



Scientific Challenges of UV-B Forecasting

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- International activities and the UV Index
- UV Index definition and forecasting requirements
- Challenges in calculation of the radiation transfer
 - effects of aerosols (relation to GEMS), of altitude, and of albedo
 - cloud modification factor
- Example, products, verification
- Conclusion



UNCED 1992, Agenda 21, and INTERSUN

Agenda 21: Undertake as a matter of urgency

- research on the effects on human health of increasing UV reaching the ground due to ozone depletion
- efforts to mitigate the effects on human beings

In response to Agenda 21, WHO and further organisations established **INTERSUN** the Global UV Project:



Objective: Primary prevention; Priority Activities:

- Filling gaps in scientific knowledge
- Assessing and quantifying health risks
- Promoting sun protection
- Promoting the UV Index



UV Index

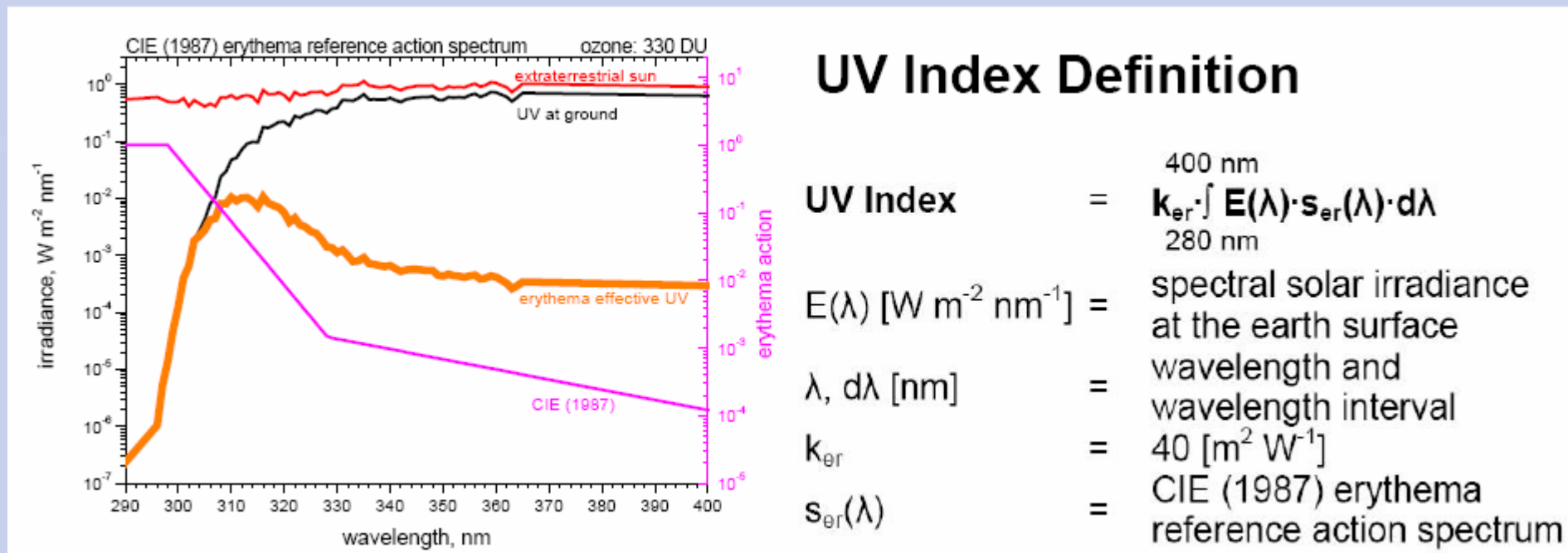
In 1994 introduced by WMO, WHO , ICNIRP, UNEP to

- describe the level of solar UV radiation reaching the ground by means of a simple physical quantity

Objective:

- raise public awareness of the risk of excessive exposure to UV and the need to adopt protective measures
- serve as an educational tool to promote sun protection

UV Index: Biologically weighted physical quantity



UV Index Definition

$$UV \text{ Index} = k_{er} \cdot \int_{280 \text{ nm}}^{400 \text{ nm}} E(\lambda) \cdot s_{er}(\lambda) \cdot d\lambda$$

$E(\lambda) [W m^{-2} nm^{-1}]$ = spectral solar irradiance at the earth surface
 $\lambda, d\lambda [nm]$ = wavelength and wavelength interval
 k_{er} = 40 [$m^2 W^{-1}$]
 $s_{er}(\lambda)$ = CIE (1987) erythema reference action spectrum

The UV Index is defined in reference to a horizontal surface



Requirements in UV Index forecasting and reporting

WMO

- to take into account cloud cover and other relevant environmental variables (surface albedo, aerosols)
- to report to the public the daily maximum value, *whenever it occurs*

COST-713 on “UV-B forecasting” recommends additionally:

- to forecast total ozone by global dynamical models, including chemistry effects and data assimilation techniques.
- to include aerosol optical depth variable in space and season.
- to model the UV Index by a multiple scattering algorithm.

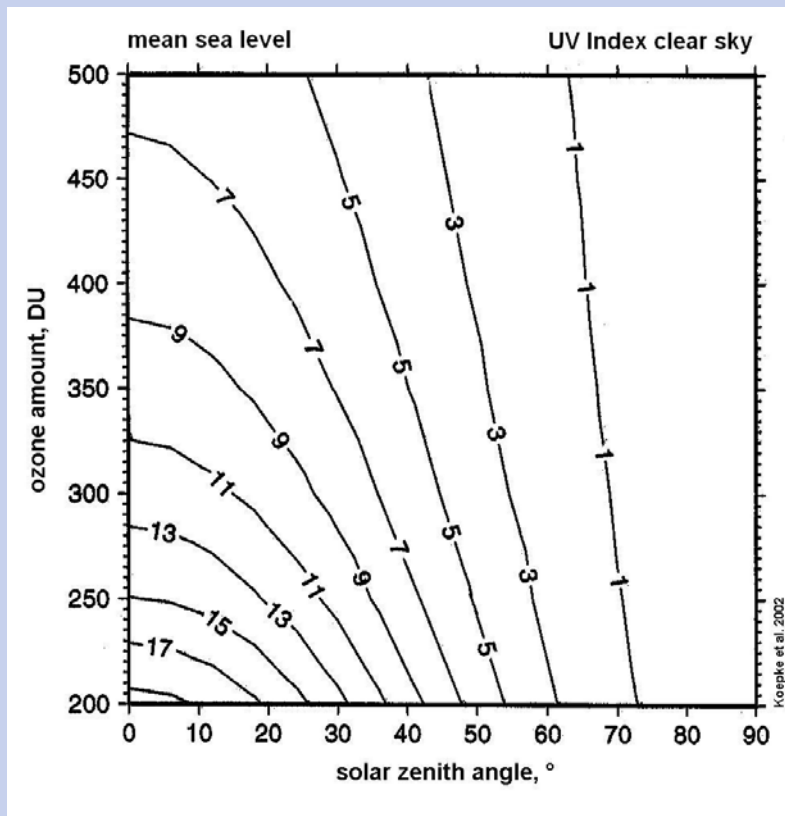


Challenges in UV Index forecasting

The challenges to accept in UV forecasting are of operational nature:

