Representation of model uncertainties for ensemble forecasts

ECMWF Ensemble Prediction Section

Martin Leutbecher, Simon Lang, Sarah-Jane Lock, Pirkka Ollinaho, Antje Weisheimer



Annual Seminar 2015: Model Uncertainty

Slide 1

Representation of model uncertainties

- Why represent model uncertainty in an ensemble forecast?
- What are the sources of model uncertainty?
- How do we represent model uncertainty?
 - 2 stochastic physics schemes in the IFS
- Impact of stochastic physics schemes in the IFS:
 - Medium-range ensemble (ENS)
 - Seasonal forecast (S4)



Forecast uncertainty via ensemble reliability

In a reliable ensemble, ensemble spread is a predictor of ensemble error



- Ensemble member
- Ensemble mean
- Observation

i.e. averaged over many ensemble forecasts,

Different ensemble members arise due to perturbations to initial conditions and to model integrations

 $e(\bar{x}) \approx \sigma(x)$



Annual Seminar 2015: Model Uncertainty

sarah-jane.lock@ecmwf.int

Forecast uncertainty via ensemble reliability

In an under-dispersive ensemble,

 $e(\bar{x}) \gg \sigma(x)$



- Ensemble member
- Ensemble mean
- Observation

and ensemble spread does not provide a good estimate of error.

What happens when the ensemble forecast includes no representation of model uncertainty?



Annual Seminar 2015: Model Uncertainty

sarah-jane.lock@ecmwf.int

What happens with no accounting for model uncertainty?





Annual Seminar 2015: Model Uncertainty

sarah-jane.lock@ecmwf.int

Model uncertainty: where does it come from?

Processes parametrised in the model:





Annual Seminar 2015: Model Uncertainty

sarah-jane.lock@ecmwf.int

Model uncertainty: where does it come from?

• Any other sources: processes missing from the underlying model?

Atmosphere exhibits upscale propagation of kinetic energy (KE)

- Occurs at ALL scales: no concept of "resolved" and "unresolved" scales
- How can the model represent upscale KE transfer from unresolved to resolved scales?



Annual Seminar 2015: Model Uncertainty

Model uncertainty: how to simulate it?

- What do the model errors look like?
- What is the relative size of model error from different sources?
- How can we represent model errors?
- Multi-model ensembles
- Multi-physics ensembles
- Perturbed parameter ensembles
- Stochastic parametrisations"



Stochastic physics schemes in IFS

- IFS ensemble forecasts (ENS and S4) include 2 model uncertainty schemes:
 - Stochastically perturbed physics tendencies (SPPT) scheme
 - Stochastic kinetic energy backscatter (SKEB) scheme
- SPPT scheme: simulates uncertainty due to sub-grid parametrisations
- SKEB scheme: parametrises a missing and uncertain process



SPPT scheme

- Initially implemented in IFS, 1998 (Buizza et al., 1999); revised in 2009:
- Simulates model uncertainty due to physical parameterisations by
 - taking the net parameterized physics tendency:

$$\boldsymbol{X} = \begin{bmatrix} X_U, X_V, X_T, X_Q \end{bmatrix}$$

coming from
radiation

gravity wave drag

vertical mixing

convection

cloud physics

• and perturbing with multiplicative noise $r \in [-1, +1]$ as:

$$X' = (1 + \mu r)X$$

where $\mu \in [0,1]$ tapers the perturbations to zero near the surface & in the stratosphere.

Shutts et al. (2011, ECMWF Newsletter); Palmer et al., (2009, ECMWF Tech. Memo.)

Slide 10

Annual Seminar 2015: Model Uncertainty

• 2D random pattern in spectral space:

First-order auto-regressive [AR(1)] process for evolving spectral coefficients \hat{r}

$$\hat{r}(t + \Delta t) = \phi \hat{r}(t) + \rho \eta(t)$$

where $\phi = \exp(-\Delta t/\tau)$ controls the correlation over timestep Δt ;

and spatial correlations (Gaussian) for each wavenumber define ρ for random numbers, η

- Resulting pattern in grid-point space r:
- clipped such that $r \in [-1, +1]$
- applied at all model levels to preserve vertical structures**
- **Except: tapered to zero at model top/bottom, avoiding:
 - instabilities due to perturbations in the boundary layer;
 - perturbations in the stratosphere due to well-constrained clear-skies radiation



