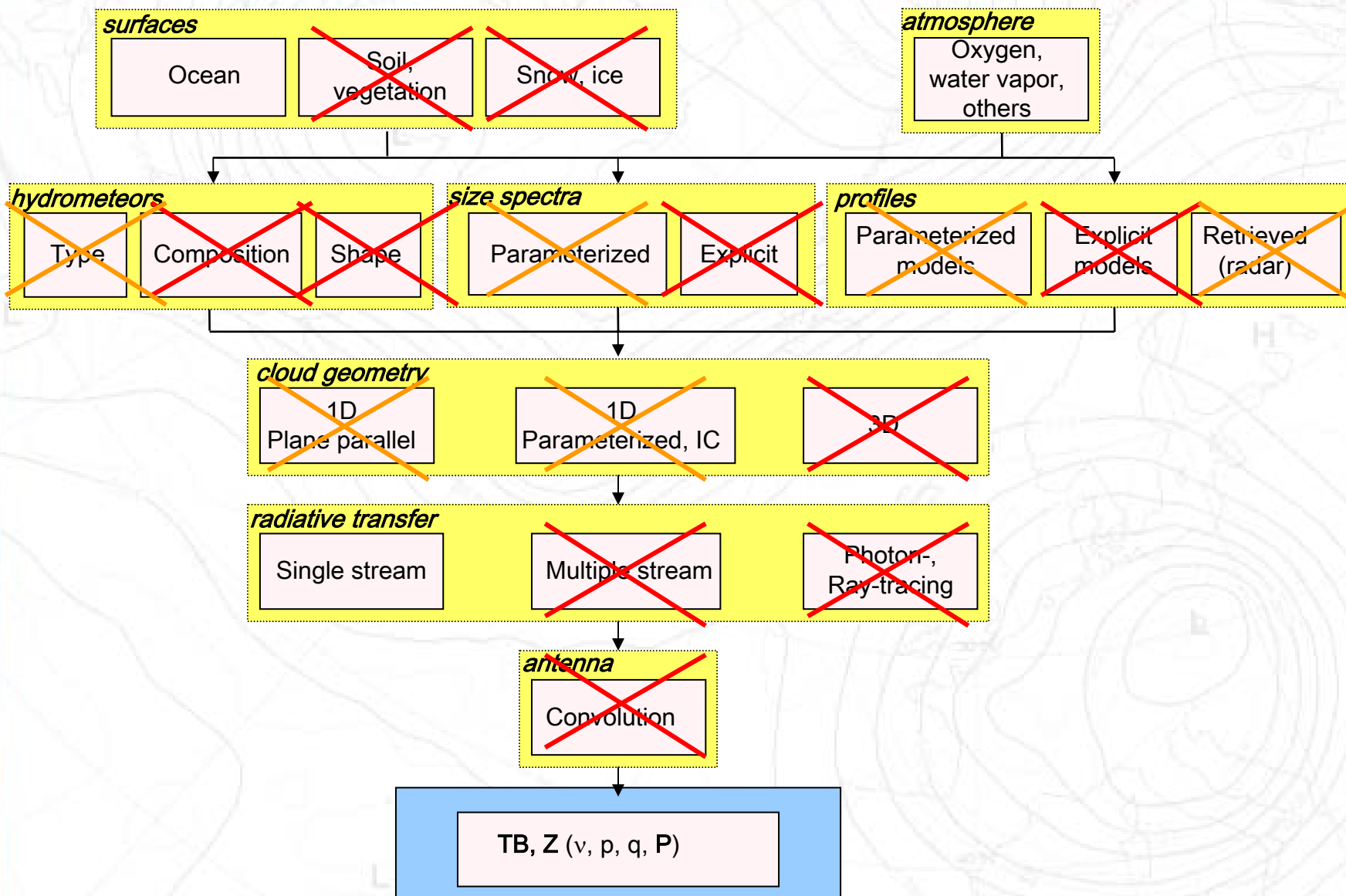
Microwave H_{TB} : Single profile over land



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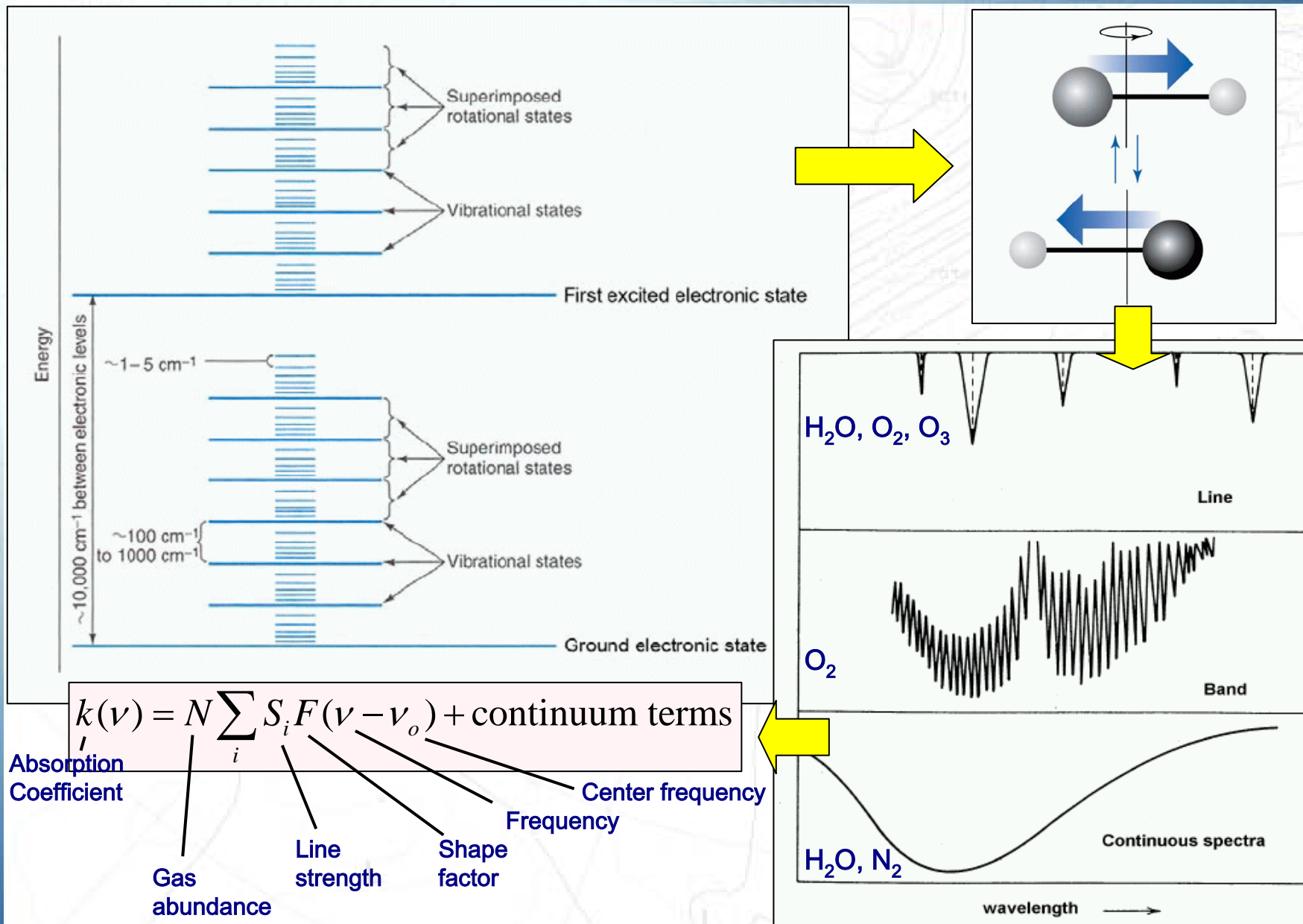
Components of μ -RT Modelling



Atmospheric absorption



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$$k(\nu) = N \sum_i S_i F(\nu - \nu_0) + \text{continuum terms}$$

Absorption Coefficient

Gas abundance

Line strength

Shape factor

Frequency

Center frequency

H₂O, N₂

Continuous spectra

wavelength →

Inter-model evaluation



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117 Profiles, RTTOV vs. LBL model

AMSU-A		RTTOV-8 117 independent set		
Channel #	NeDT degK	Mean bias degK	St. dev. degK	Max diff degK
1	0.20	0.00	0.01	0.04
2	0.24	-0.01	0.02	0.12
3	0.19	-0.02	0.03	0.18
4	0.13	0.01	0.01	0.07
5	0.13	0.02	0.01	0.08
6	0.11	0.01	0.01	0.06
7	0.12	0.00	0.01	0.03
8	0.13	0.00	0.00	0.01
9	0.15	-0.01	0.01	0.07
10	0.19	0.19	0.16	0.39
11	0.20	0.00	0.04	0.27
12	0.31	-0.02	0.07	0.44
13	0.42	-0.05	0.09	0.58
14	0.70	-0.04	0.06	0.41
15	0.10	0.07	0.10	0.34

117 Profiles, RTTOV vs. LBL model

AMSU-B		RTTOV-8 117 independent set		
Channel #	NeDT degK	Mean bias degK	St. dev. degK	Max diff degK
1	0.32	0.07	0.10	0.33
2	0.71	-0.02	0.08	0.73
3	1.05	-0.01	0.07	0.74
4	0.69	-0.01	0.04	0.44
5	0.57	0.00	0.05	0.40

Inter-model >> LBL-parameterized
model differences model differences

All vs. CIMSS MWLBL model

Model	AMSU-06		AMSU-10		AMSU-14		AMSU-18	
	std	bias	std	bias	std	bias	std	bias
RTTOV-7/8	0.04	-0.06	0.15	0.25	1.36	0.90	0.35	-0.39
OPTRAN	0.09	0.00	0.05	-0.04	0.73	-1.97	0.10	0.00
AER_OSS	0.06	0.13	0.04	0.03	0.09	0.14	0.14	-0.16
MIT	0.01	0.00	0.04	-0.04	0.08	-0.09	0.19	-0.40
RAYTHEON	0.42	-0.57	0.17	0.24	0.20	0.60	0.50	-0.07
AER_LBL	0.06	0.13	0.05	0.03	0.09	0.16	0.14	-0.15
MSCMWLBL	0.03	0.05	0.03	0.04	0.20	0.51	0.32	-0.36
ATM	0.19	0.46	0.07	0.08	0.11	0.23	0.24	-0.28

(Saunders et al. 2005)

Evaluation with measurements



Table 2.8. Evolution of O₂ and H₂O line and continuum parameters in the Millimeter-wave Propagation Model of Liebe (1977-1993). The O₂ parameters include line-mixing coefficients.

Year of publ.	O ₂ parameters	H ₂ O lines	H ₂ O continuum
1977	M	O	O
1978	M	-	-
1981	R	O	-
1984	-	O	M
1987	-	O	M
1989	R	O	-
1992	M	-	-
1993	R	O	O

key
 M: new measurements by Liebe and co-workers
 R: revised analysis of earlier measurements by Liebe and co-workers
 O: revision based on measurements by others
 -: no change from previous version

Table 2.10. Calculated minus measured brightness-temperature differences (K) at Wallops Island, VA from England *et al.* (1993).*

frequency, GHz	PWV	average T _B , K	MPM89	W76
20.7	low	14.1	+0.38	+0.80
20.7	high	30.0	+0.15	+1.00
22.2	low	18.8	+0.56	-0.49
22.2	high	45.0	-0.76	-4.30
31.4	low	13.5	-0.33	+0.31
31.4	high	20.0	-1.10	+0.22

low PWV: <1 cm; high PWV: 2.1 cm.

Table 2.11. Calculated minus measured brightness-temperature differences (K) at Nauru Island (tropical western Pacific) from Westwater *et al.* (2003). Average PWV = 4.7 cm. The values in parentheses are 99% confidence intervals for the final digit.

frequency, GHz	average T _B , K	MPM87	MPM93	R98
23.8	65	+0.80 (67)	+3.90 (70)	+0.69 (67)
31.4	32	-0.16 (32)	+3.37 (36)	+0.86 (33)

Table 2.12. Calculated minus measured 31.4-GHz brightness-temperature differences (K) at three sites, from Marchand *et al.* (2003).

location	average PWV, cm	MPM87	R98	MonoRTM
Nauru	4.0	-0.35	+0.65	-0.51
Oklahoma	1.0	-0.04	+0.42	+0.05
Alaska	0.7	-0.79	-0.34	-0.84

Ground-based intercomparison

Satellite data based intercomparisons

Meissner and Wentz (2003)

SSM/I: retrievals tuned with 19-37 GHz observations, verified against 85 GHz to within 1.2 K

Pumphrey and Bühler (2000)

MLS/MAS: 183 GHz line shift verified to within 0.2 MHz/torr

Bühler (2005)

ASUR: 626 GHz line shift verified to within 0.15 MHz/torr

Rosenkranz and Barnett (2005)

HSB: 0.2-0.8 K

Rosenkranz (2003)