- forecasting not only for local noon
- fast calculation of radiation transfer accounting for
 - aerosol optical depth
 - aerosol absorption (single scattering albedo)
 - altitude effects
 - surface albedo
 - clouds

Radiation Transfer for clear sky conditions



Large Scale UV Index (ls_UVI)

Realised by Lookup Tables (LUT) for

- 12 months
- 5 zonal belts per hemisphere

Resolution 1° in solar zenith angle (SZA), 10 DU in total ozone (TOZ)

Range: 0 - 14 (mean sea level)

Vertical distribution of ozone and temperature: **effects < 3 %**



Adjusting for variable atmospheric conditions and altitude

$$\text{UVI} = I_s_UVI(\text{SZA}, \text{TOZ})$$

$$* f_{\text{aerosol}}(\text{AOD}, \text{SSA}, \text{SZA})$$

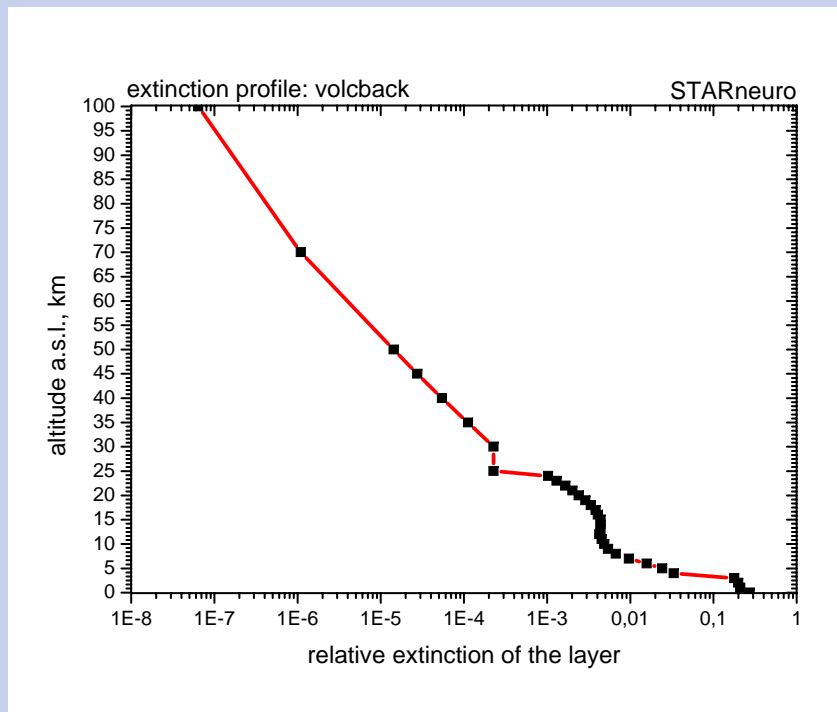
$$* f_{\text{altitude}}(\text{AOD}, \text{SSA}, \text{SZA}, z)$$

$$* f_{\text{albedo}}(\text{ALB}, \text{SZA}, z)$$

[UV Index clear sky]

$$* f_{\text{clouds}}(\text{cloud_cover}, \text{cloud_type}[, \text{sw-rad}])$$

Aerosol effects: $f_{\text{aerosol}}(\text{AOD}, \text{SSA}, \text{SZA})$



Aerosol effects modelled by

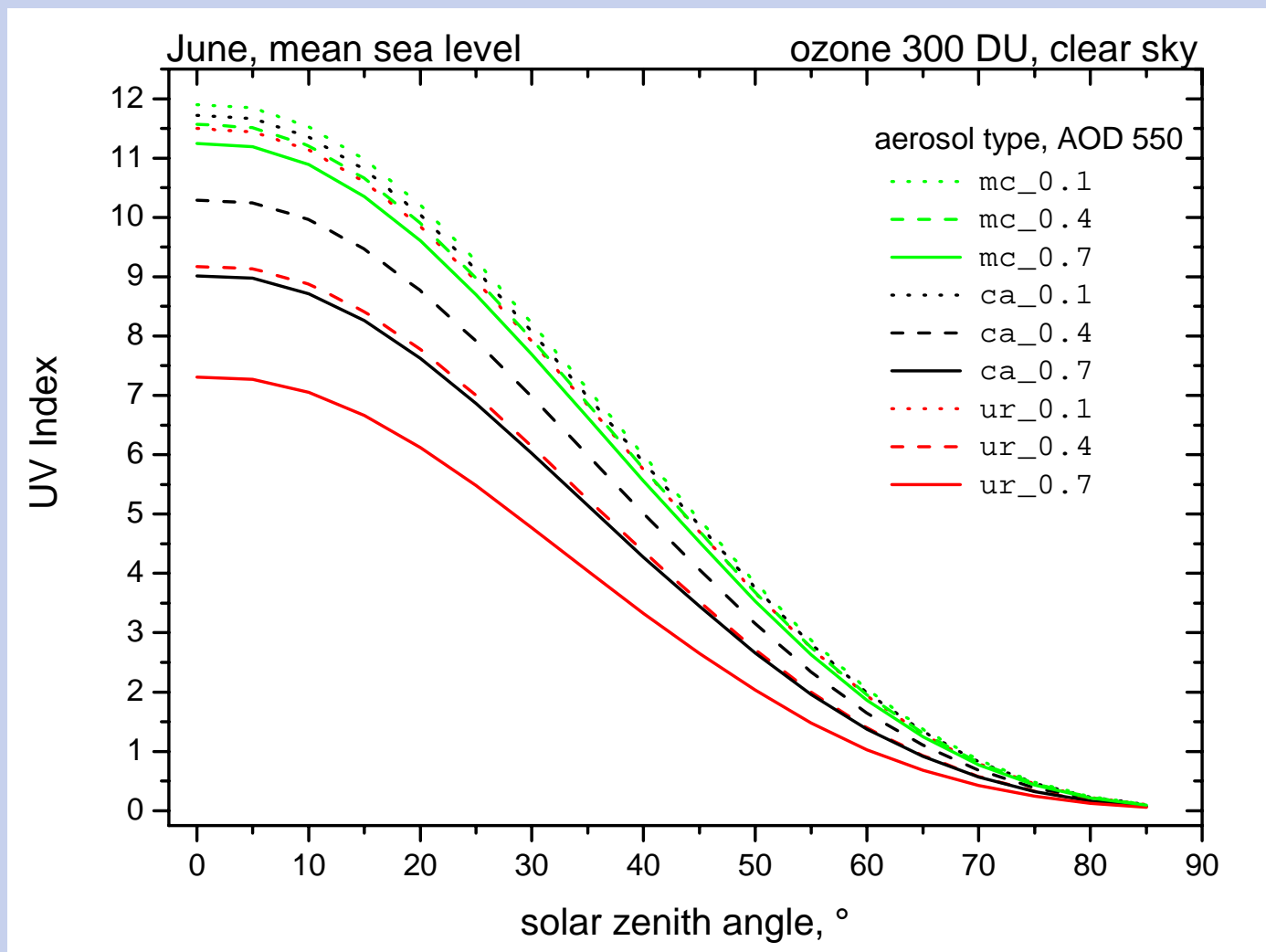
- set type: stratosphere (>12 km)
- set type: free troposphere (<12 km)
- variable types: (0-3 km)