Slide 11

- 2D random pattern of spectral coefficients, *r*:
- Time-correlations: AR(1)
- Space-correlations: Gaussian
- Clipped such that $r \in [-1, +1]$
- Applied at all model levels to preserve vertical structures**
- **Except: tapered to zero at model top/bottom



3 correlation scales:

- i) 6 hours, 500 km, $\sigma = 0.52$
- ii) 3 days, 1 000 km, $\sigma = 0.18$
- iii) 30 days, 2 000 km, $\sigma = 0.06$









Annual Seminar 2015: Model Uncertainty

sarah-jane.lock@ecmwf.int

SKEB scheme

Introduced into IFS, 2010:

Attempting to simulate a process otherwise absent from the model –

upscale transfer of energy from sub-grid scales

• Represents backscatter of Kinetic Energy (KE) by adding perturbations to U and V via a forcing term to the streamfunction:

$$F_{\varphi} = \left(\frac{b_R D_{\text{tot}}}{B_{\text{tot}}}\right)^{1/2} F^*$$

where F^* is a 3D random pattern field, B_{tot} is the mean KE input by F^* alone, D_{tot} is an estimate of the total dissipation rate due to the model, b_R is the backscatter ratio – a scaling factor.

> Shutts et al. (2011, ECMWF Newsletter); Palmer et al., (2009, ECMWF Tech. Memo.); Shutts (2005, QJRMS); Berner et al. (2009, JAS)

SKEB pattern

$$F_{\varphi} = \left(\frac{b_R D_{\text{tot}}}{B_{\text{tot}}}\right)^{1/2} F^*$$

• 3D random pattern field F^* :

- First-order auto-regressive [AR(1)] process for evolving F^*

$$F^*(t + \Delta t) = \phi F^*(t) + \rho \eta(t)$$

where $\phi = \exp(-\Delta t/\tau)$ controls the correlation over timestep Δt ;

and spatial correlations (power law) for wavenumbers define ρ for random numbers, η

- vertical space-(de)correlations: random phase shift of η between levels



SKEB perturbations

$$F_{\varphi} = \left(\frac{b_R D_{\text{tot}}}{B_{\text{tot}}}\right)^{1/2} F^*$$

• *D*_{tot} is an estimate of sub-grid scale production of KE, and includes:

- D_{num} = numerical dissipation from
 - explicit horizontal diffusion (bi-harmonic, ∇^2); and
 - estimate due to semi-Lagrangian interpolation error
- *D*_{con} = estimated KE generated by updraughts and detrainment within sub-grid deep convection



How are the perturbation patterns determined?

- Characteristics of model errors cannot be determined from observations:
 - uncertain processes are small-scale (space and time)
 - lack of observational coverage

Can attempt to use models: coarse-graining studies (e.g. Shutts and Palmer, 2007)

- take high-resolution model simulations as "truth"
- coarse-grain: average high-res model fields and tendencies (or streamfunction) to a grid-resolution typical of the forecast model
- compare the contribution of "sub-grid" scales in the coarse-grained simulation with parametrisations in the forecast model
- coarse-graining studies have been used to justify and inform scales in SPPT and SKEB perturbation patterns



IFS ensembles: ENS and System 4 (S4)

- **ENS** = ensemble prediction system for
 - medium-range forecasts (up to 15 days) and
 - monthly forecasts (up to 46 days)
- S4 = seasonal forecasting system
 - up to 7 months
- Both systems represent model uncertainty with SPPT and SKEB

• ENS:

- 1 control forecast + 50 perturbed members
- T639 (~32 km) resolution to day 10; T319 (~65 km) days 10-15
- 91 vertical levels, up to 0.01hPa