AMSU-A: -0.23 – 0.42 K

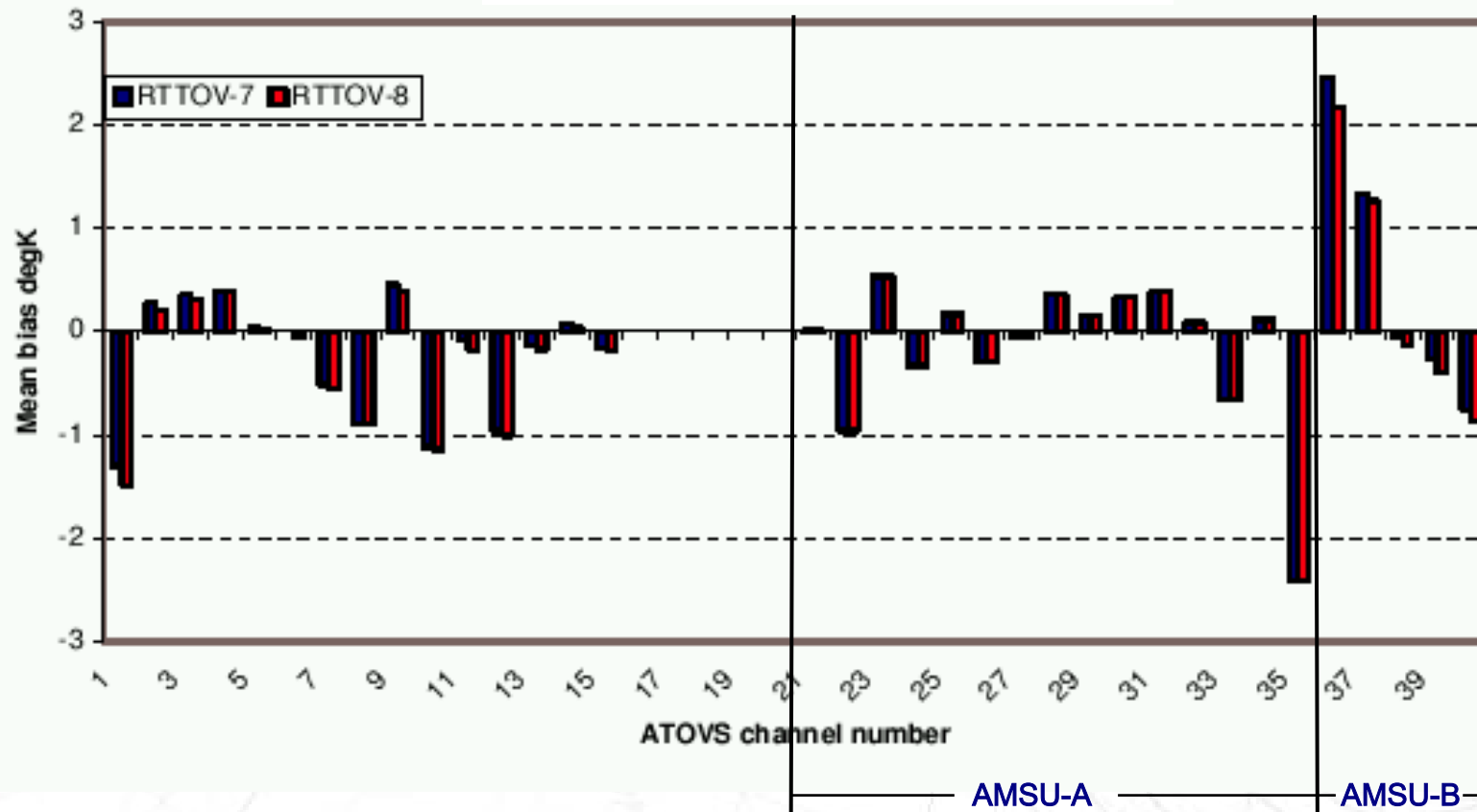
(Mätzler et al. 2005)

Evaluation inside NWP system



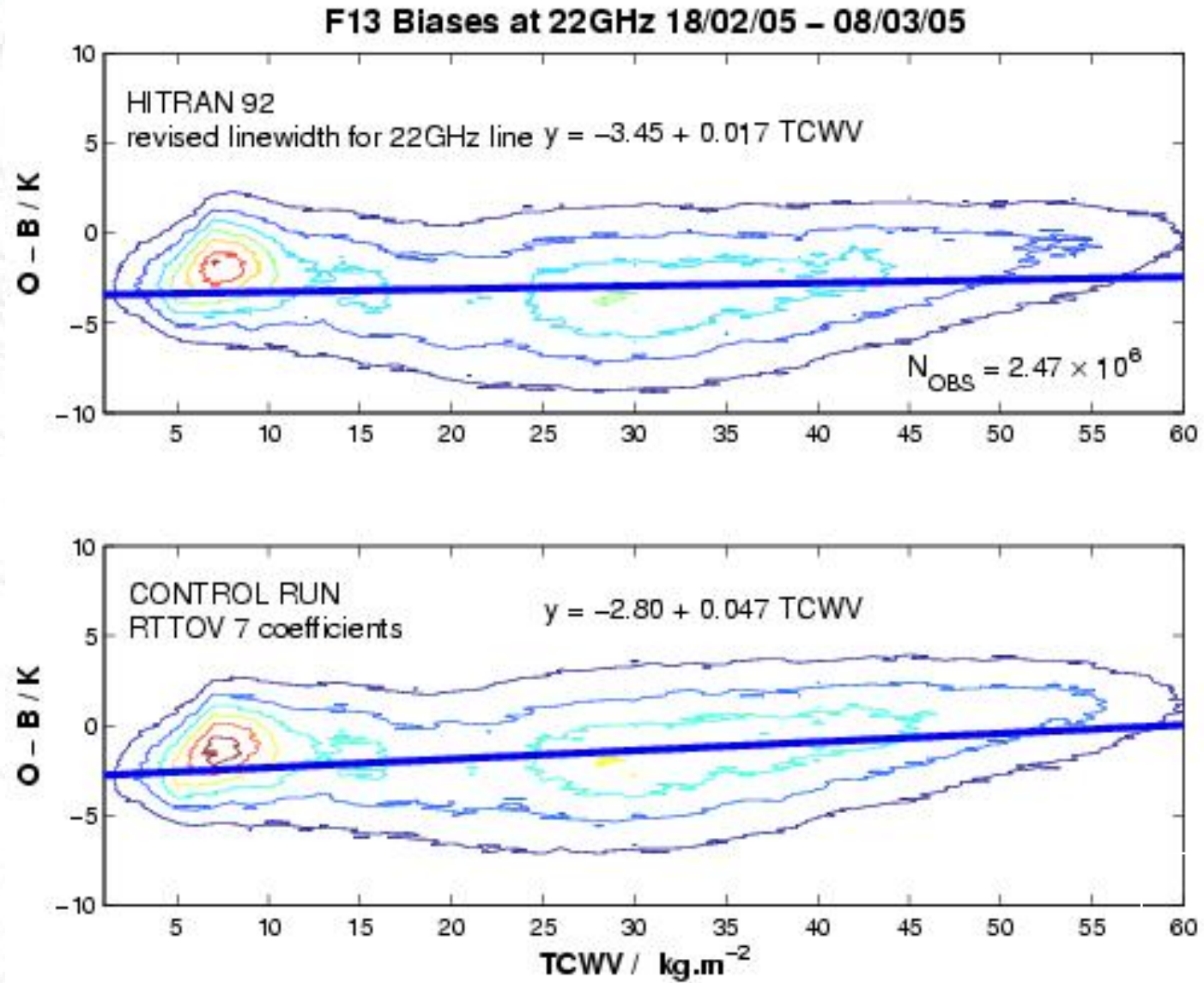
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21 days O-B (ECMWF) Mean Departures



(Saunders et al. 2005)

22.235 GHz line-width



LBL

Line inventory

HITRAN
MONORTM/LBLRTM
MPM89/92
Rosenkranz
ATM
STRANSAC
ARTS

Pressure/temperature dependence

Natural broadening (small)
Pressure broadening
Doppler broadening (at low pressures)

Continuum absorption (mainly H₂O)

Photoionization (IR)
Photodissociation (IR)
Far wings vs. H₂O clusters (MW windows)

- line intensities agree within 1%
- line frequencies accurate to within 0.1 kHz
- line widths/shifts modelling-measurements agree within a few %
- H₂O continuum controversial, modelling-measurements agree within 10-20 %

All the above is function of molecule and frequency!

Parameterizations

Profile datasets

Representativeness

Predictors

Temperature, pressure, gas concentration

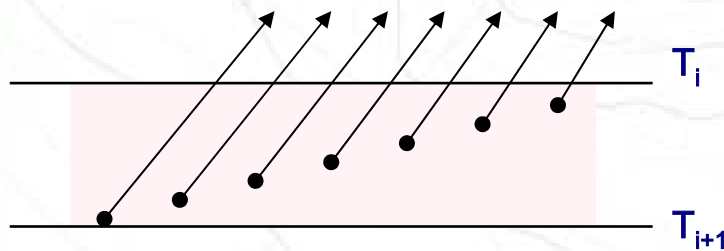
Integrated layer emission



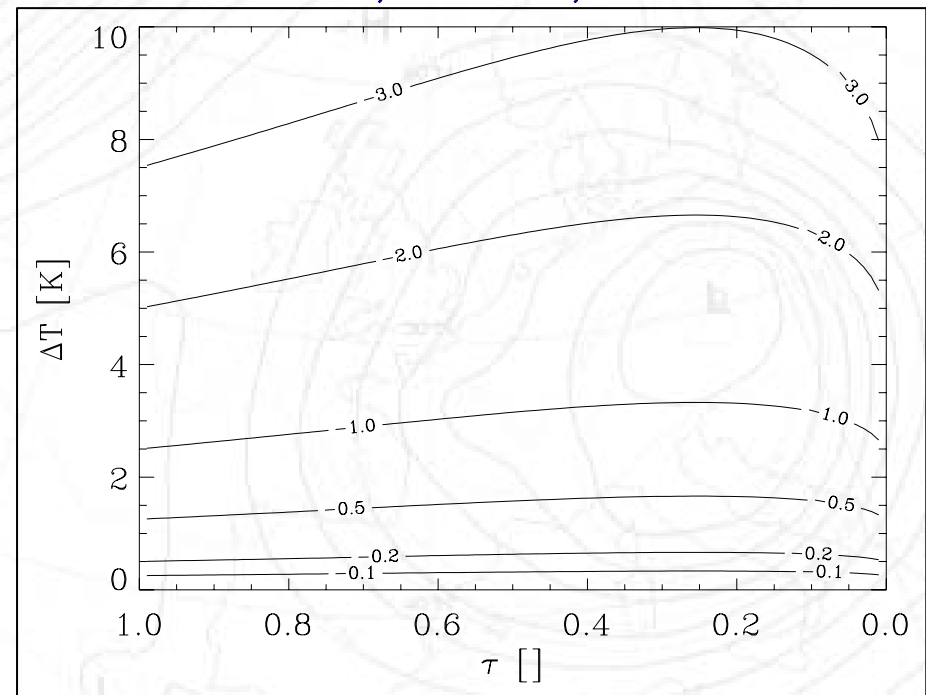
$$L_{\downarrow} = \int_0^{\Delta\delta} B[T(\delta)] \exp[-(\Delta\delta - \delta)/\mu] d\delta, \quad L_{\uparrow} = \int_{\Delta\delta}^0 B[T(\delta)] \exp[-\delta/\mu] d\delta$$

RTTOV: $L_{\uparrow} = L_{\downarrow} = \frac{B(T_i) + B(T_{i+1})}{2} [1 - \tau], \quad \tau = \exp(-\Delta\delta/\mu)$

Exact:
$$\begin{cases} L_{\uparrow} = B(T_i) - B(T_{i+1})\tau + B'[1 - \tau], & B' = \frac{B(T_{i+1}) - B(T_i)}{\Delta\delta} \\ L_{\downarrow} = B(T_{i+1}) - B(T_i)\tau - B'[1 - \tau] \end{cases}$$



$L_{\uparrow,RTTOV} - L_{\uparrow,Exact}$

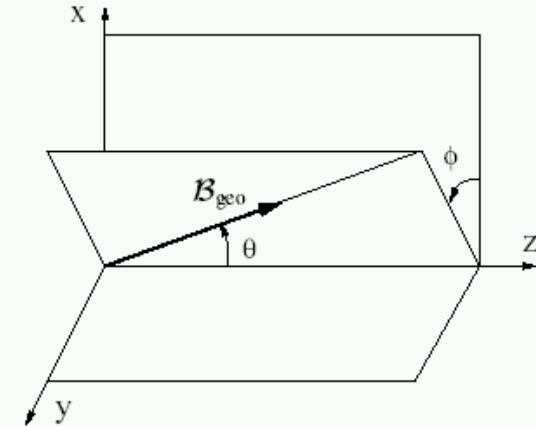


Zeeman splitting



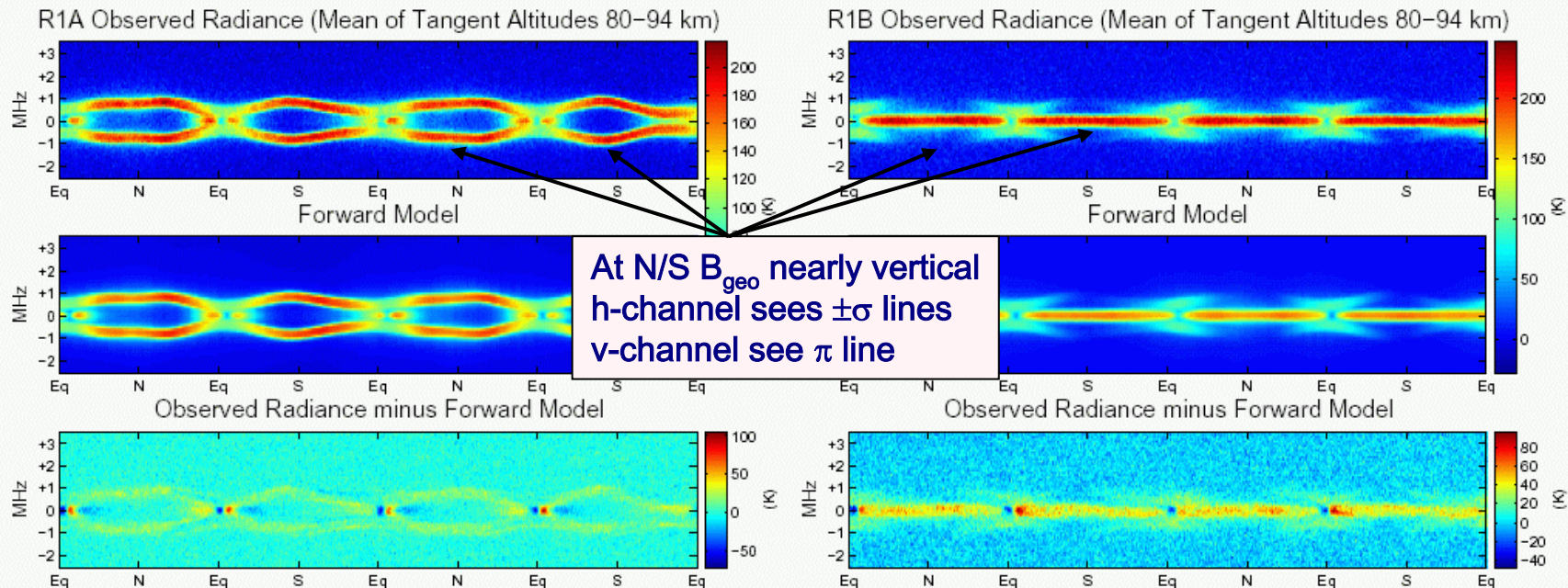
Splitting of O₂ lines in 60 GHz band and 118.75 GHz line through interaction of O₂ electronic spin (with magnetic dipole moment) with Earth's magnetic field:

- e.g. 118.75 GHz line has 3 components ($\Delta\nu \sim 1$ MHz)
- RT becomes polarization dependent
- RT becomes dependent on magnetic field orientation
- SSMIS, MLS



MLS Channel H_h at tangent point

MLS Channel H_v at tangent point

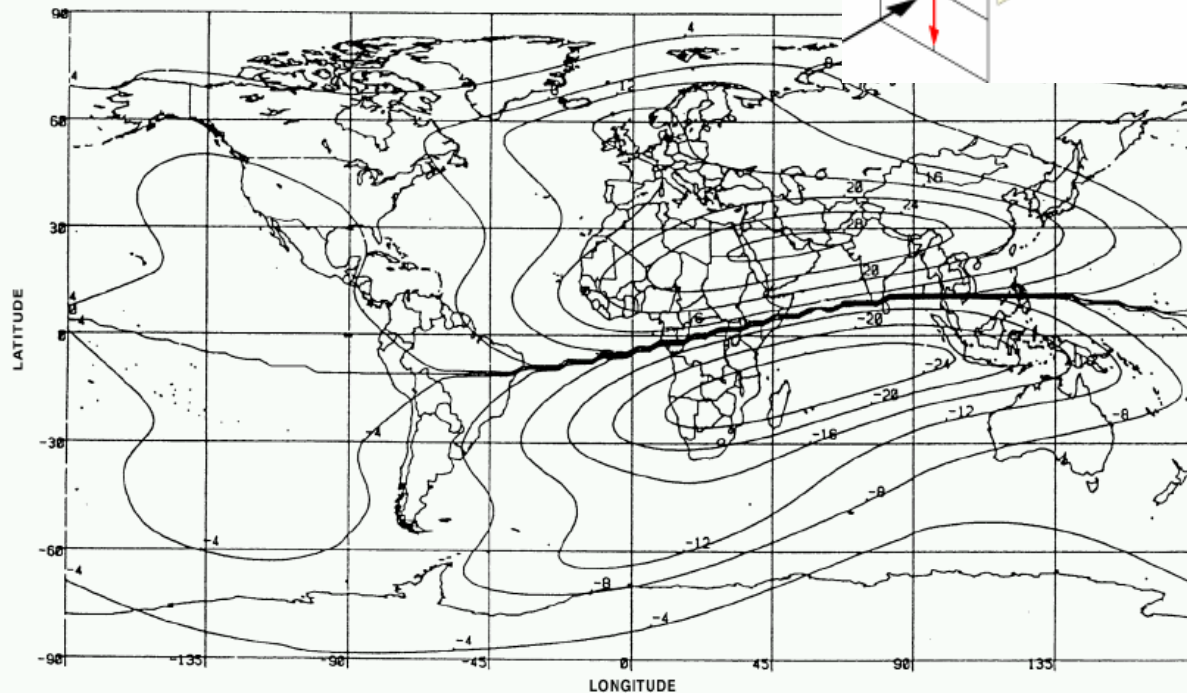
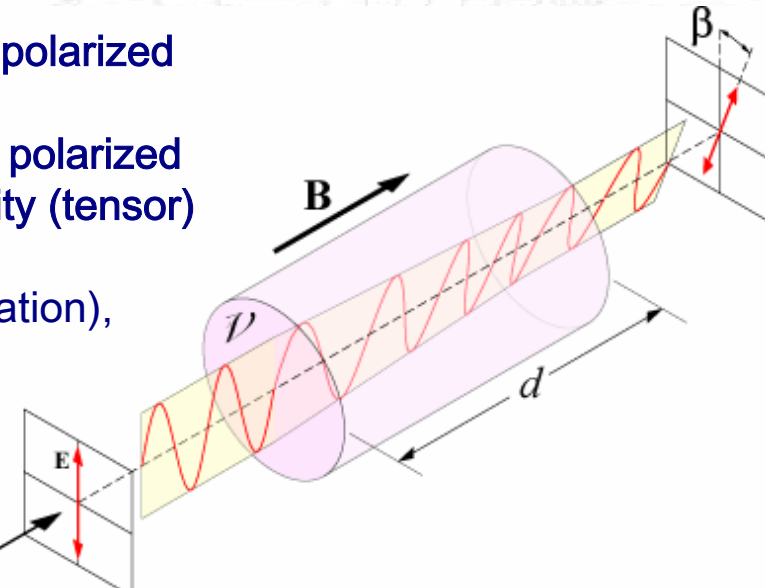


(Schwartz et al. 2005)

Faraday Rotation

Rotation of polarization through interaction of polarized light passing through strong magnetic field (in ionosphere). Splitting of wave into 2 circularly polarized rays due to polarization dependent permeability (tensor) causing phase delay.

- affects polarized light (surface sensitive radiation), also 3rd Stokes vector
- $\beta \approx 17 / \nu^2$ (β in degrees, ν in GHz)
- SMOS, AMSR, Windsat



Rotation angle simulation for 1.4 GHz (Svedlind 1986)

Refractive index $n = \sqrt{\epsilon\mu} = n' + in'' = f(\text{material, frequency, temperature})$

Water permittivity models are based on Debye model + fits to observational datasets:

- 2 datasets (3-20, 30-100 GHz)
- no sea-water data above 105 GHz
- 1 dataset for 9.62 GHz for super-cooled water (-18°C)
- fits required for $T \in [250-300 \text{ K}]$ and $\nu \in [1, 1000 \text{ GHz}]$

Ice permittivity models are mainly based on empirical fits to observational data:

- Real part rather constant
- Imaginary part very uncertain

Snow/ice permittivity:

- From mixing formulae based on air/water/ice and inclusion shape/orientation

Vegetation permittivity:

- Mainly function of water content but complex organic structure limits applicability of conventional mixing theory
- Experimental (field) observations available but extrapolation to satellite scale and wide frequency range difficult

Soil permittivity:

- Function of frequency, temperature, and salinity, volumetric water content, volume fraction of bound and free water related to the specific soil surface area, soil bulk material, and shape of the water inclusions
- Experimental (field) observations available but extrapolation to satellite scale and wide frequency range difficult

(Mätzler et al. 2005)

Surface emissivity - Oceans



Plane surface:

Sea-water permittivity
Fresnel equations (I, Q)

Wind roughened surface:

Sea-water permittivity
Fresnel equations (I, Q)

Large-scale waves

Gravity-capillary, capillary waves (> 2m/s)

Whitecaps (> 7 m/s)

Foam (> 10-12 m/s)

Directional wind roughened surface:

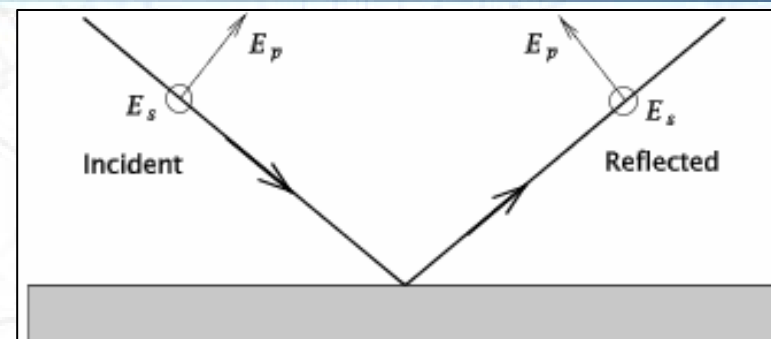
Sea-water permittivity
Fresnel equations (I, Q, **U**, **V**)

Large-scale waves

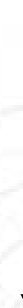
Gravity-capillary, capillary waves (> 2m/s)

Whitecaps (> 7 m/s)

Foam (> 10-12 m/s)



RTTOV FASTEM-2



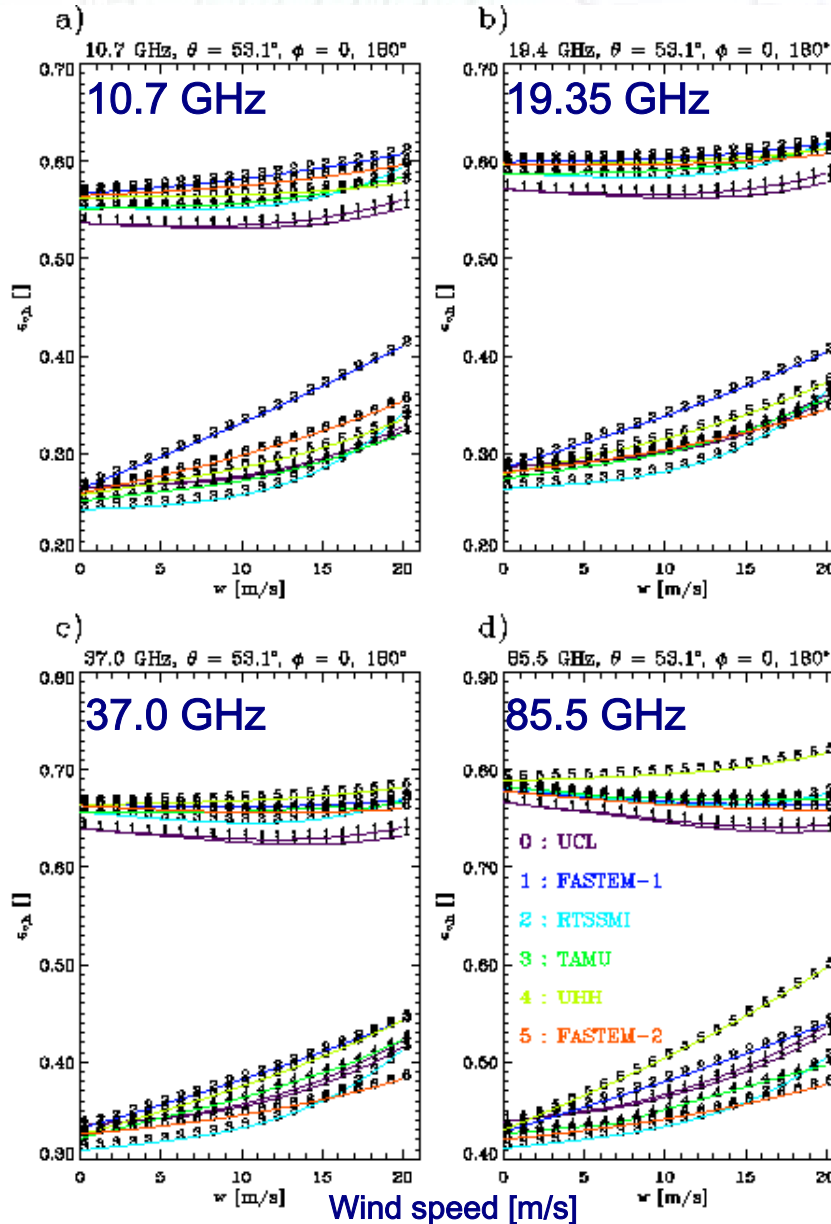
RTTOV FASTEM-3

$$\mathbf{L} = \begin{pmatrix} I \\ Q \\ U \\ V \end{pmatrix} = \frac{1}{2\sqrt{\mu/\varepsilon}} \begin{pmatrix} \langle E_v E_v^* \rangle + \langle E_h E_h^* \rangle \\ \langle E_v E_v^* \rangle - \langle E_h E_h^* \rangle \\ 2\text{Re} \langle E_v E_h^* \rangle \\ 2\text{Im} \langle E_v E_h^* \rangle \end{pmatrix} \approx \begin{pmatrix} I_o + I_1 \cos \varphi + I_2 \cos 2\varphi + \dots \\ Q_o + Q_1 \cos \varphi + Q_2 \cos 2\varphi + \dots \\ U_1 \sin \varphi + U_2 \sin 2\varphi + \dots \\ V_1 \sin \varphi + U_2 \sin 2\varphi + \dots \end{pmatrix}$$

Modelled emissivity - Oceans



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$\theta = 53.1^\circ$
 $\phi = 0, 180^\circ$

v-pol.

h-pol.

Sensitivity of surface emission ($E_0 T_s$) to real and imaginary part of dielectric constant (Meissner and Wentz 2004)

ν [GHz]	$\frac{\Delta(E_0 T_s)}{\Delta \text{Re}(\epsilon)} \Big _{T_s=273.15K}$ [K]	$\frac{\Delta(E_0 T_s)}{\Delta \text{Im}(\epsilon)} \Big _{T_s=273.15K}$ [K]
19.35	-0.056	+1.571
37.0	+0.300	-2.705
85.5	+0.082	+4.470

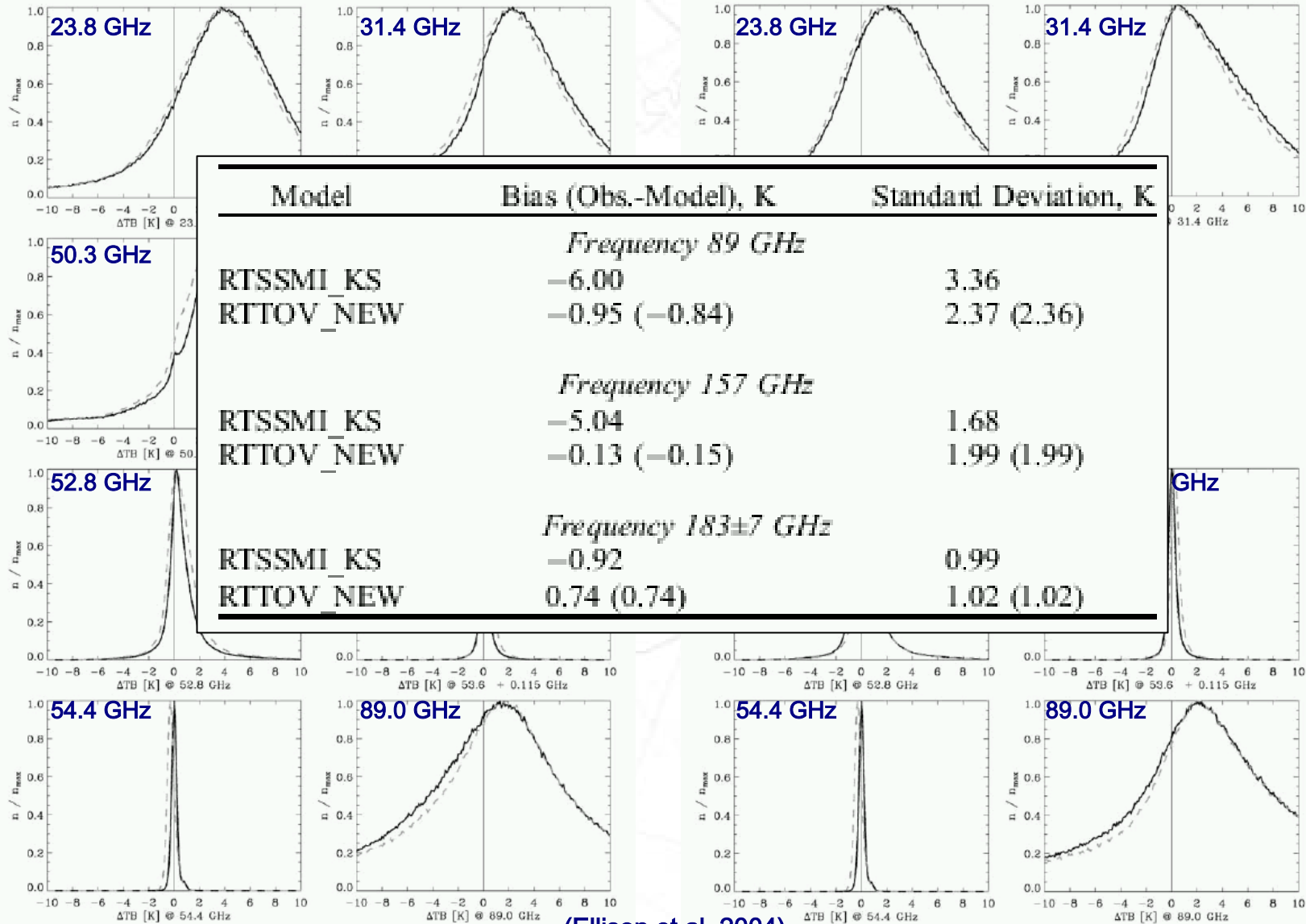
AMSU-A FG Departures - Oceans



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NOAA-16 AMSU-A before FASTEM-2

