

# From small-scale turbulence to large-scale convection: a unified scale-adaptive EDMF parameterization

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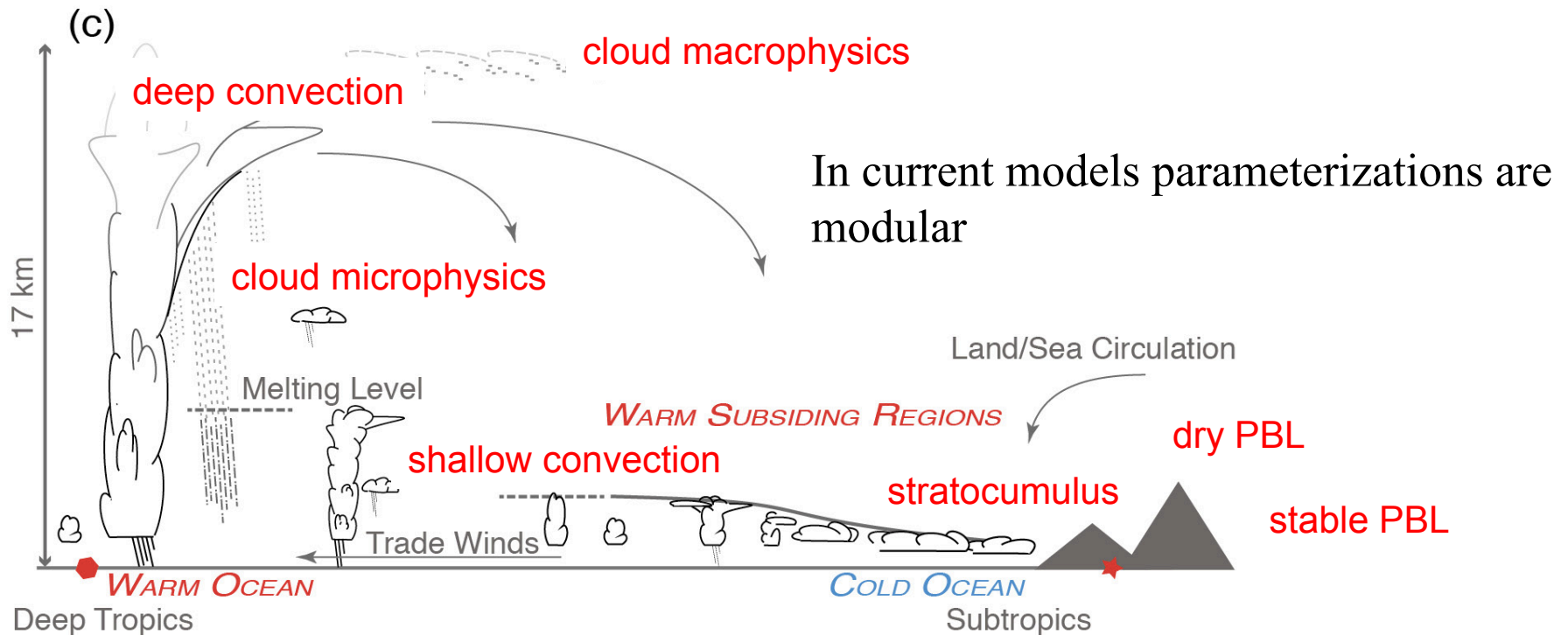
*Shedding light on the greyzone*

*ECMWF, Nov 14, 2017*

# From small-scale turbulence to large-scale convection: a unified scale-adaptive EDMF parameterization

1. **Unified approach to parameterization**
  - Eddy-Diffusivity/Mass-Flux model
2. **Scale adaptive eddy-diffusivity approach**