Impact of SPPT and SKEB in ENS





Annual Seminar 2015: Model Uncertainty

sarah-jane.lock@ecmwf.int

Impact of SPPT and SKEB in ENS





Annual Seminar 2015: Model Uncertainty

sarah-jane.lock@ecmwf.int

- Adding SPPT + SKEB perturbations:
 - increases ensemble "spread" (= ensemble standard deviation), i.e. ensemble members describe greater region of the parameter space
 - some reduced ensemble mean errors
- In the extra-tropics:
 - SPPT and SKEB each have a similar impact, i.e. perturbations are successfully adopted and evolved by the model
 - Experiments: perturbations in days 0-5 contribute most effect
- In the tropics:
 - SPPT has a much greater impact (in terms of both spread and error) than SKEB, i.e. SPPT perturbations more able to excite modes that the model can evolve
 - Experiments: effect of perturbations rapidly lost at all times



Impact of SPPT and SKEB in S4

- System 4 (S4), November 2011: introduction of revised SPPT and SKEB
- Operational configuration:
 - T255 (~80 km), 91 vertical levels (up to 0.01 hPa)
 - Coupled ocean model: NEMOv3.0, 1 degree (~110 km), 42 vertical levels
 - 51 members
 - Initialised on 1st of each month
 - Forecast lead times: to 7 months
- Recent work with S4 to assess impact of stochastic schemes
- For longer time-scales, consider impact in terms of:
 - Noise-induced drift, i.e. change in model mean
 - Noise-activated regime transition, e.g. Pacific-N. American region regimes



Impact of SPPT and SKEB in S4

- Recent work with S4 to assess impact of stochastic schemes:
 - Hindcast period: 1981-2010
 - Start dates: May, Aug & Nov
 - Ensemble size: 51
 - Forecasts to lead times: 4-7 months
- Considers impact of SPPT + SKEB on:
 - Systematic errors
 - Madden-Julian Oscillation (MJO) statistics
 - Circulation regimes over the Pacific-North American region



Impact of SPPT and SKEB in S4: systematic errors

Outgoing Longwave Radiation (DJF 1981-2010) stochphysOFF – ERA-I





- SPPT+SKEB: reduction of overly active tropical convection
- Similar reductions in excessive:
 - Total cloud cover
 - Total precip
 - Zonal winds (850 hPa)
- SPPT is responsible for most of the difference; SKEB has little impact





Impact of SPPT and SKEB in S4: MJO



Wheeler and Hendon Index: projection of daily data on 2 dominant combined EOFs of OLR, u200 and u850 over 15° N-15° S

SPPT+SKEB:

- **Increased frequency** of events
- Improved amplitude distribution

S4 – ERA-I

3

4

Impact of SPPT & SKEB in S4: Pacific North America (PNA) circulation regimes



sarah-jane.lock@ecmwf.int

-250-200-150-90 -80 -30 30 60 90 150 200 250 m

Model uncertainty representation: brief outlook for IFS

Exploring alternative stochastic perturbations:

Currently, in SPPT, we perturb:

```
X = X_{RAD} + X_{GWD} + X_{MIX} + X_{CON} + X_{CLD}
```

with zero perturbations near the surface and in the stratosphere.

Instead, identify and perturb individual uncertain parameters.

e.g. Boundary layer:



Total (physics) temperature tendency (K)

Model uncertainty representation: brief outlook for IFS

Exploring alternative stochastic perturbations:

Currently, in SPPT, we perturb:

 $\boldsymbol{X} = \boldsymbol{X}_{RAD} + \boldsymbol{X}_{GWD} + \boldsymbol{X}_{MIX} + \boldsymbol{X}_{CON} + \boldsymbol{X}_{CLD}$

with zero perturbations near the surface and in the stratosphere.

Instead, identify and perturb individual uncertain parameters.

e.g. Boundary layer:



Total (physics) temperature tendency (K)

Representing model uncertainty: summary

- Model uncertainty arises due to unresolved and misrepresented processes
 - finite-resolution of a discrete numerical model
 - parametrisations must describe multi-scale sub-grid processes in bulk
- Difficult to characterise sources of model errors due to lack of observations
- Without representing model uncertainty, ensembles are under-dispersive
- ECMWF ensembles include 2 stochastic model uncertainty schemes:
 - SPPT: representing uncertainty due to sub-grid physics parameterisations
 - SKEB: simulating upscale transfer of kinetic energy from unresolved scales
- Medium-range: increased ensemble spread, greater probabilistic skill
- Seasonal: reduction in biases; better representation of MJO, PNA regimes
- Outlook: Seeking to focus perturbations on individual uncertain parameters