NOAA-16 AMSU-A after FASTEM-2

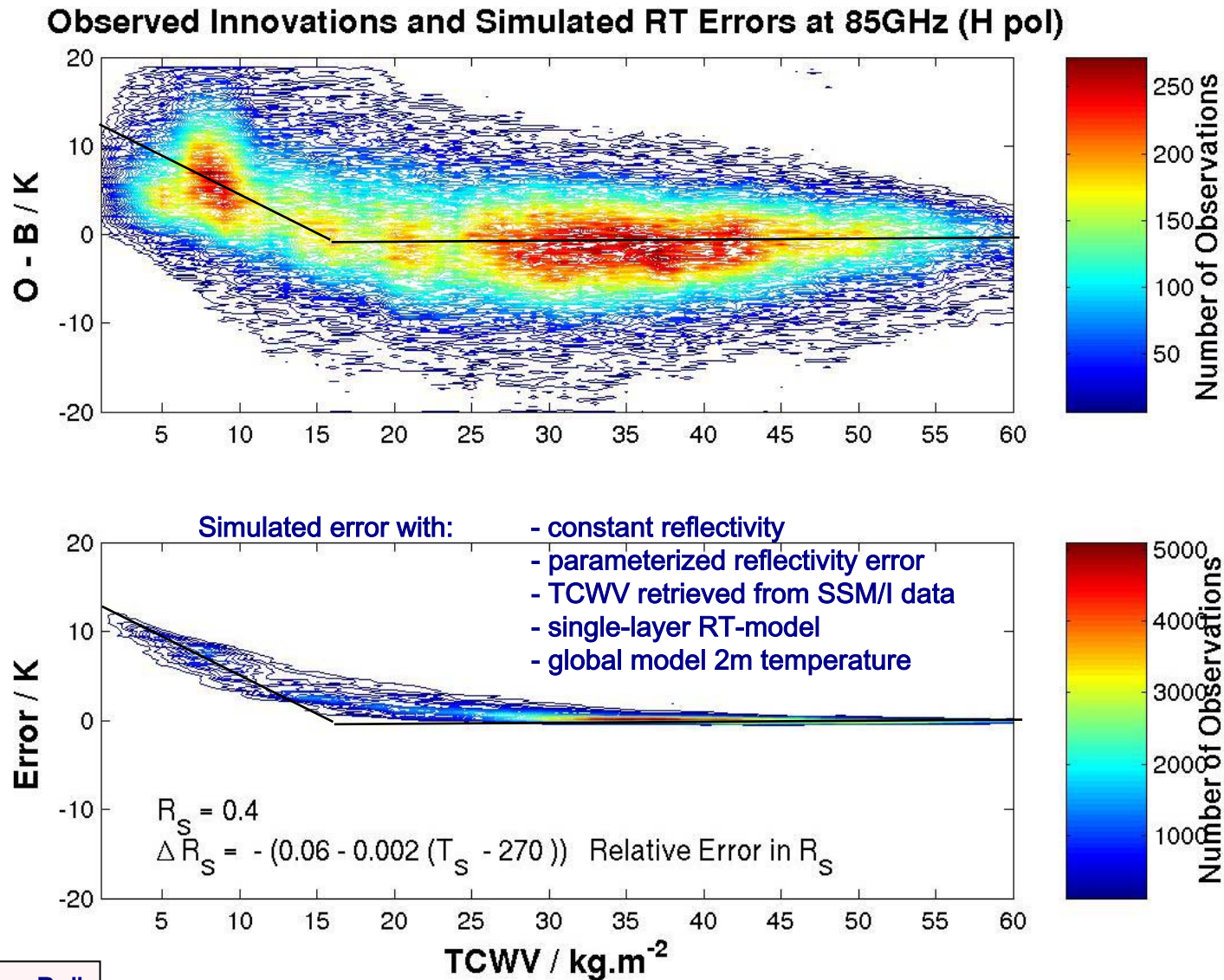


(Ellison et al. 2004)

SSM/I FG-Departure bias in dry/cold environments



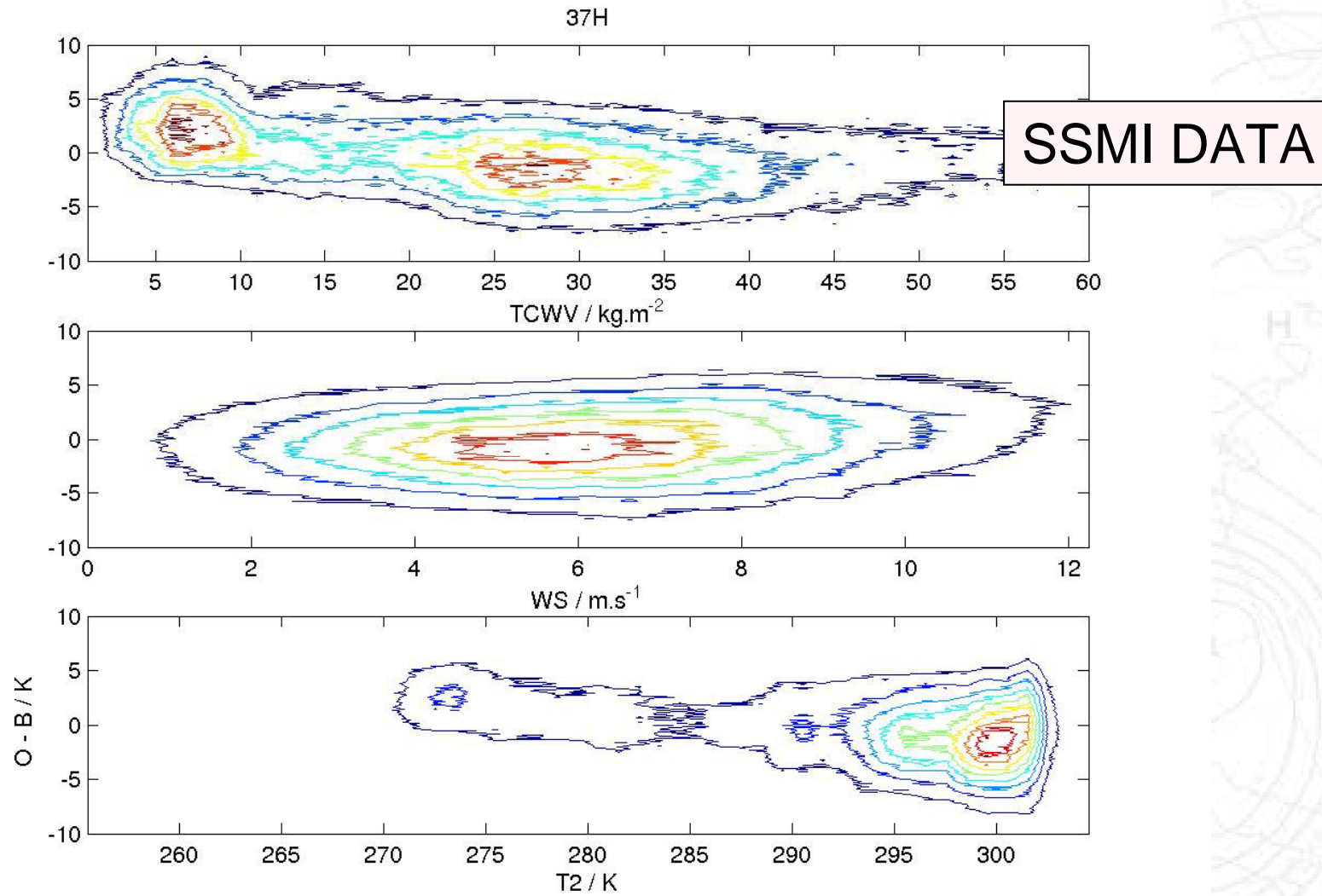
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William Bell
Met Office

Bias at low T: 37 GHz H pol

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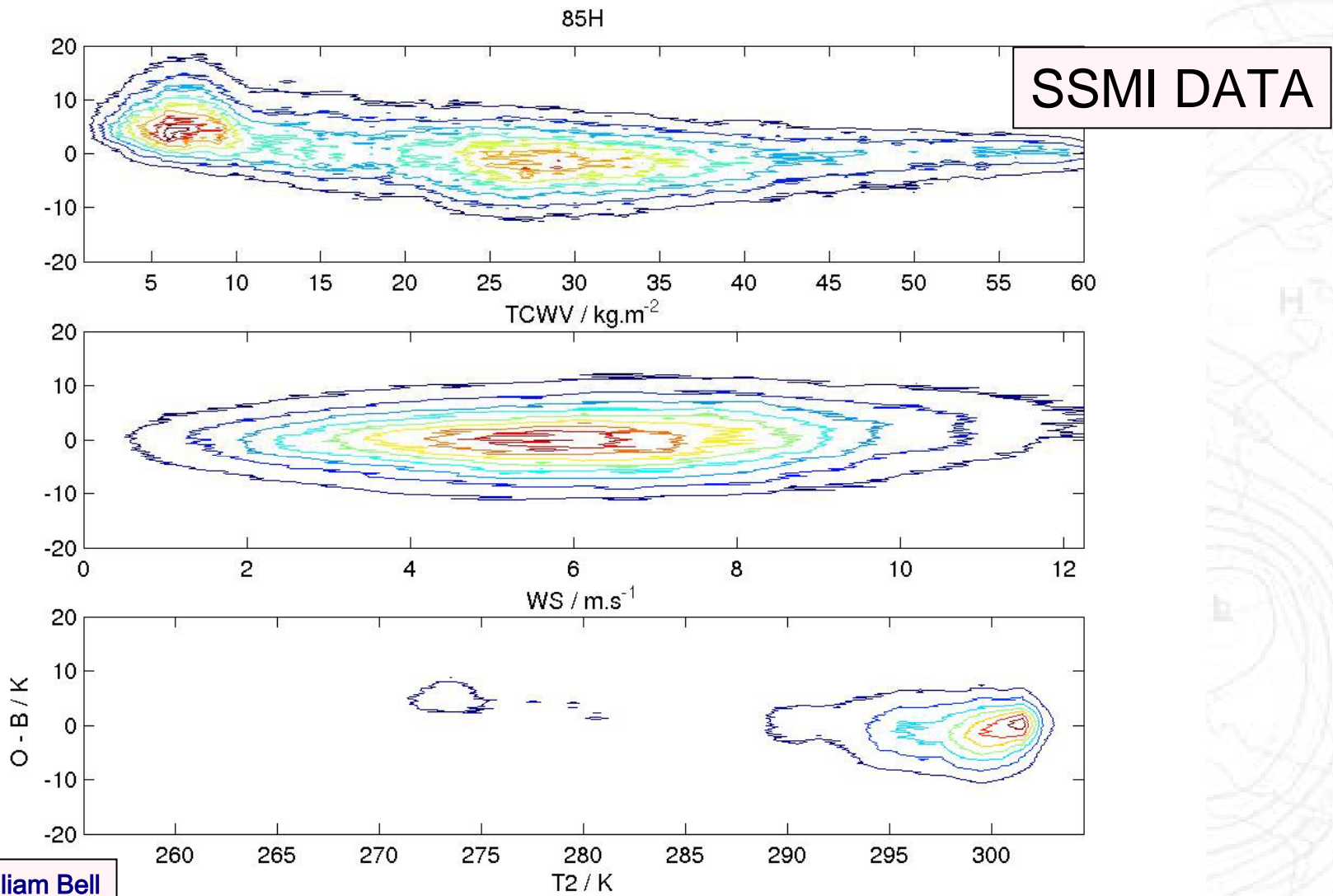