Type of relative aerosol profile:

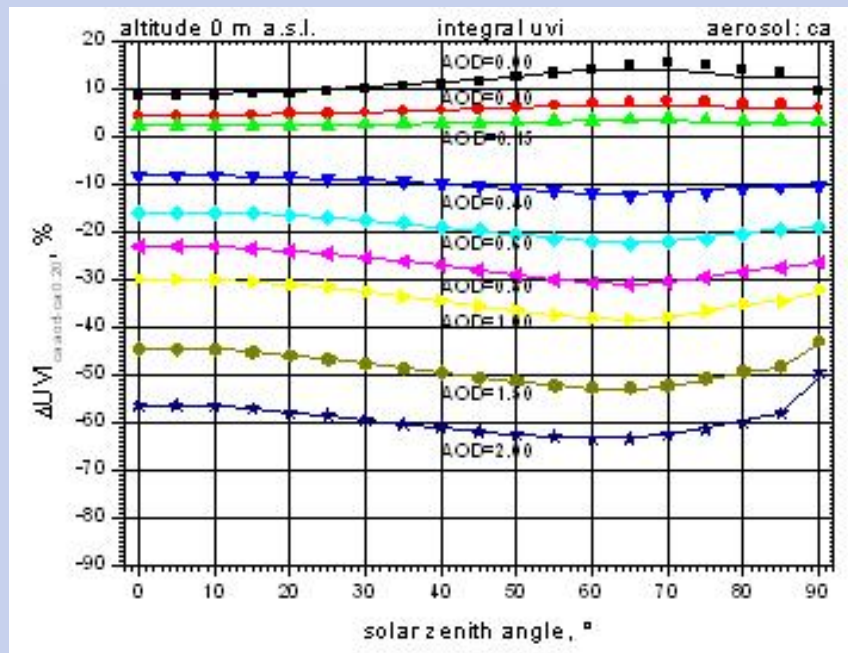
- volcanic background conditions

Phase function, asymmetry factor, and SSA are defined via the aerosol types. Effects of varying profiles <3%

Aerosol effects



Aerosol effects: $f_{\text{aerosol}}(\text{AOD}, \text{SSA}, \text{SZA})$



UVI, maximum abs. error: 0.05

Accounting for aerosol effects:

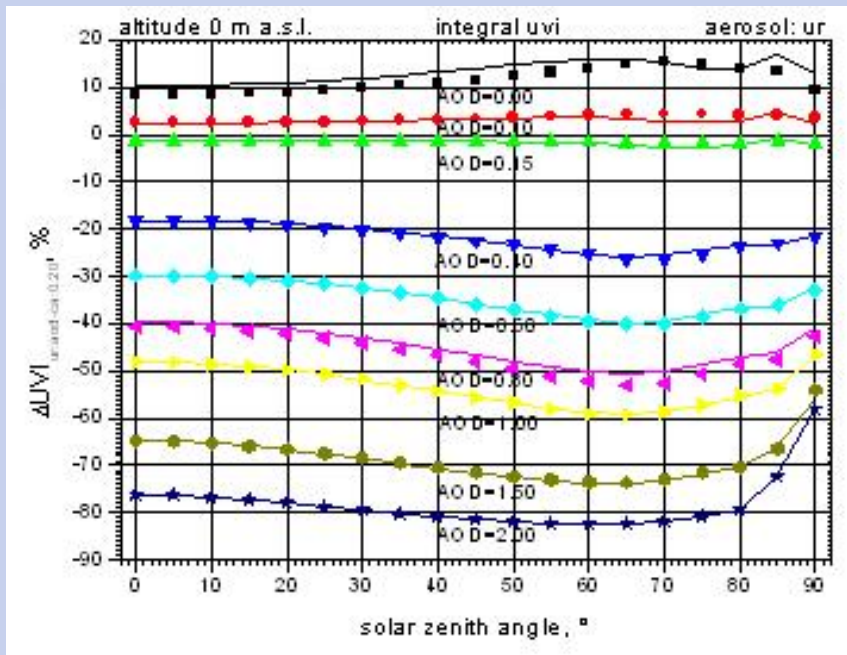
- influence of AOD increases with decreasing SSA
- dependence on SZA is low, strongest at $\sim 60^\circ$

Large-scale UV Index

Adjusted to variable aerosol by a factor dependent on

- AOD, SSA, SZA

Aerosol effects: $f_{\text{aerosol}}(\text{AOD}, \text{SSA}, \text{SZA})$



UVI, maximum abs. error: 0.21

Accounting for aerosol effects:

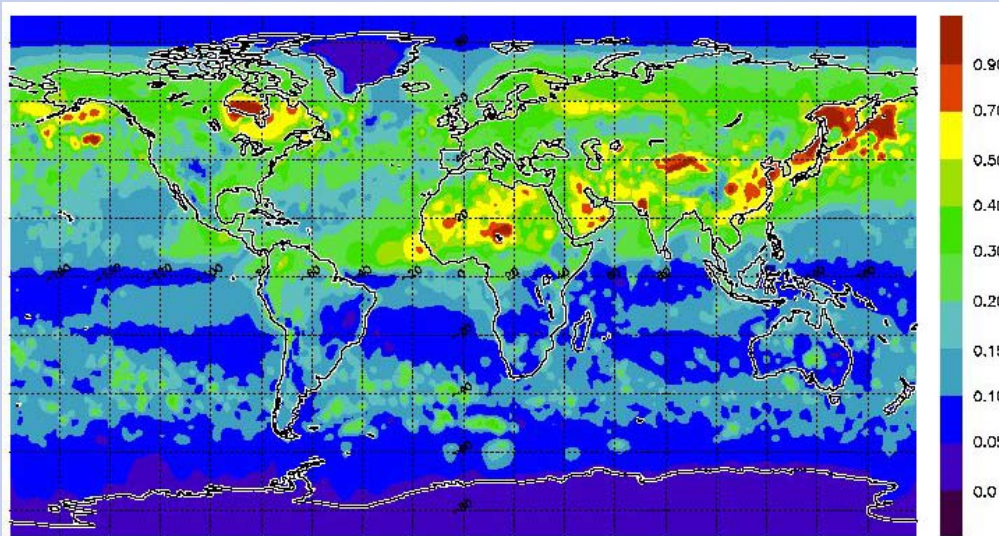
- influence of AOD increases with decreasing SSA
- dependence on SZA is low, strongest at $\sim 60^\circ$