# Turbulence, convection and cloud parameterizations



Artificial modularity leads to many problems: interfaces, transition

**Goal: unified parameterization for boundary layer, convection and macro/micro-physics**

# Turbulence and convective parameterizations

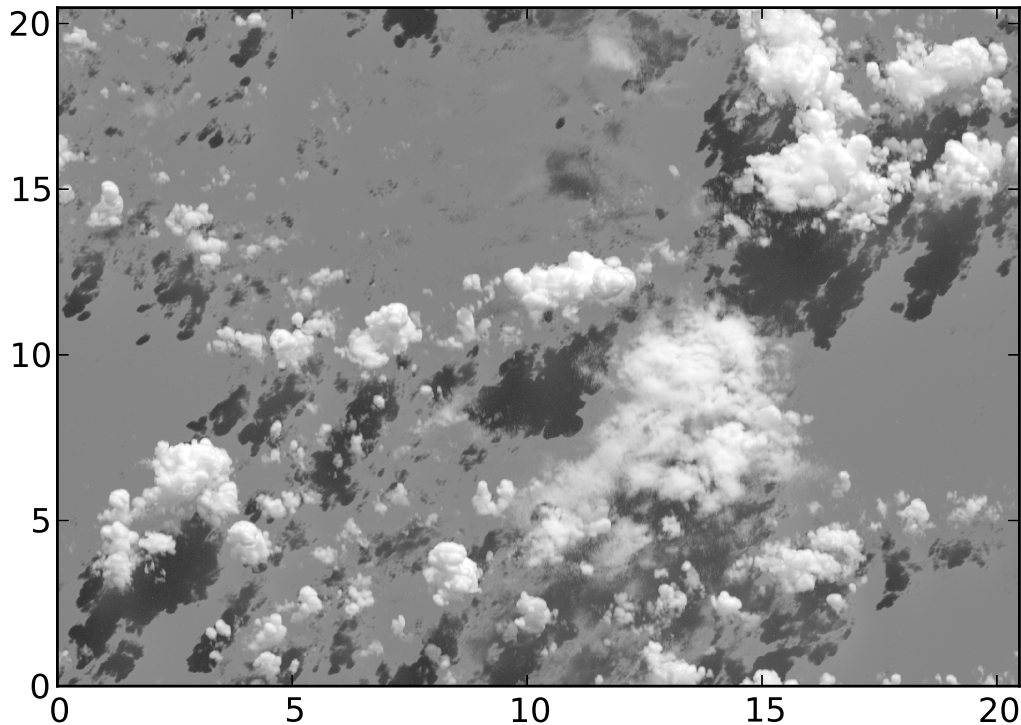
Reynolds-averaged conservation equations:

$$\frac{\partial \bar{\phi}}{\partial t} + \frac{\partial}{\partial x}(\overline{u\phi}) + \frac{\partial}{\partial y}(\overline{v\phi}) + \frac{\partial}{\partial z}(\overline{w\phi}) = -\frac{\partial}{\partial z}(\overline{w'\phi'}) + \bar{S},$$

Boundary Layer	Shallow Convection	Deep Convection
$\overline{w'\phi'} \cong -K \frac{\partial \bar{\phi}}{\partial z}$ <p>Eddy-Diffusivity (ED)</p>	$\overline{w'\phi'} \cong M(\phi_u - \bar{\phi})$ <p>Mass-Flux (MF)</p>	
$\frac{\partial}{\partial t}(\overline{w'\phi'}) = -\frac{\partial}{\partial z}(\overline{w'w'\phi'}) + \dots$ <p>Higher-Order Closure, e.g. CLUBB</p>		
$\overline{w'\phi'} = -k \frac{\partial \bar{\phi}}{\partial z} + M(\phi_u - \bar{\phi})$ <p>Eddy-Diffusivity/Mass-Flux (EDMF)</p>		