William Bell
 Met Office

Bias at low T: 85 GHz H pol



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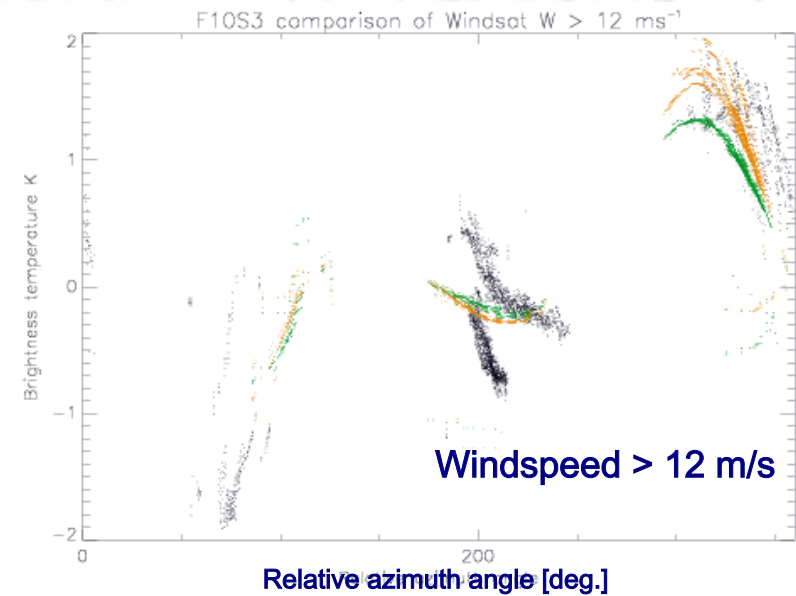
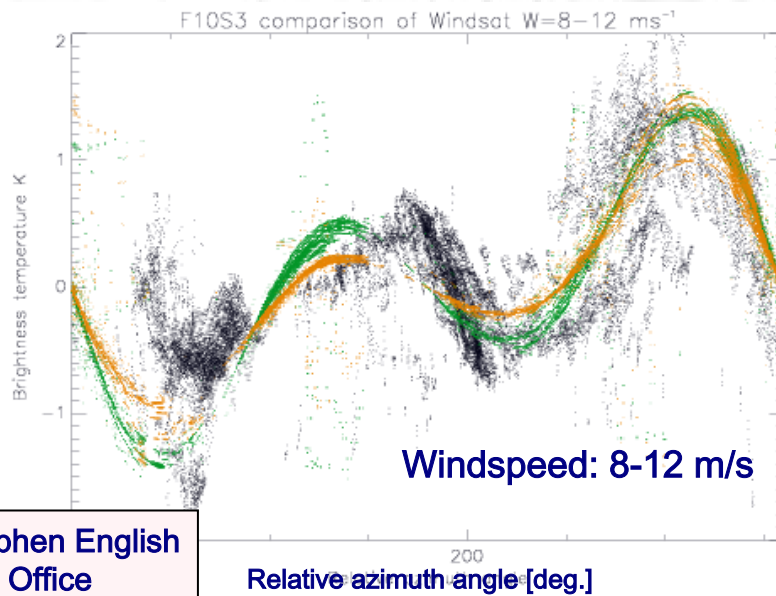
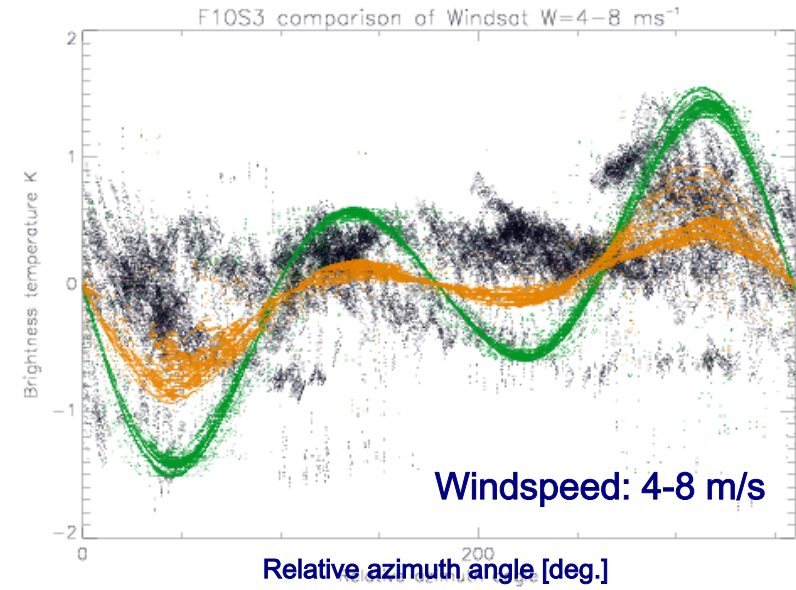
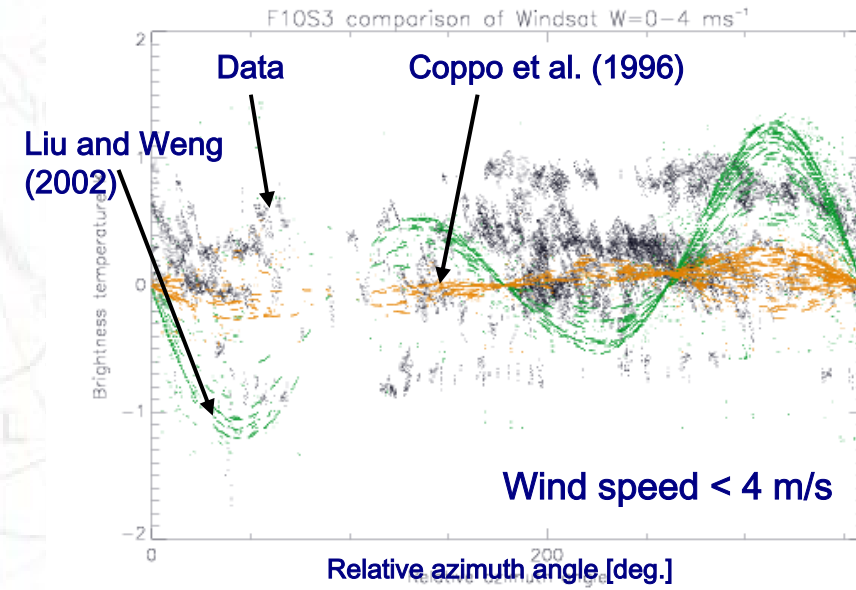


William Bell
Met Office

3rd Stokes Vector at 10.7 GHz: Models vs. Windsat Data



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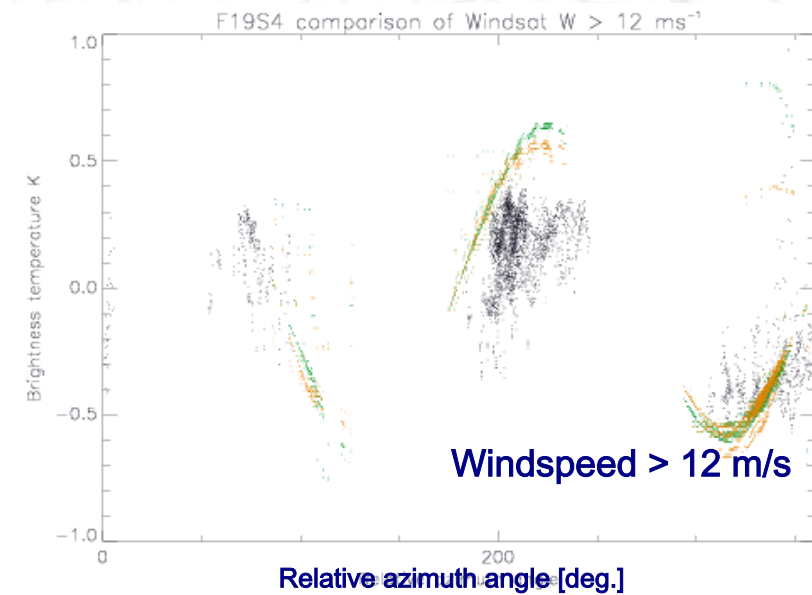
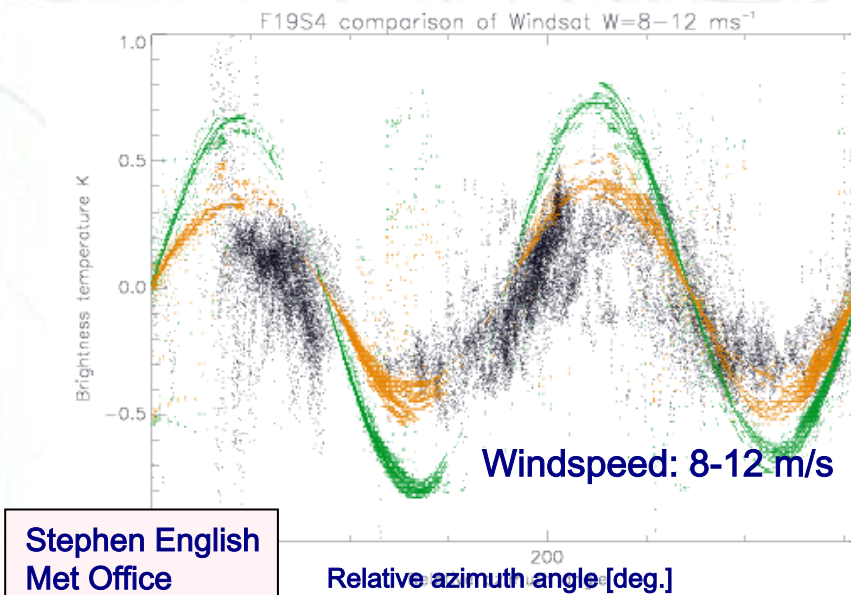
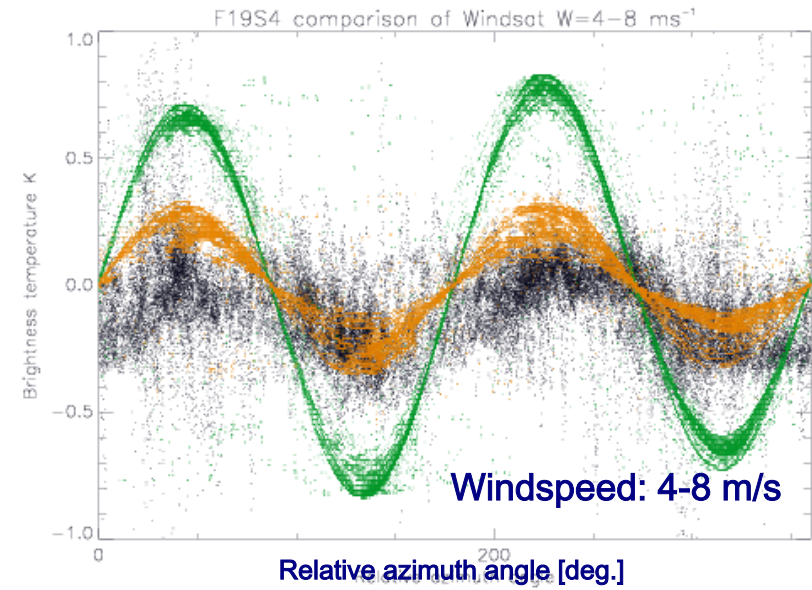
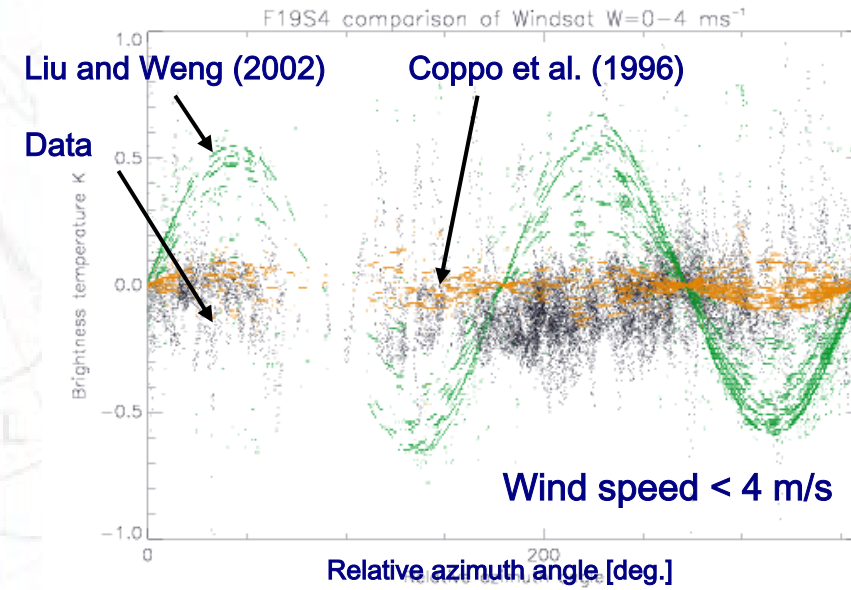


Stephen English
Met Office

4th Stokes Vector at 19.35 GHz: Models vs. Windsat Data



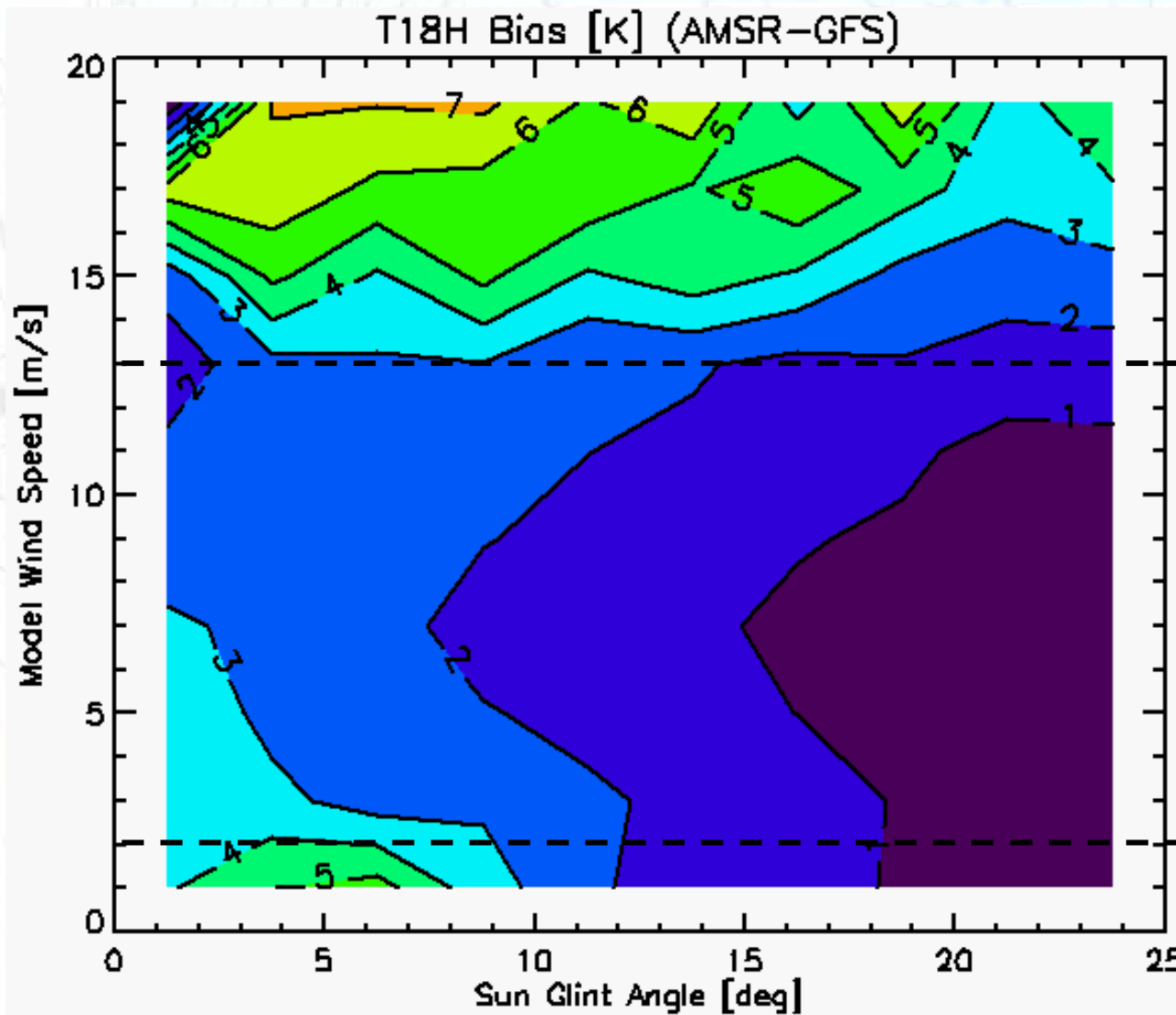
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Met Office