Large-scale UV Index

Adjusted to variable aerosol by a factor dependent on

- AOD, SSA, SZA

Seasonal and Regional Variations of AOD and SSA

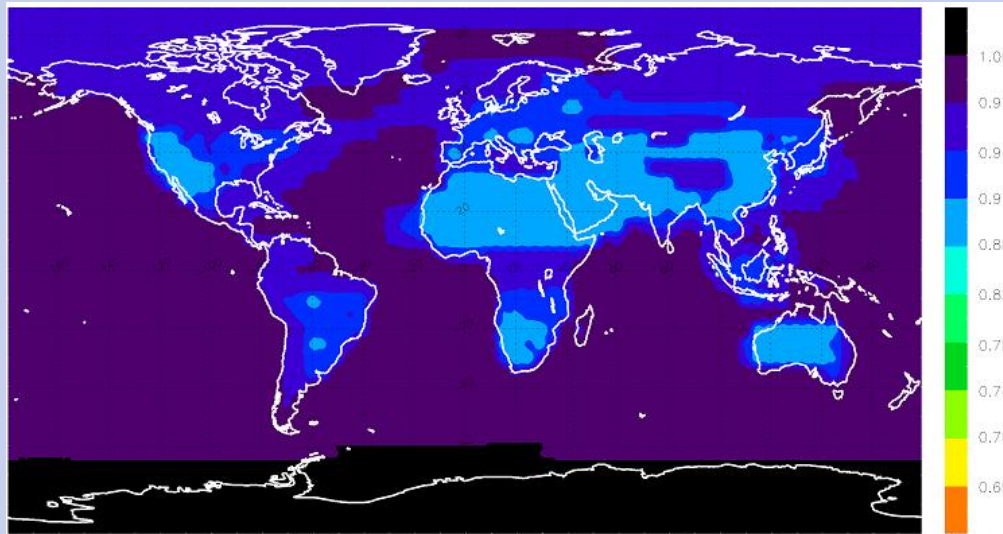


“Climatology”: Average of same months and slight smoothing, April

Accounting for seasonal changes in AOD at 550 nm:

- NASA MOD08_M3, original
- Reduction of pixel mean with increased standard deviation
- Polar regions: set background from GADS and further reading
- Fill data gaps by NASA TOMS AOD, 1979 - 2001
- Average of same months

Seasonal and Regional Variations of AOD and SSA

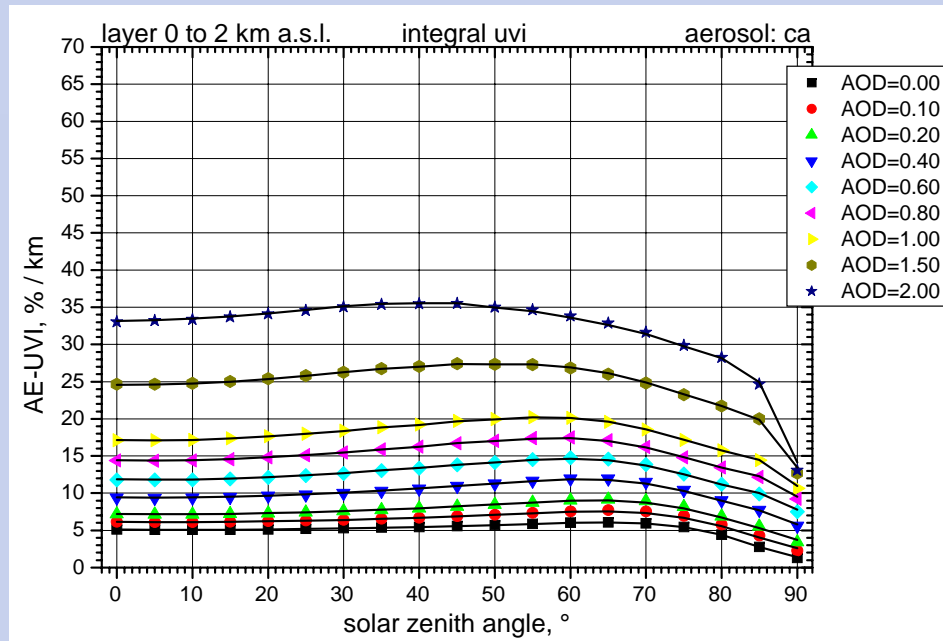


SSA 300 nm (relative Humidity 70%):
GADS limited to SSA > 0.887: Summer

Accounting for semi-annual changes in single scattering albedo (SSA) at 300 nm:

- Global Aerosol Data Set (GADS), University Munich
- spatial resolution: $5^\circ * 5^\circ$

Adjusting for Altitude: $f_{\text{altitude}}(\text{AOD}, \text{SSA}, \text{SZA}, z)$



UVI, maximum abs. error: 0.04

Altitude Effect = relative difference in the irradiance between two altitudes, scaled to 1 km.

Parameterisation formula:

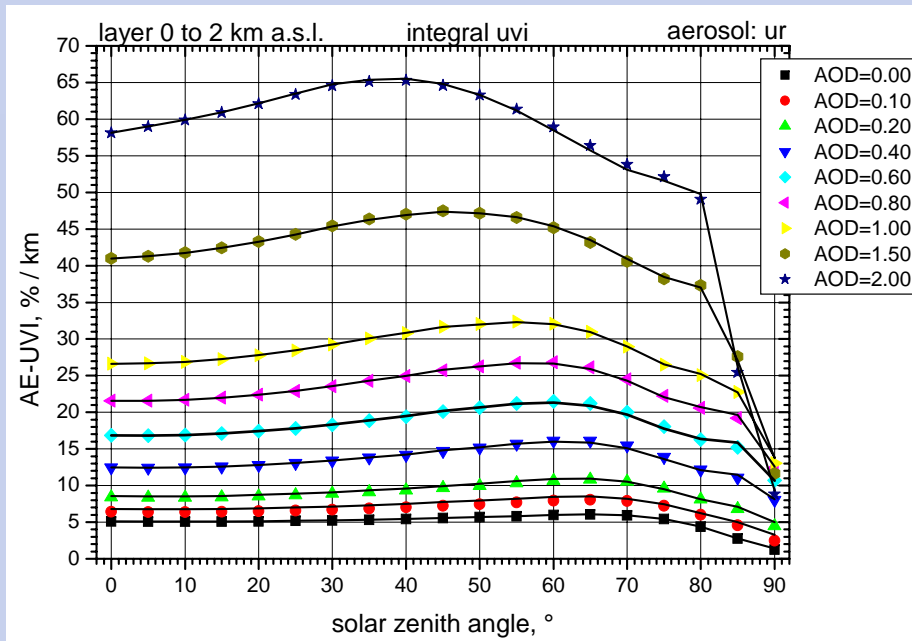
$$\ln(\text{UVI}_z / \text{UVI}_{\text{msl}}) =$$

$$+ \text{Ray}(\text{SZA}, z)$$

$$+ \text{Aer}(\text{SZA}, z, \text{AOD}, \text{SSA})$$

Ray(SZA, z) ~ air pressure
special case for AOD = 0.0

Adjusting for Altitude: $f_{\text{altitude}}(\text{AOD}, \text{SSA}, \text{SZA}, z)$