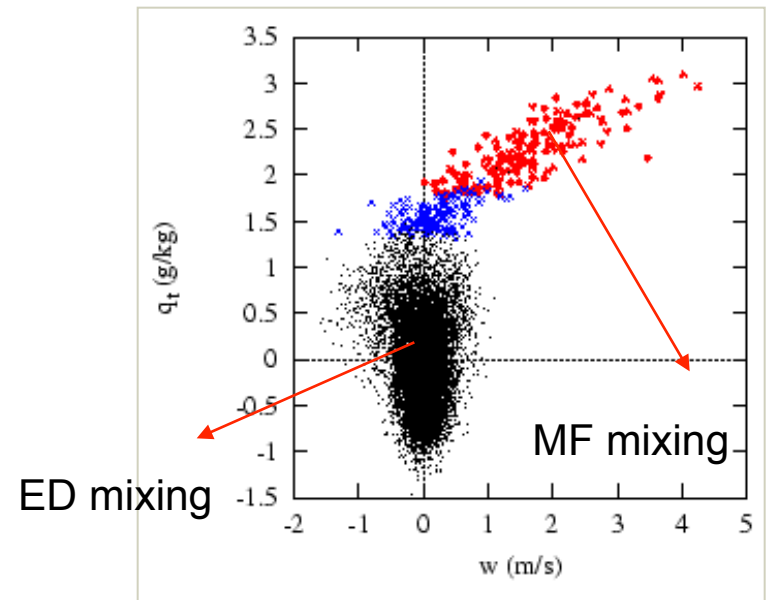
# LES model-informed parameterization development

- LES models solve filtered version of conservation (e.g. Navier-Stokes) equations
- High-resolutions ( $\sim 1 - 100\text{m}$ ) in all 3 dimensions
- LES models resolve most of the essential turbulence/convection
- Closures still needed for scales  $< 10\text{m}$  (but simpler than GCMs)



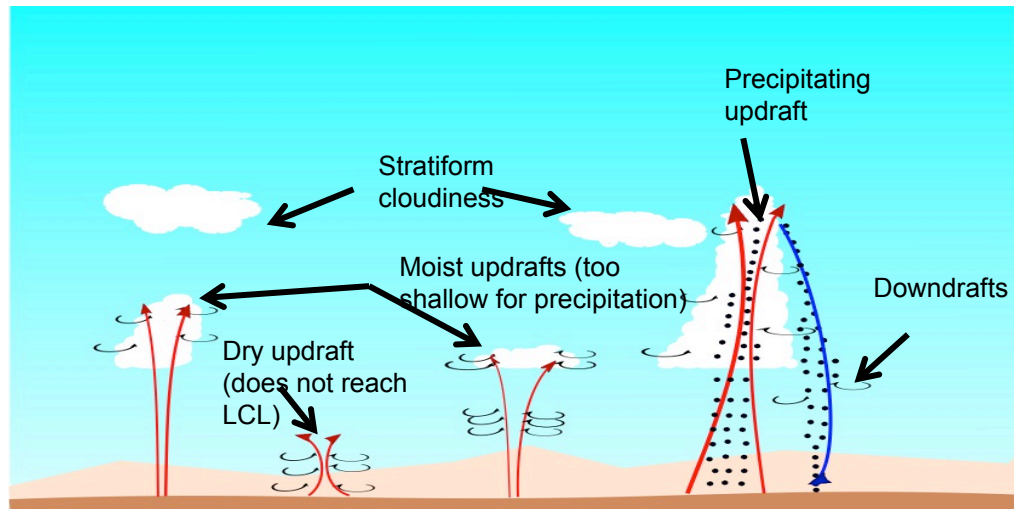
Courtesy of G. Matheou

Bimodal joint pdf( $w, q_t$ ) for convective case



# Our EDMF approach

- Multiple convective plumes – Mass-Flux (MF) model, sum of uniform PDFs
  - Surface driven updrafts
  - Precipitation driven downdrafts
- Non convective environment - Eddy-Diffusivity (ED) model, joint normal PDFs



$$\bar{\varphi} = a_e \varphi_e + \sum_i a_i \varphi_i$$

$i$	Convective elements
$e$	Environment
$a_e, a_i$	Fractional areas

$$\overline{\varphi' \psi'} = \underbrace{a_e \overline{\varphi' \psi'}|_e}_{\text{ED}} + \underbrace{a_e (\varphi_e - \bar{\varphi})(\psi_e - \bar{\psi})}_{\text{Compensating sub.}} + \cancel{\sum_i a_i \overline{\varphi' \psi'}_i} + \underbrace{\sum_i a_i (\varphi_i - \bar{\varphi})(\psi_i - \bar{\psi})}_{\text{Multiple MF}}$$