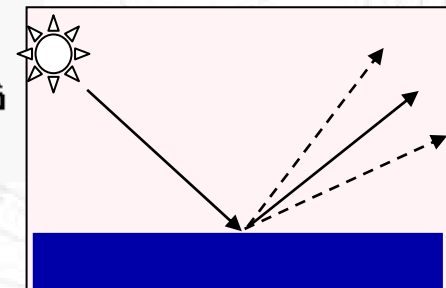
Sun-glint

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Sufficient
Data sample

Ralf Bennartz
University of Wisconsin

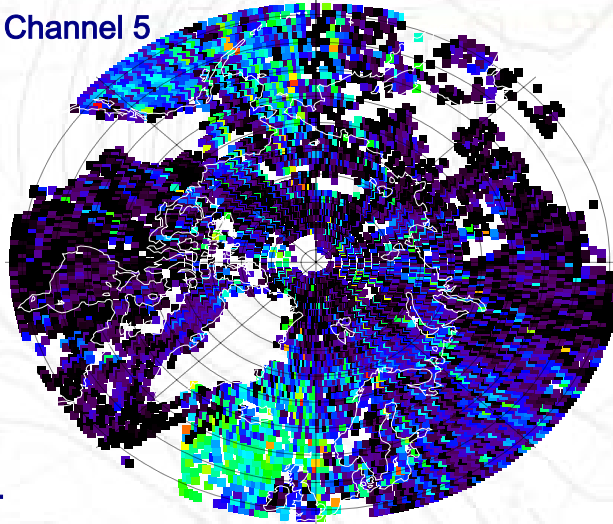


Surface emission - Land



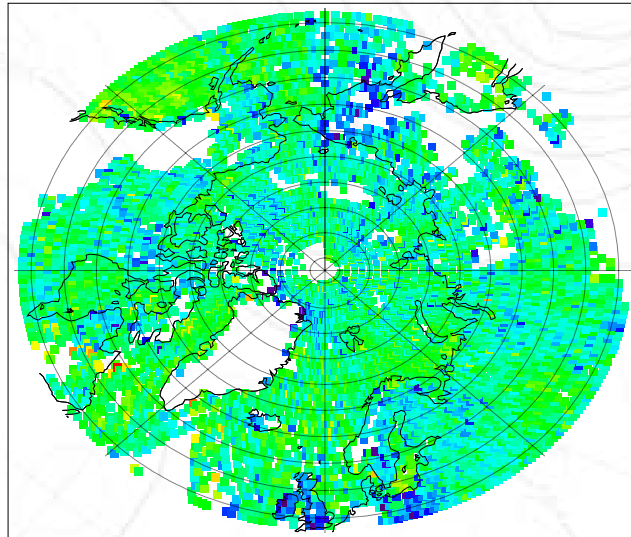
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AMSU-A Channel 5

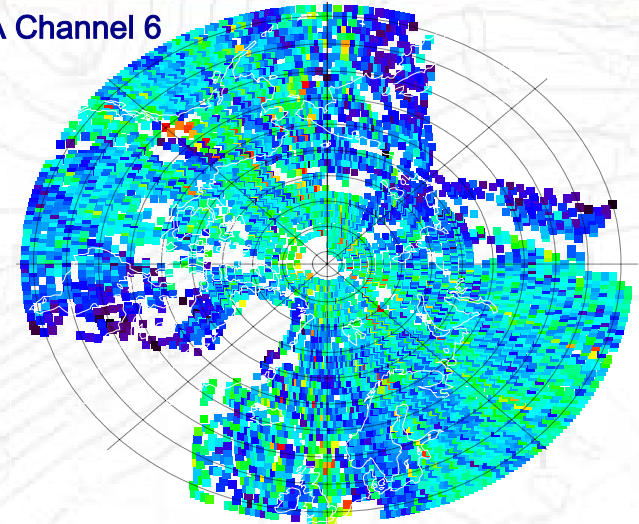


Before ...

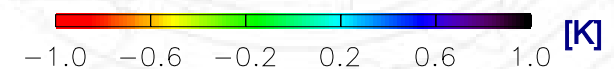
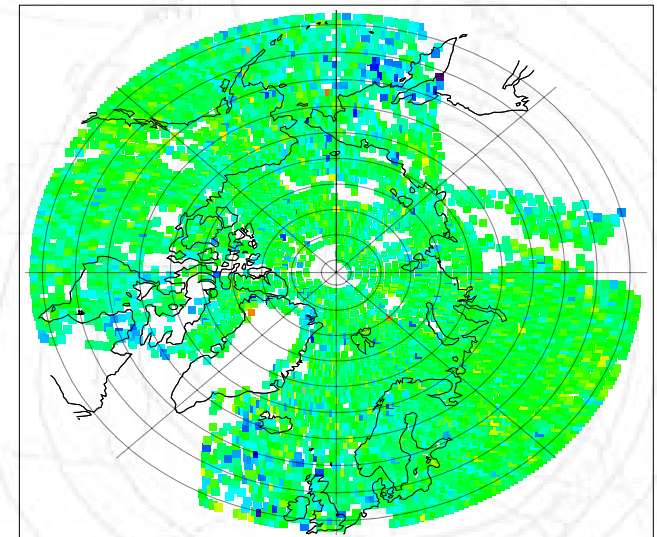
... after bias correction



AMSU-A Channel 6

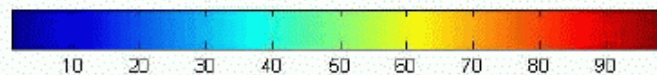
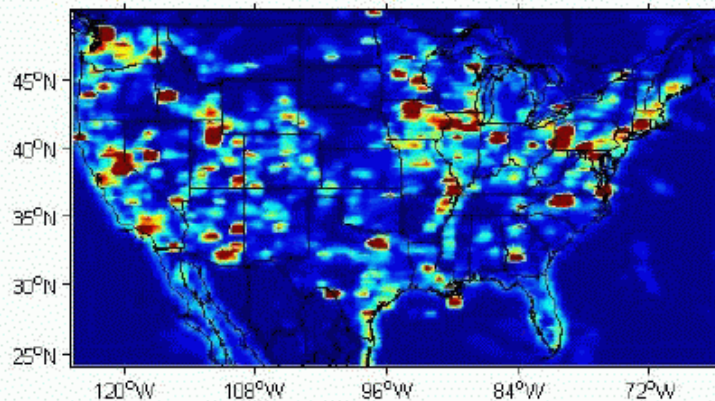


Used data 20050801-03

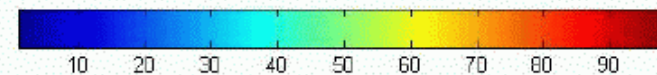
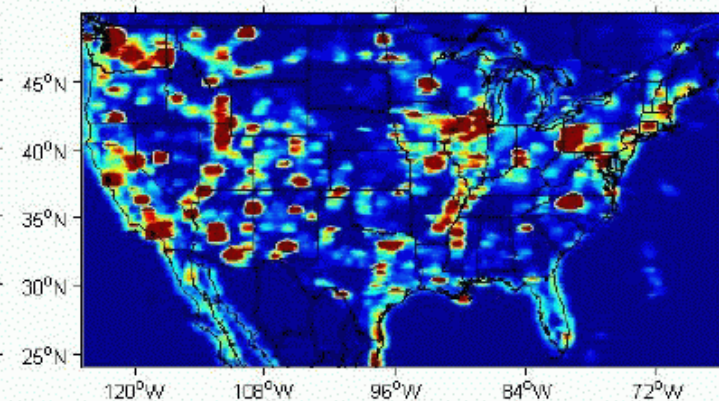


“Spectral Index” $T_{B, 6.8 \text{ GHz}} - T_{B, 10.7 \text{ GHz}}$

- This index is nominally $< 5 \text{ K}$ (typically, negative) for the geophysical signal
- Values $> 5\text{K}$ indicate RFI
- May be a more sensitive test for RFI than absolute T_B 's



V: $T_{B, 6.8 \text{ GHz}} - T_{B, 10.7 \text{ GHz}}$ [K]
Max Hold over 6 Months



H: $T_{B, 6.8 \text{ GHz}} - T_{B, 10.7 \text{ GHz}}$ [K]
Max Hold over 6 Months

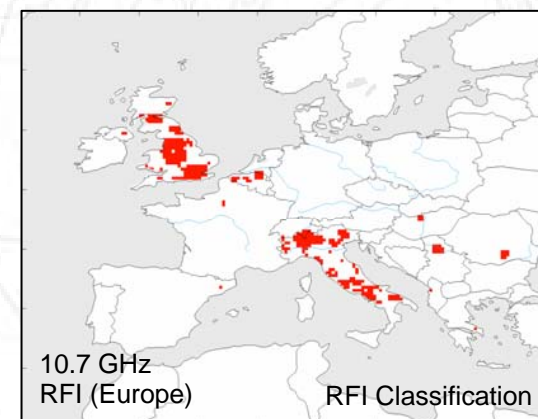
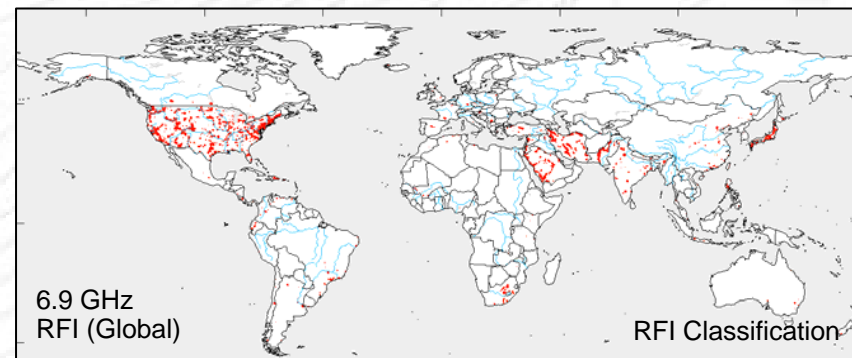
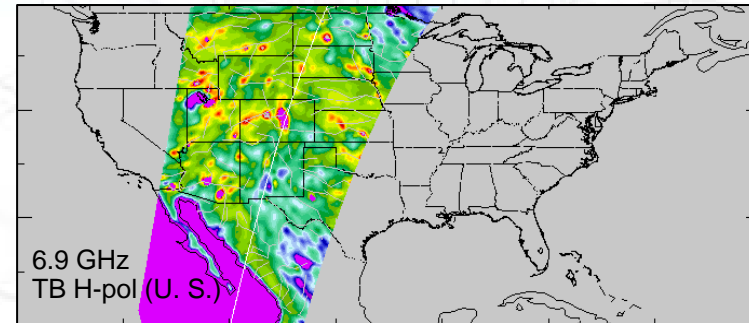
Surface emission - RFI



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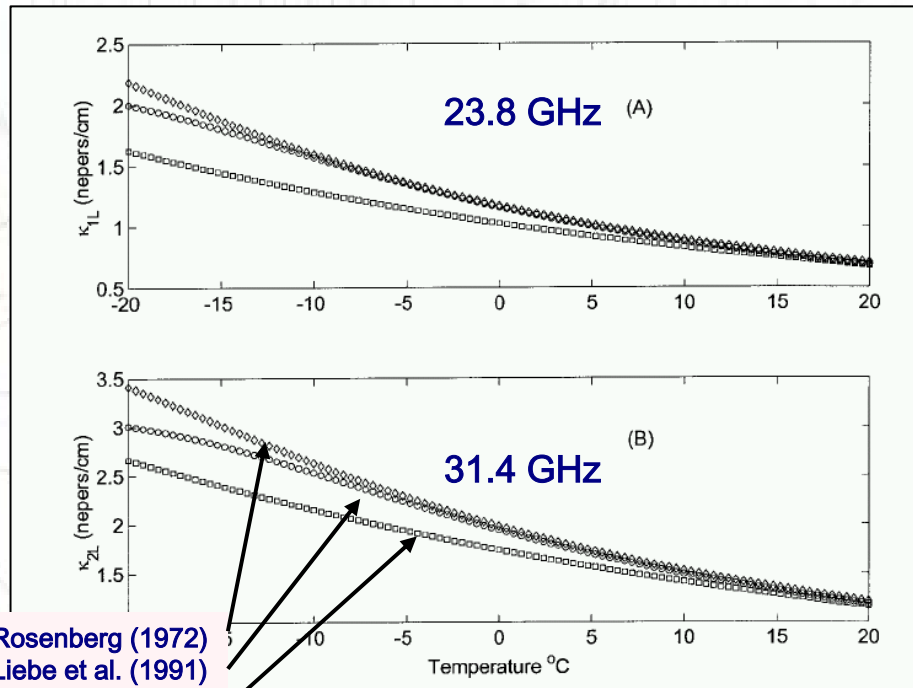
- RFI is observed in the AMSR C-band (6.75–7.1 GHz) and X-band (10.6–10.7 GHz) data
 - C-band is unprotected
 - X-band is protected from 10.68–10.7 GHz
- Classification algorithms can identify and filter strong RFI for AMSR-E geophysical algorithms
 - But, weak RFI cannot reliably be separated from geophysical signals
- C-band RFI mostly in the U. S., Japan, Middle East, some in Europe, Asia, S. America, Africa
- X-band RFI mostly in Japan, England, Italy, some in U. S.
- Situation at C-band has worsened considerably since Seasat and Nimbus-7 SMMR 1978-1987 (6.6 GHz)
- NPOESS/CMIS also operates at C- and X-bands
 - Re-assessment of radiometer design is in progress