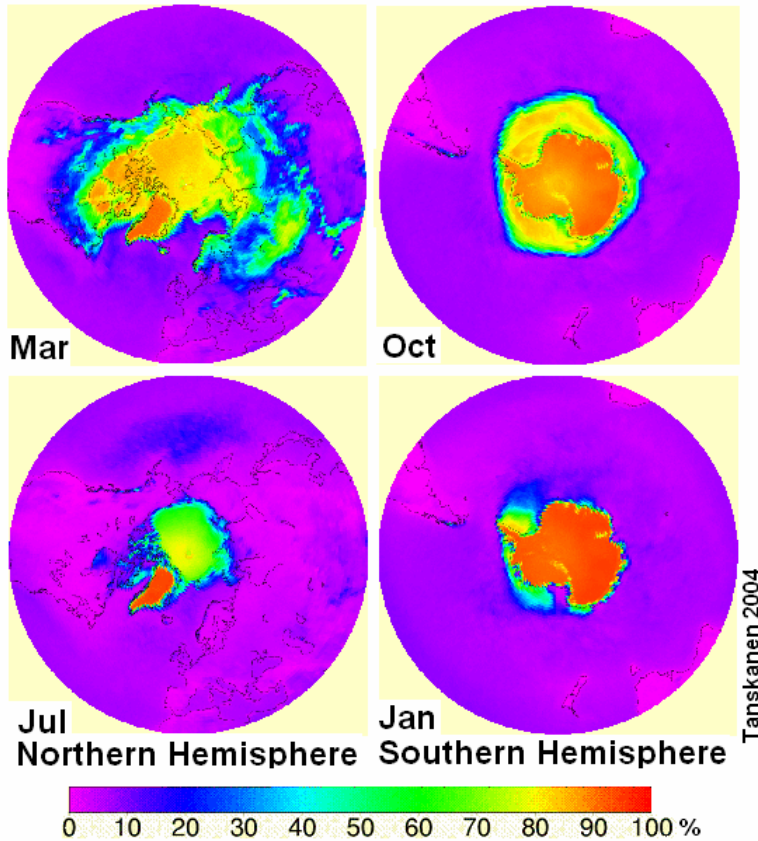
UVI, maximum abs. error: 0.21

Altitude effects of aerosol:

- increasing with increasing AOD
- strongly increasing with decreasing SSA
- slightly dependent on SZA with a maximum at about 50 to 70°

Effect: UVI increase, % per 1 km aerosol “ca”, AOD=0.20: 7 - 8%
aerosol “ca”, AOD=0.40: 10 - 11%
aerosol “ur”, AOD=0.40: 14 - 16%

Lambertian Surface Albedo Climatology
MTW algorithm applied on TOMS at 360 nm



Effects on UVI: 0 - 40 %

Adjusting for Surface Albedo

$$f_{\text{albedo}}(\text{ALB}, \text{SZA}, z)$$

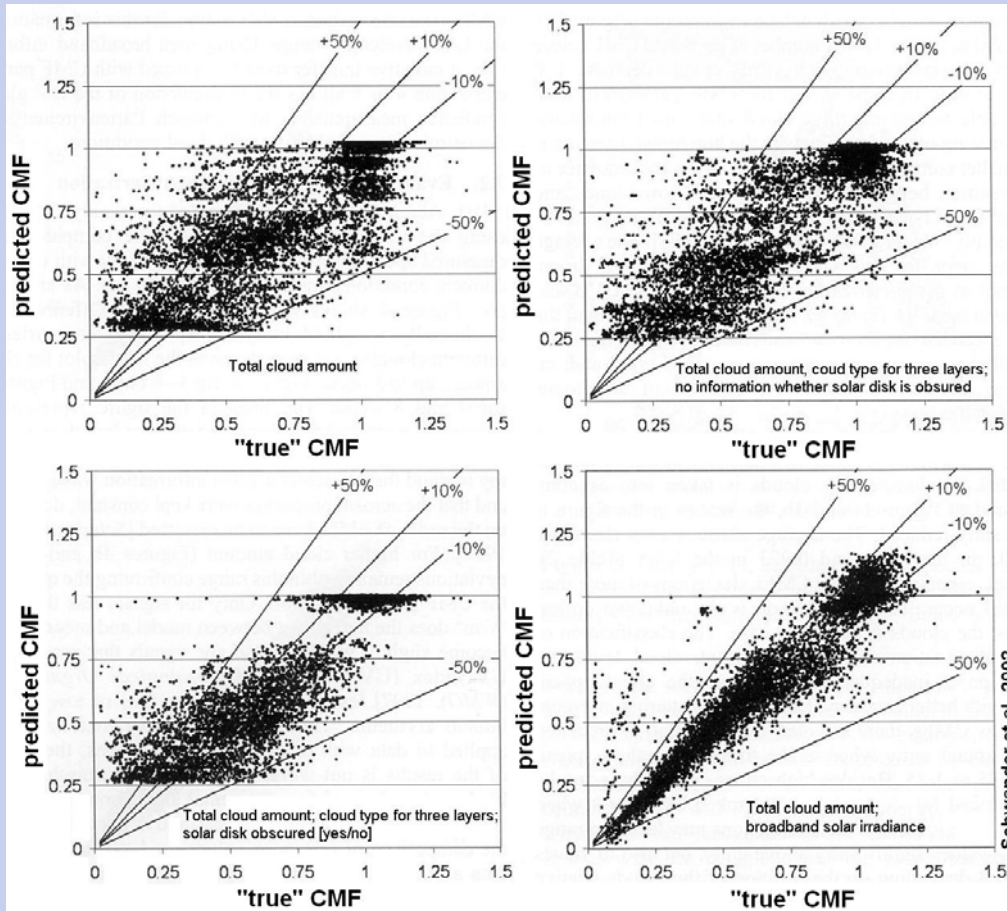
Local surface albedo defined by

$$\text{alb} [\%] = 100. * (\text{irr} \uparrow / \text{irr} \downarrow)$$

UV surface albedo generally is low, exception:

- snow/ice, fresh : ~95 %
- old/wet : ~70 %

The irradiance at a receiver is determined by the albedo of an area of up to 30 km².



Cloud Modification Factor (CMF)