# Shallow and deep version of EDMF

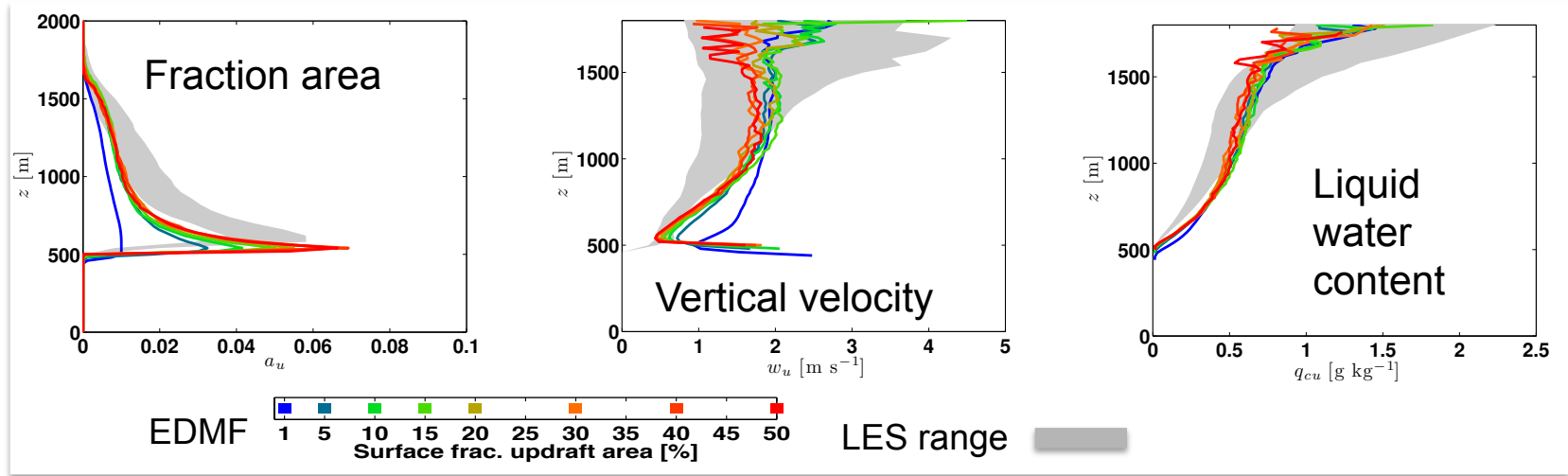
- Non-convective environment:
  - TKE-based eddy-diffusivity approach
- Mass-flux
  - Multiple surface-forced plumes, starting from surface PDF
  - Stochastic entrainment rate
  - Simple Kessler-type microphysics coupled to updraft dynamics
  - Downdrafts driven by evaporation of rain
  - Precipitation-driven cold pools
  - Cold pools impact on updraft entrainment rates and surface PDF