(Njoku 2004)



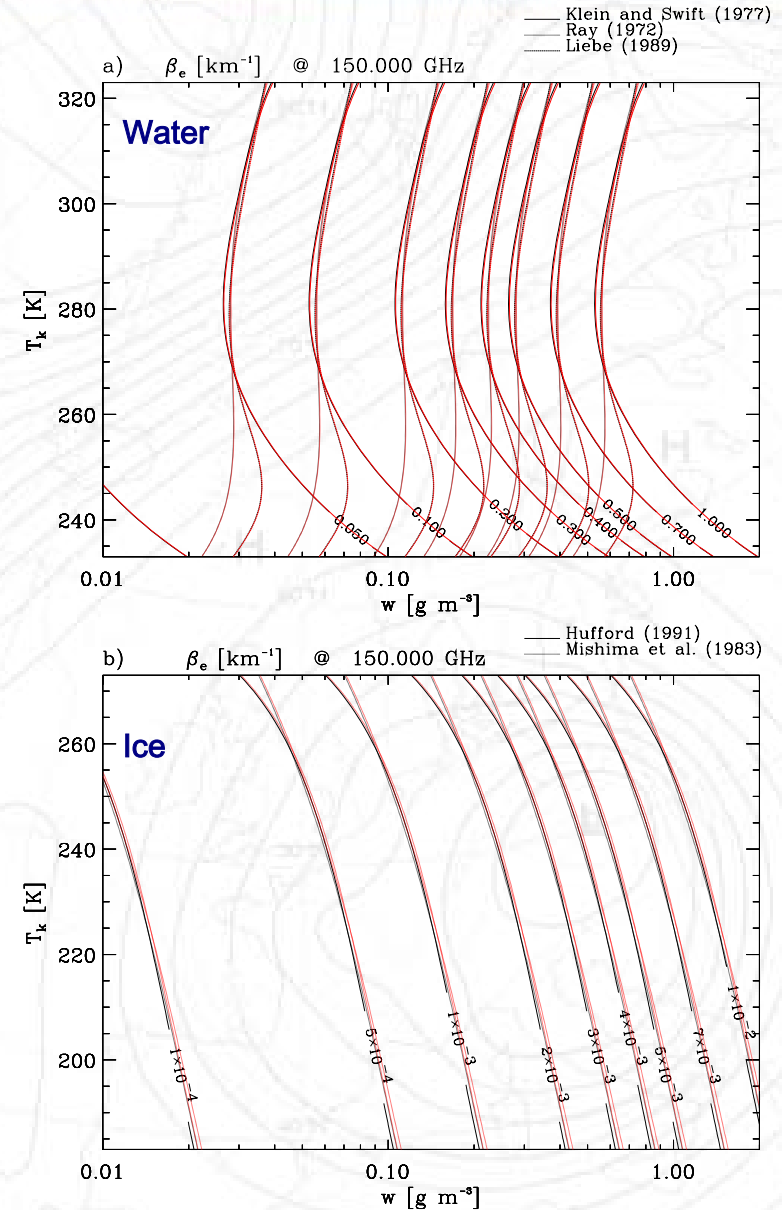
Cloud water/ice absorption

Cloud water mass absorption coefficient, κ ,
from various models
(Westwater et al. 2001)



Rosenberg (1972)
Liebe et al. (1991)
Grant et al. (1957)

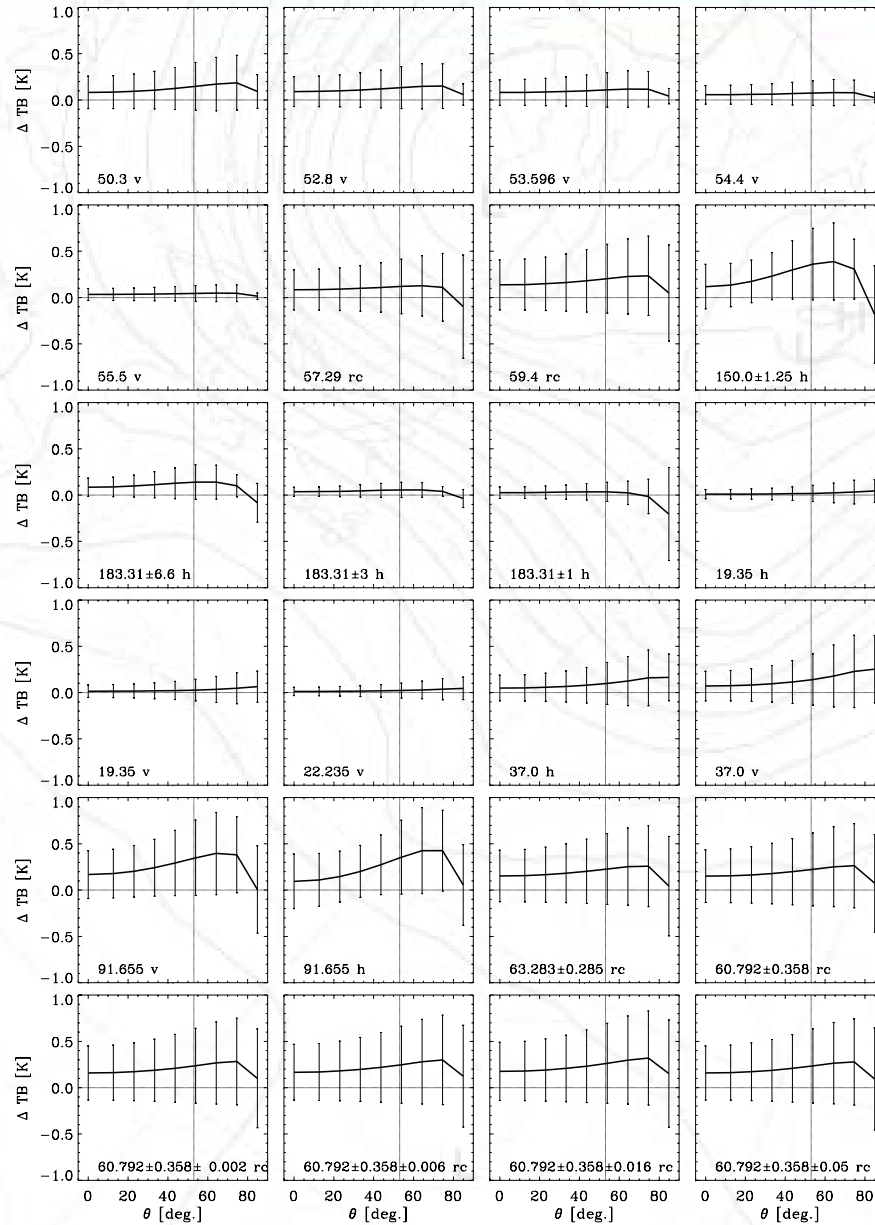
Cloud water/ice volume extinction coefficient, β ,
from various models and Rayleigh (black) and
Mie (red) calculations at 150 GHz



Multiple-scattering RT-Model comparison



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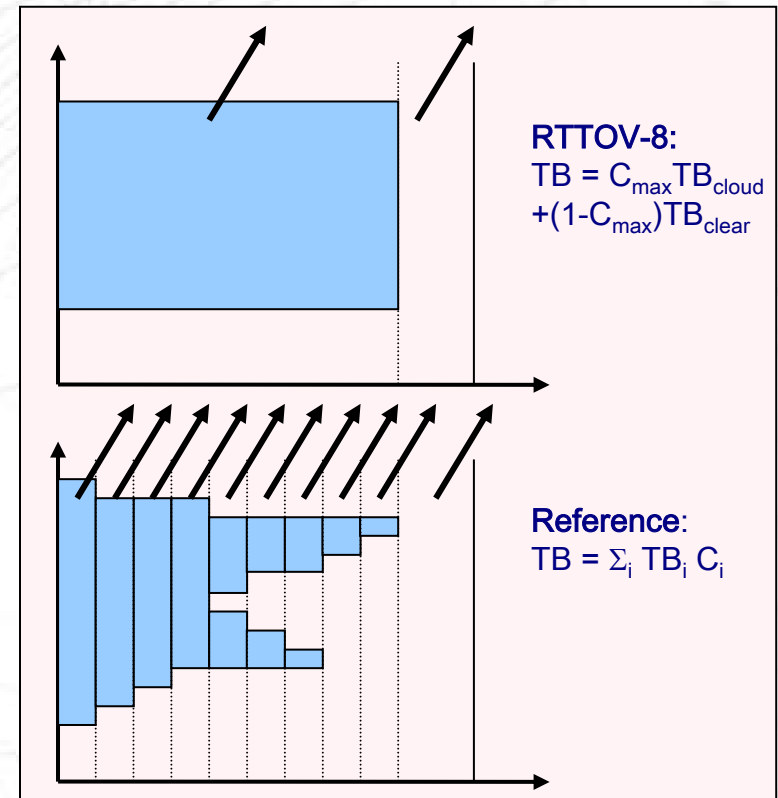
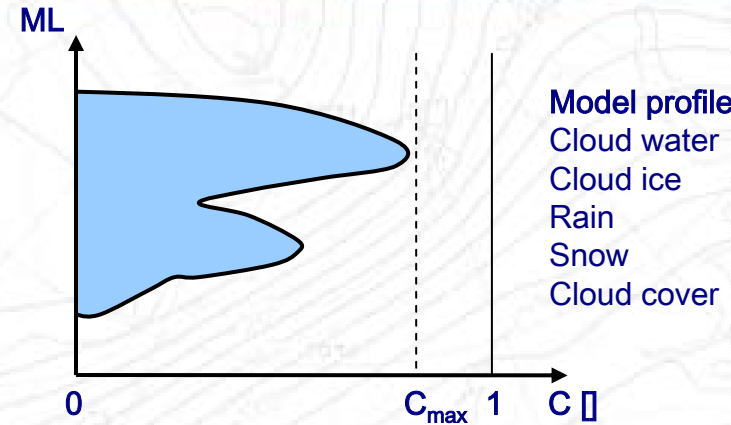
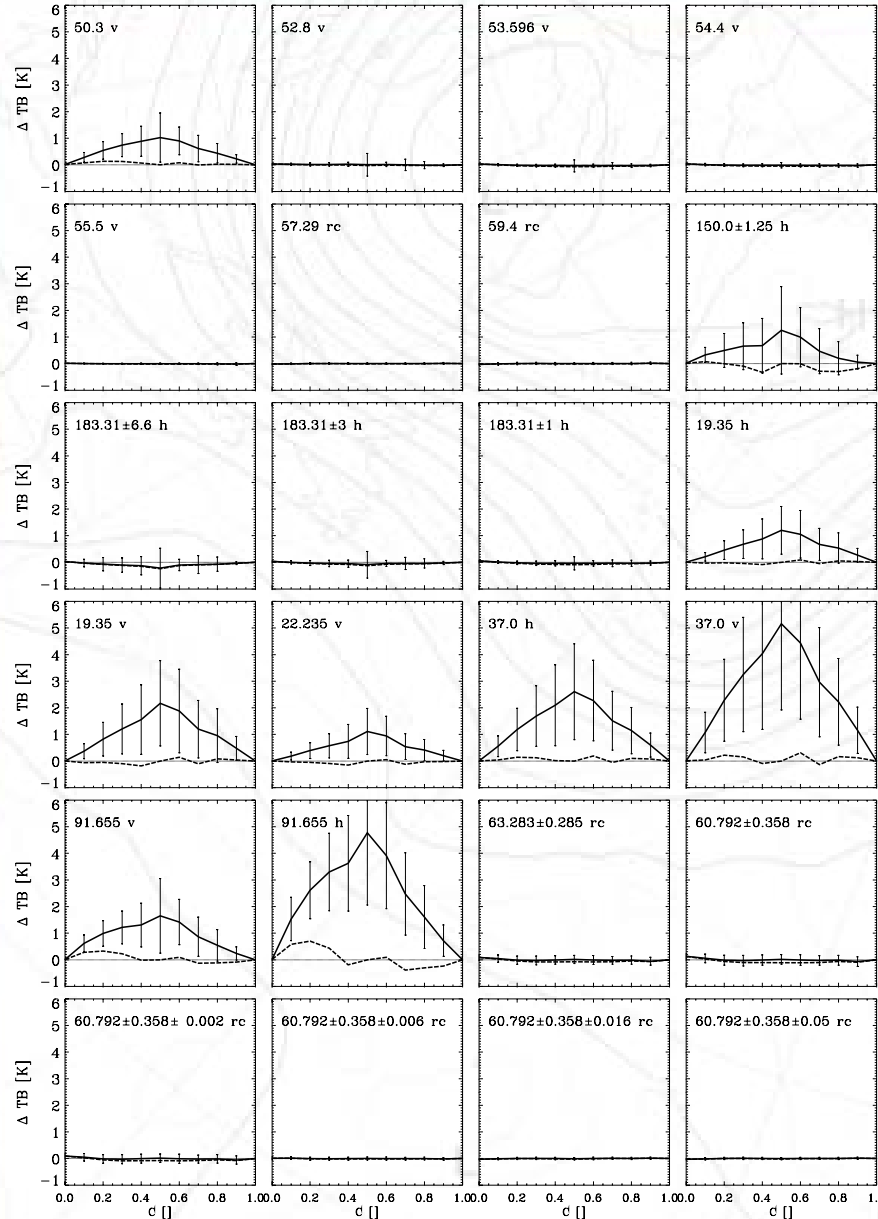


RT-biases < 0.5 K!