$$CMF = UVI_{cloudy} / UVI_{clear_sky}$$

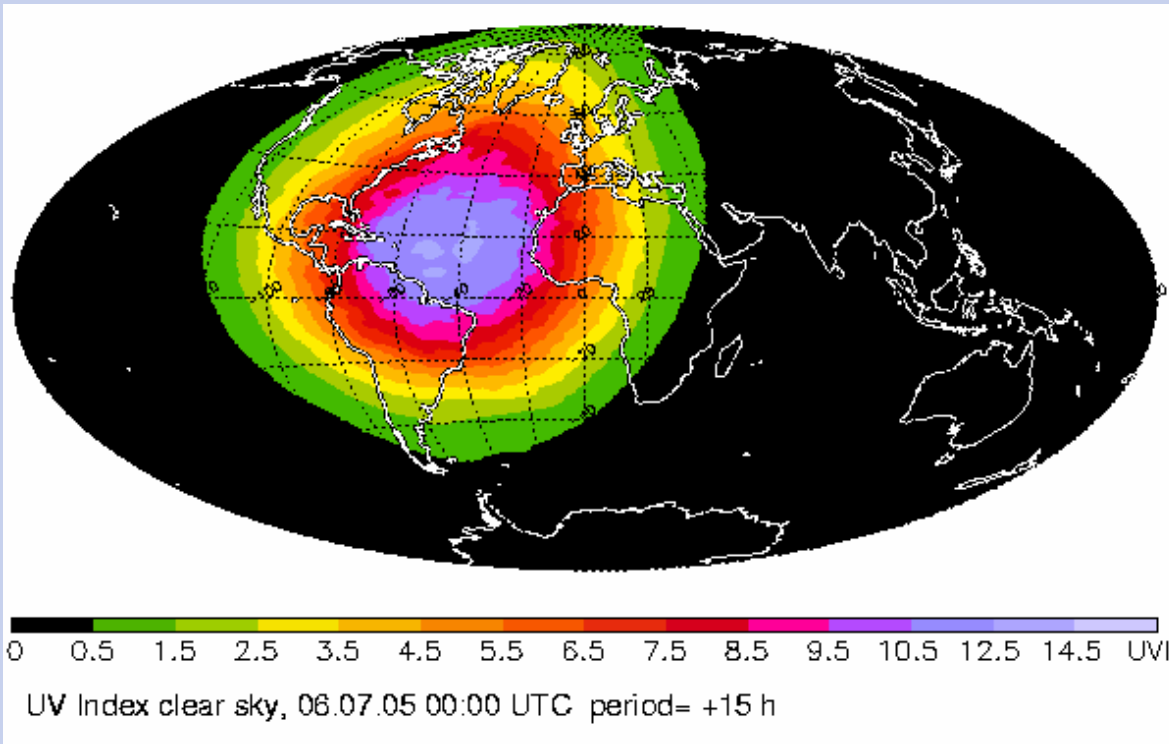
Function of

- cloud cover
- cloud type
- short-wave irradiance

Challenging is the high micro- and macro-physical variability in space and in time

Range of CMF: 1.00 - 0.20

Examples of UV Index Calculation and Derived Products



Radiation transfer:

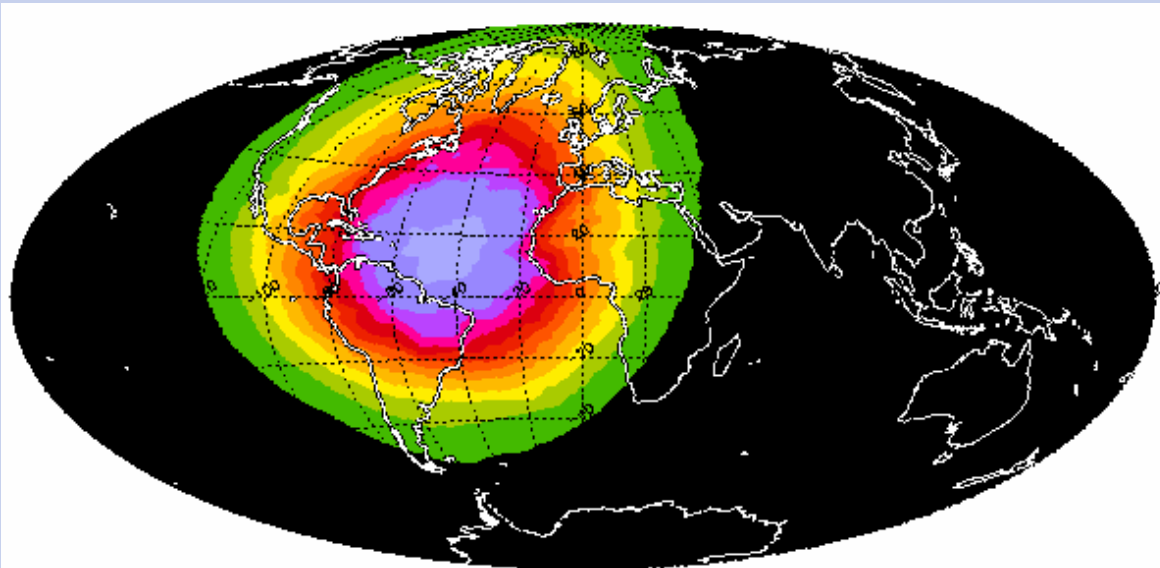
- lookup tables
- depending on
- solar zenith angle
 - total column ozone

Resolution:

- 12 month
- 10 zonal belts

Set : clear sky , albedo, mean sea level , AOD/SSA
Variable : total column ozone, solar zenith angle

Examples of UV Index Calculation and Derived Products



0 0.5 1.5 2.5 3.5 4.5 5.5 6.5 7.5 8.5 9.5 10.5 12.5 14.5 UVI

UV Index clear sky, 06.07.05 00:00 UTC period= +15 h

Set : clear sky, albedo, mean sea level
Variable : AOD/SSA

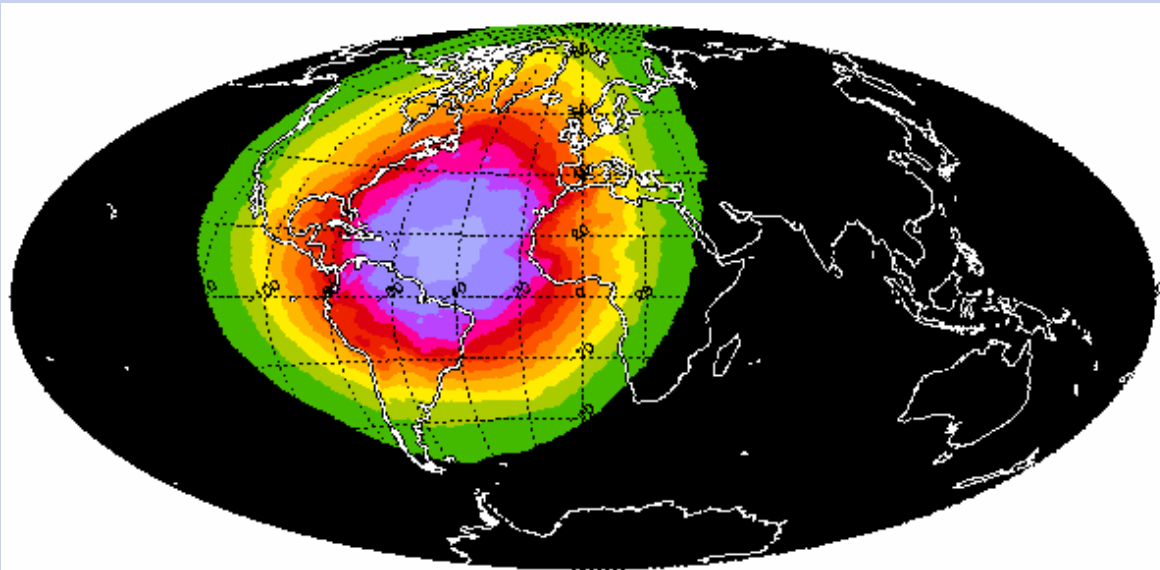
Radiation transfer:

- lookup tables

Adjustment for

- aerosol optical depth (AOD)
- single scattering albedo (SSA)