## **Main advantages:**

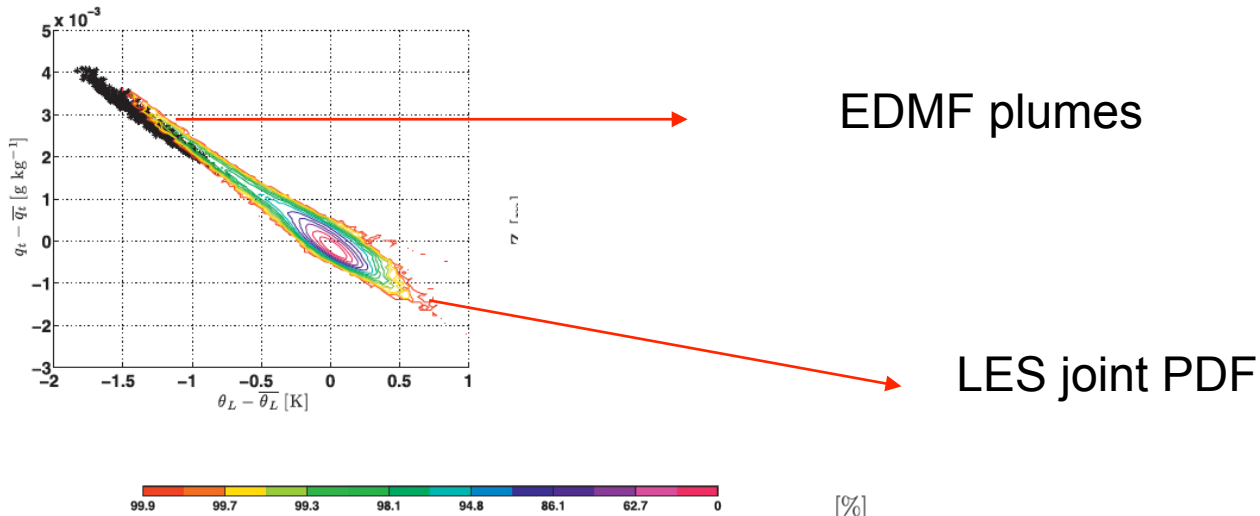
- **Different types of convection within one grid-box**
- **No need for trigger functions and explicit convective closures**
- **Smooth transition between convective regimes (dry, shallow, deep)**