Effect of fractional cloud cover



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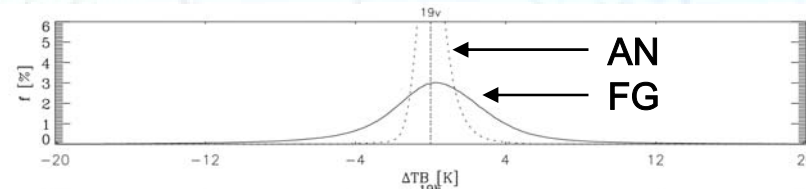


SSM/I TB FG/AN-Departures, September 2004

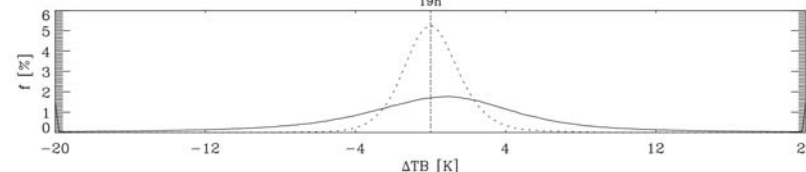


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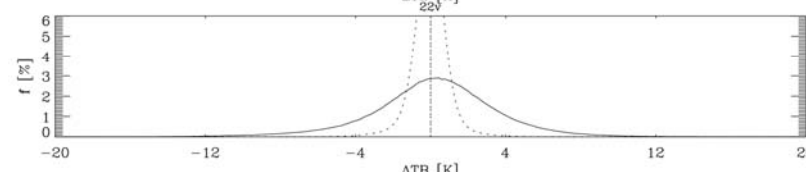
SSM/I channel:
1: 19.35 GHz (v)



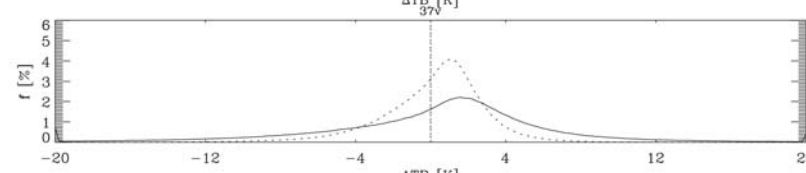
2: 19.35 GHz (h)



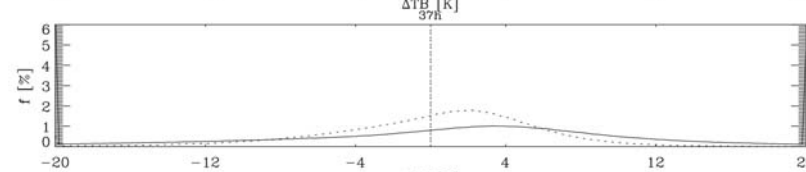
3: 22.235 GHz (v)



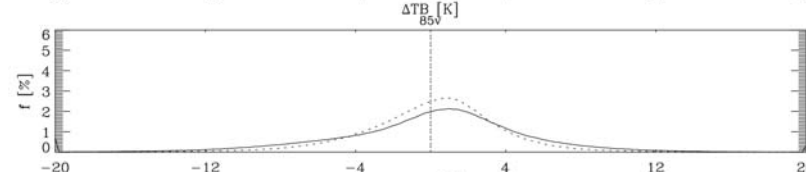
4: 37.0 GHz (v)



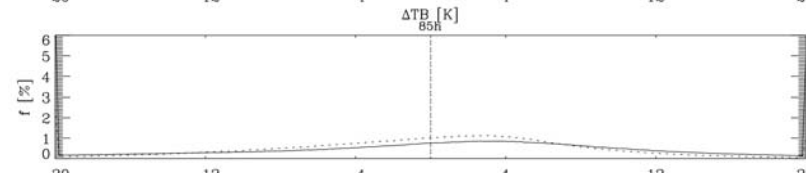
5: 37.0 GHz (h)



6: 85.5 GHz (v)



7: 85.5 GHz (h)



Bias corrected with
TCWV-predictor only

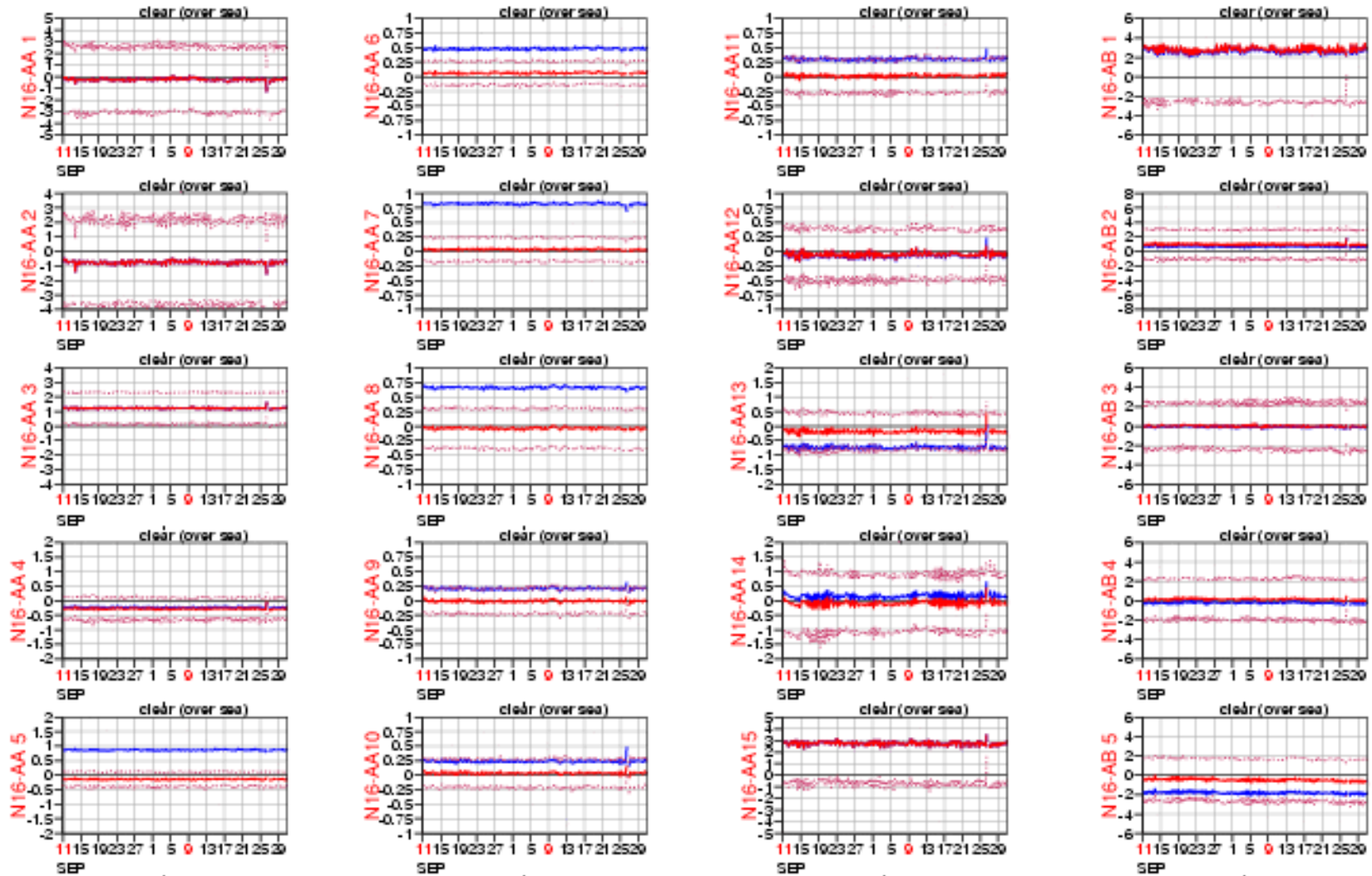
Higher order effects
(C-structure, PSD's etc.)
become important

Departure time series NOAA-16 AMSU-A/B (Clear-sky)



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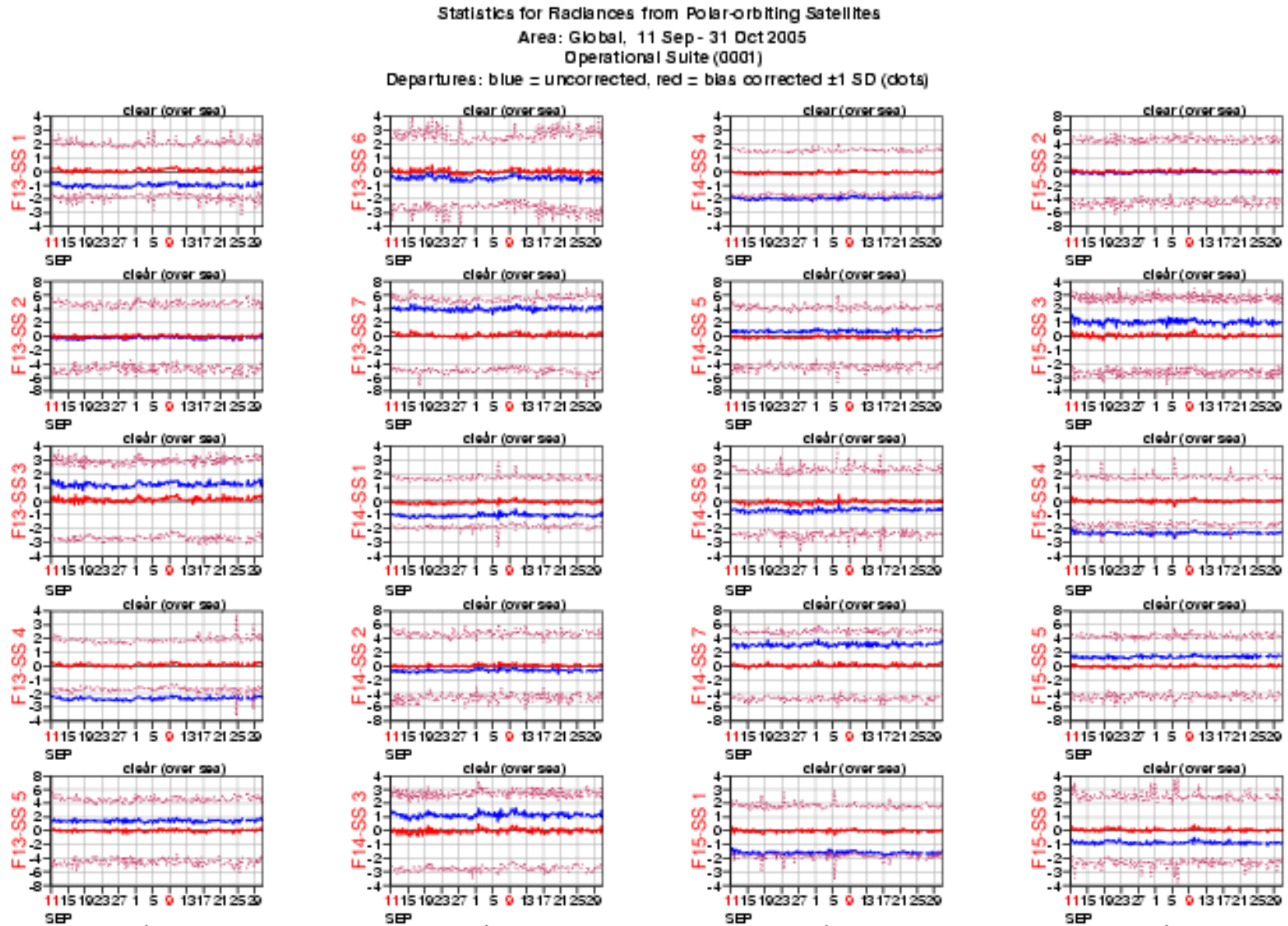
Statistics for Radiances from Polar-orbiting Satellites
Area: Global, 11 Sep- 31 Oct 2005
Operational Suite (0001)
Departures: blue = uncorrected, red = bias corrected ± 1 SD (dots)



Departure time series DMSP F-13/14/15 SSM/I (Clear-sky)



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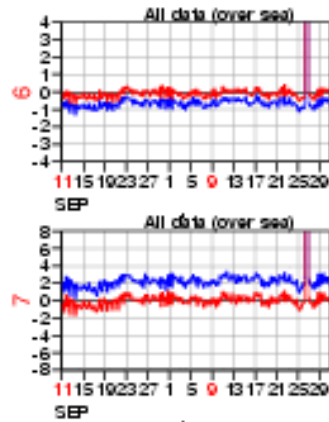
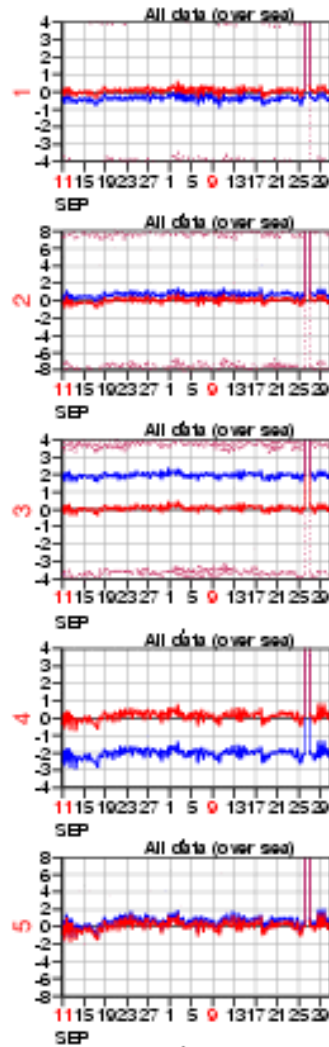


Departure time series DMSP F-13/14/15 SSM/I (RAIN)



ECMWF/NWP-SAF Workshop on bias estimation and correction in data assimilation
8-11 November 2005, Reading, UK

Statistics for Rain Affected Microwave Radiances from DMSP-11 / SSM/I
Area: Global, 11 Sep - 31 Oct 2005
Operational Suite (0001)
Departures: blue = uncorrected, red = bias corrected ± 1 SD (dots)



Summary



Contribution/Effect	Frequencies	Comments
Atmospheric absorption: Spectroscopy, LBL Parameterized models Zeeman splitting Faraday rotation	all all O ₂ lines 1.4 20-300 GHz	H ₂ O continuum problematic very accurate relative to LBL models limited applicability ($p < 10$ hPa), SSMIS, MLS limited applicability, SMOS Clear-sky <i>atmospheric</i> TB's accurate within 1-3%
Surface emission: Sea surface Permittivity Polarimetry Land surfaces Soil, vegetation Snow/ice Type, age, etc.	all 10-37 all All	1 K between 20-150 GHz 10% for $4 \text{ m/s} < \text{SWS} < 12 \text{ m/s}$ uncertain uncertain
Cloud droplet emission: Permittivity	all 5-500 GHz	well modelled between 5-500 GHz and $T > 273$ K for $T < 273$ K models differ by 20-30%
Precipitation emission/scattering: PSD, Permittivity, Shape	all	uncertain
Radiative transfer modelling: Clear-skies Clouds/precipitation multiple scattering Layer inhomogeneity	all all all	biases up to 3 K due to $T(z)$, otherwise accurate biases < 0.5 K biases up to 5 K, can be parameterized