Examples of UV Index Calculation and Derived Products



0 0.5 1.5 2.5 3.5 4.5 5.5 6.5 7.5 8.5 9.5 10.5 12.5 14.5 UVI

UV Index clear sky, 06.07.05 00:00 UTC period= +15 h

Set : clear sky, albedo
Variable : AOD/SSA, topography

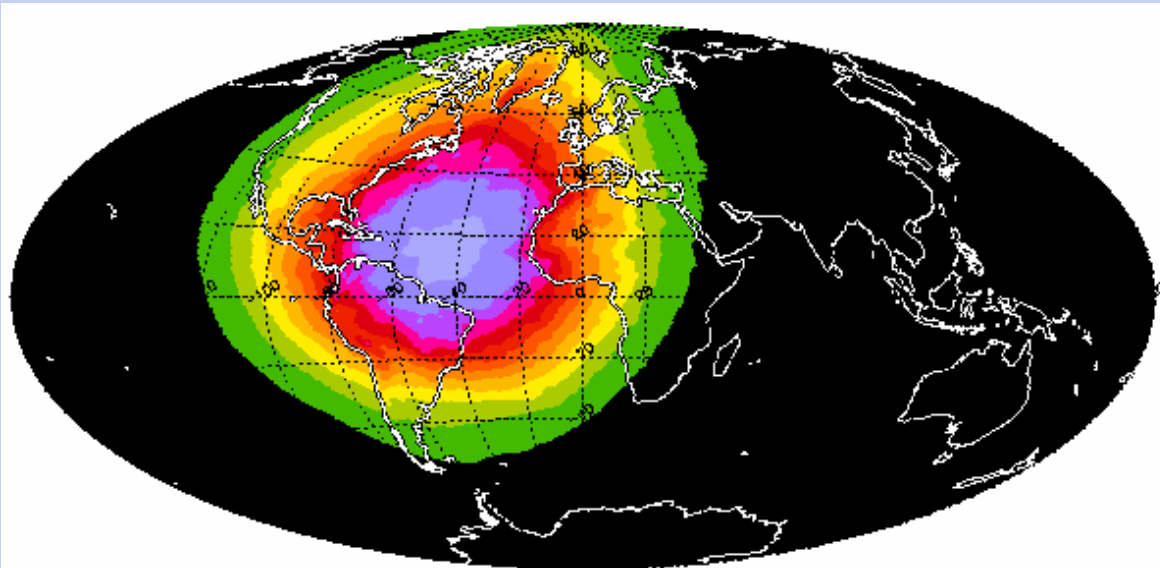
Radiation transfer:

- lookup tables

Adjustment for

- AOD and SSA
- altitude a.m.s.l

Examples of UV Index Calculation and Derived Products



0 0.5 1.5 2.5 3.5 4.5 5.5 6.5 7.5 8.5 9.5 10.5 12.5 14.5 UVI

UV Index clear sky, 06.07.05 00:00 UTC period= +15 h

Set : clear sky
Variable : AOD/SSA, topography, albedo

Radiation transfer:

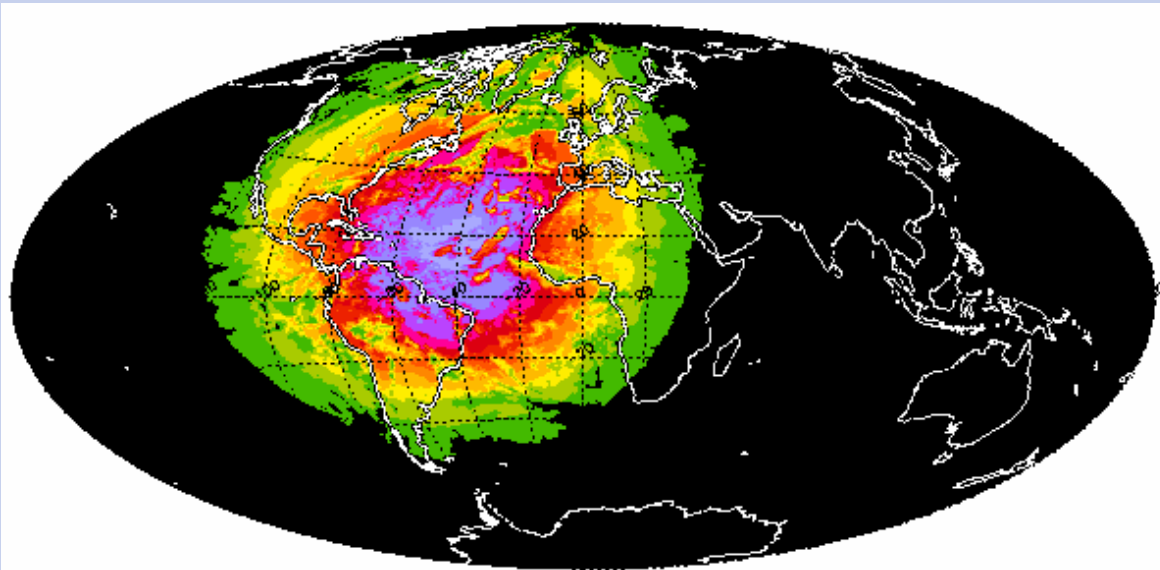
- lookup tables

Adjustment for

- AOD and SSA
- altitude a.m.s.l
- albedo

UV Index clear sky

Examples of UV Index Calculation and Derived Products



0 0.5 1.5 2.5 3.5 4.5 5.5 6.5 7.5 8.5 9.5 10.5 12.5 14.5 UVI

UV Index cloudy, 06.07.05 00:00 UTC period= +15 h

Set : ---

Variable : AOD/SSA, topography, albedo, clouds

Radiation transfer:

- lookup tables

Adjustment for

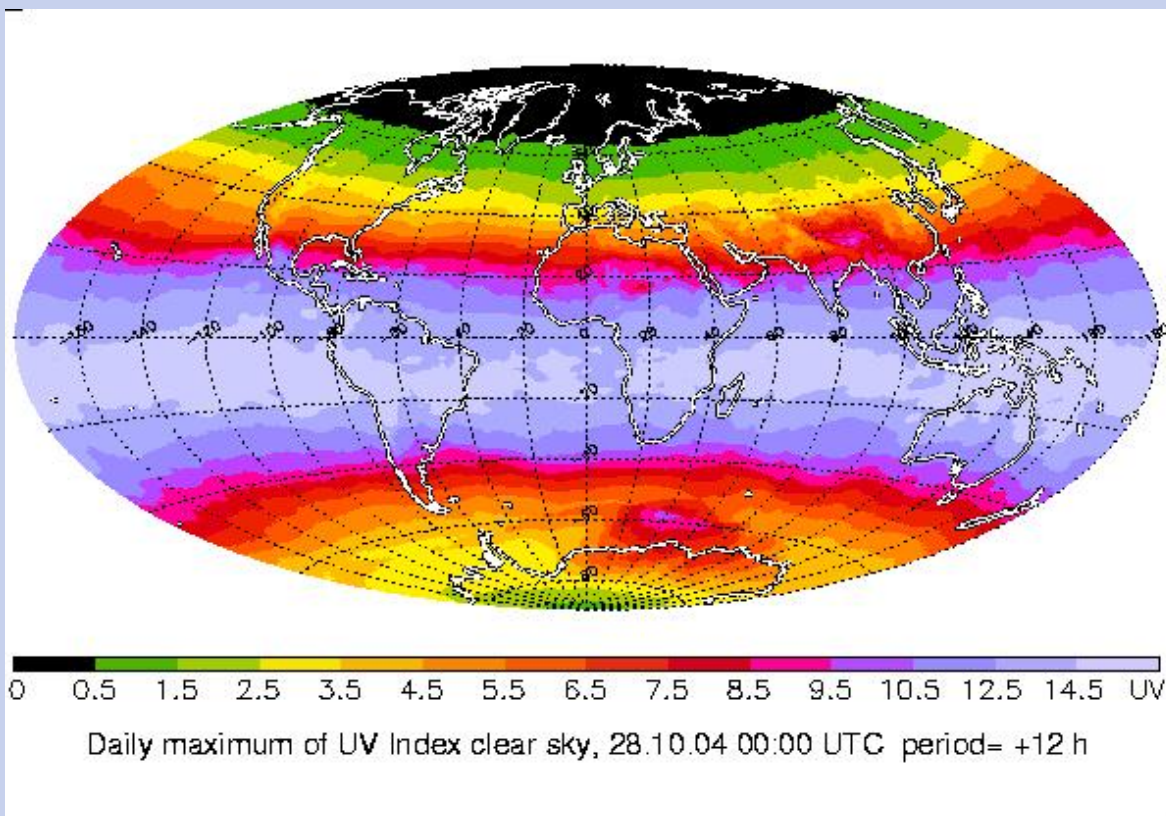
- AOD and SSA
- altitude a.m.s.l
- albedo