# Shallow convection case - BOMEX

BOMEX: Comparison of EDMF moist updraft properties against LES results

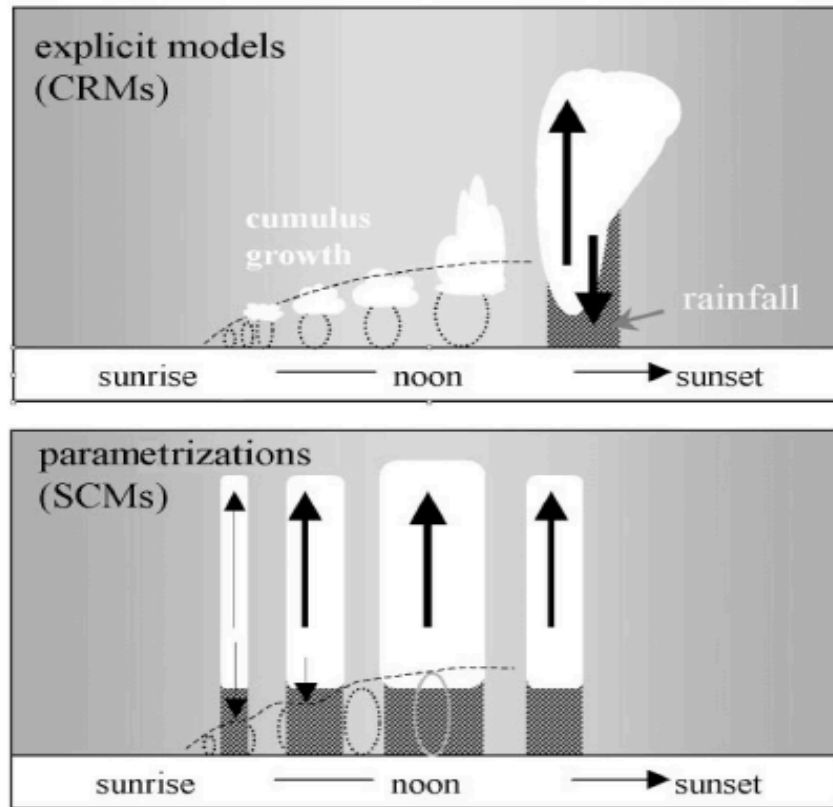


Low sensitivity of multiple-plume EDMF to surface updraft area





# Diurnal cycle of continental convection



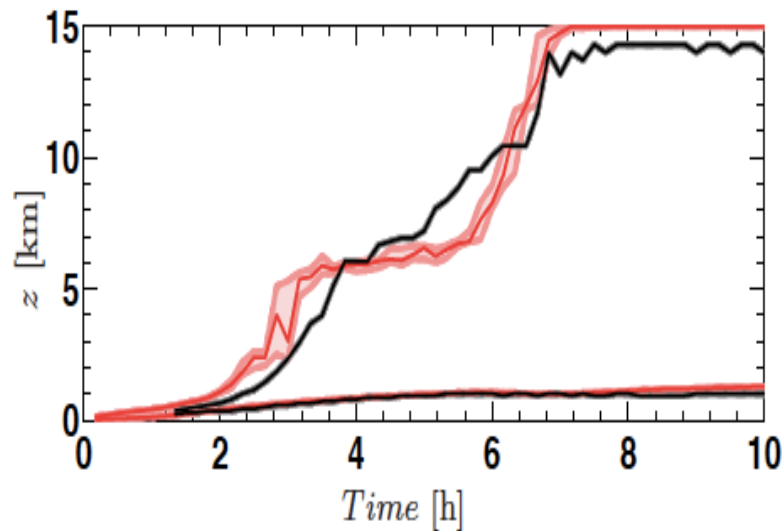
Guichard et al,  
QJRMS, 2004

Climate and weather models often struggle to represent transition between convective regimes

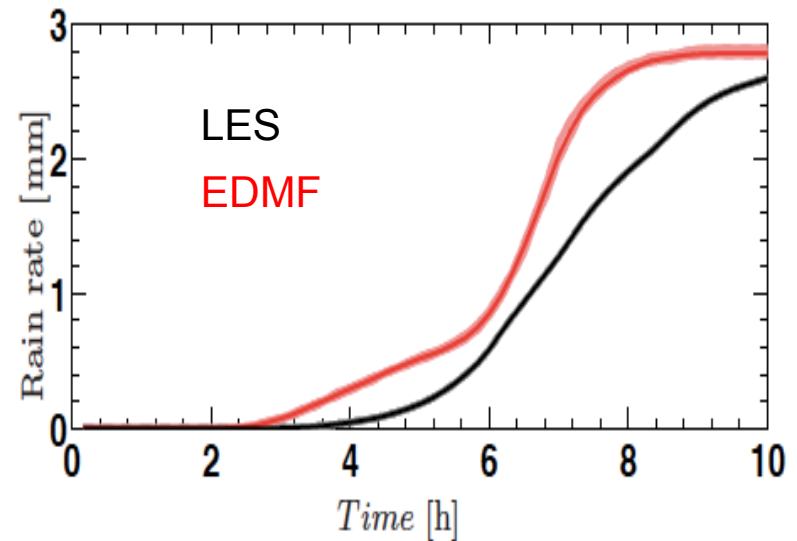
# Unified EDMF, diurnal cycle of convection over land

- New fully unified (PBL + shallow + deep convection) EDMF
- EDMF with cloud microphysics
- LBA diurnal cycle of precipitating convection

Cloud base and



Cumulative surface



Realistic transition with EDMF from shallow to deep convection

# Scale adaptive ED closure for the dry convective boundary layer: from LES to climate scales

$$\overline{\varphi' u'_i} = -K \frac{\partial \bar{\varphi}}{\partial x_i} \quad \text{where} \quad K = l \sqrt{tke}$$

$$l^{-2} = l_{3d}^{-2} + l_{1d}^{-2}$$

Merging the 3D (LES-scale) and 1D (GCM-scale) limits

$$l_{3d} = (\Delta x \Delta y \Delta z)^{1/3}$$

LES scale length ( $\Delta x \sim 10$  m)