UV Index clear sky

- CMF

UV Index cloudy

Examples of UV Index Calculation and Derived Products



Daily Maximum of UV Index clear sky

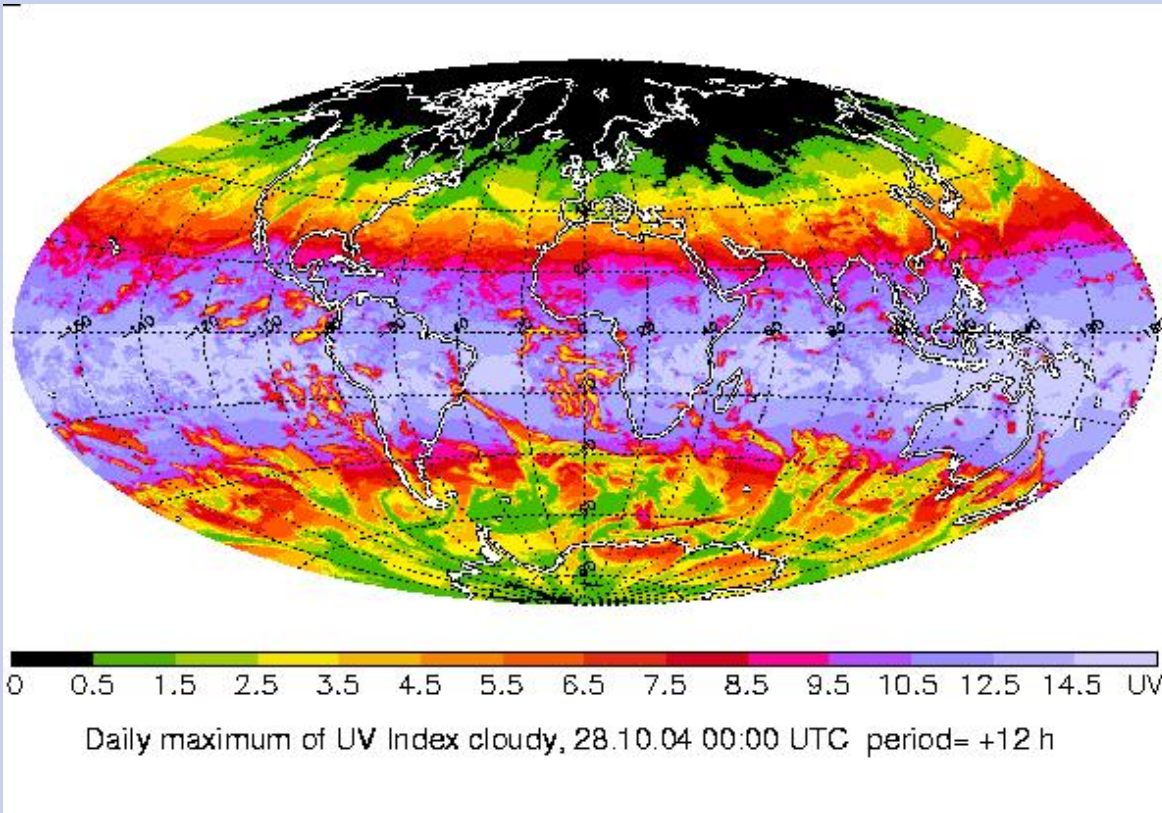
- reassembly of hourly data

Antarctic Ozone Hole:

additional effects:

- sun elevation mid spring
- albedo of snow/ice

Examples of UV Index Calculation and Derived Products



Daily Maximum of UV Index cloudy

- reassembly of hourly data

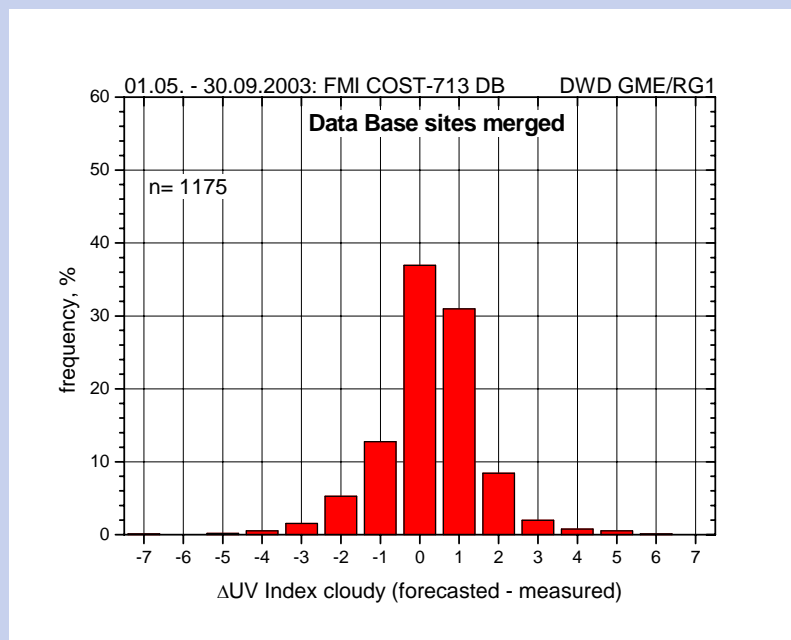
Antarctic Ozone Hole:

additional effects:

- sun elevation mid spring
- albedo of snow/ice
- in essential parts masked by clouds

Verification of forecasted UV Index

Measured UV Index: FMI COST-713 UV Index data base: 21.11.03



sites merged:	11	
co-ordinates:	Europe	
time series:	01.05. - 30.09.2003	
measurements:	1175	
	UVI	ozone, DU
measured mean:	5.23	323.6
forecasted mean:	5.50	331.7
bias:	+0.27	+8.1
rms error:	1.29	18.6
correlation:	0.88	.88

Acknowledgement: Thanks to all contributors and to FMI hosting the data base



Conclusion

- Aerosols can influence the UV Index up to an amount comparable to that of day to day changes in ozone column.
- UV Index forecasting is prepared to account for aerosol effects.
- GEMS could be a source providing operationally:
 - current values of aerosol optical depth and aerosol type as near real time assimilated fields of aerosol properties having a global coverage.
 - future: forecasts.
- The largest uncertainties are due to cloudiness.
 - Improvements for the future will be possible.
 - COST-726 on UV climatology and trends can contribute to more accurate Cloud Modification Factors.

WMO EC-LV (2003) approves a UV Index function for RSMC Offenbach