$$l_{1d} = f(kz, \tau \sqrt{tke})$$

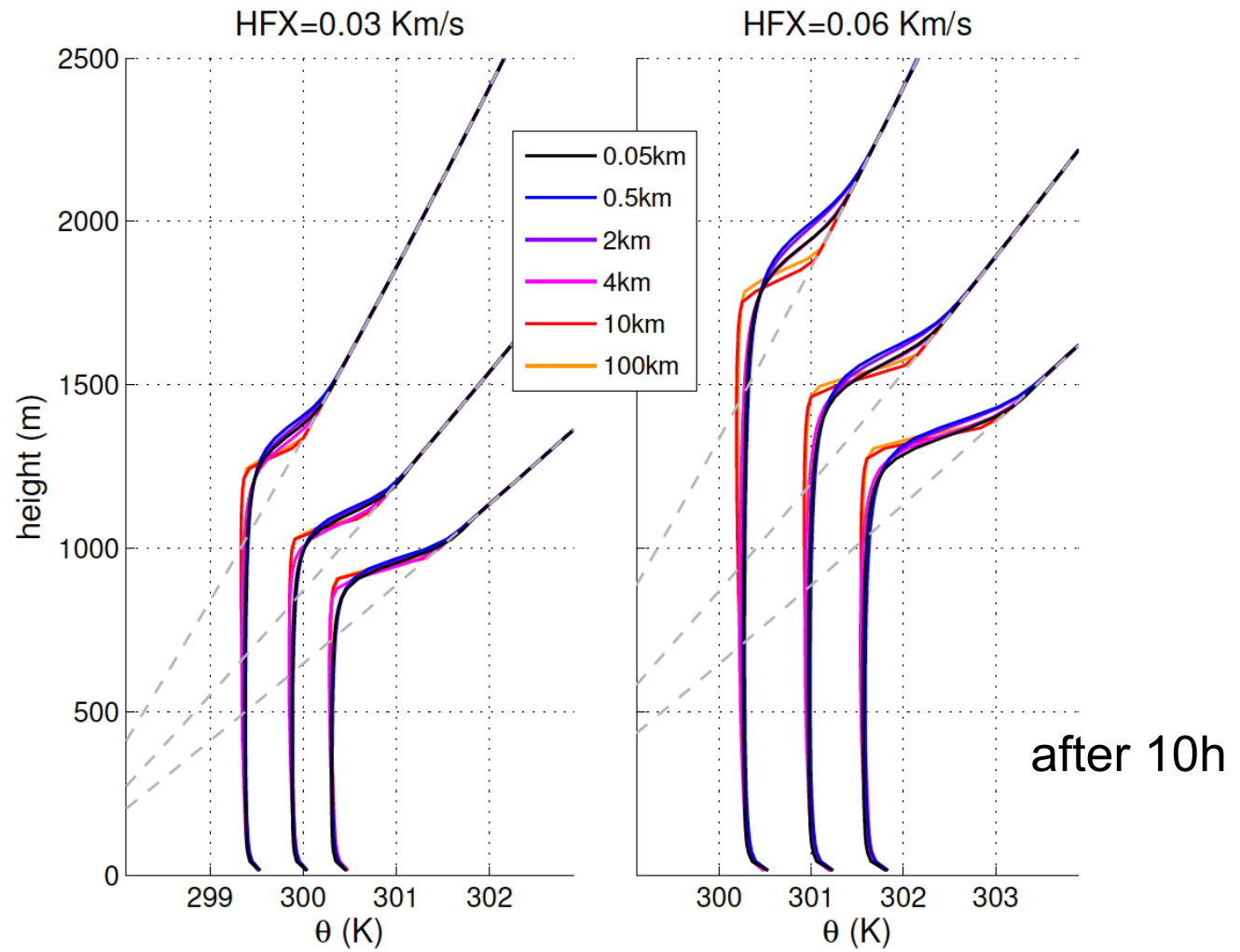
GCM scale ( $\Delta x \sim 100$  km):

# Dry convective boundary layer

WRF model from LES ( $\Delta x=50$  m) to NWP/Climate ( $\Delta x=100$  km)

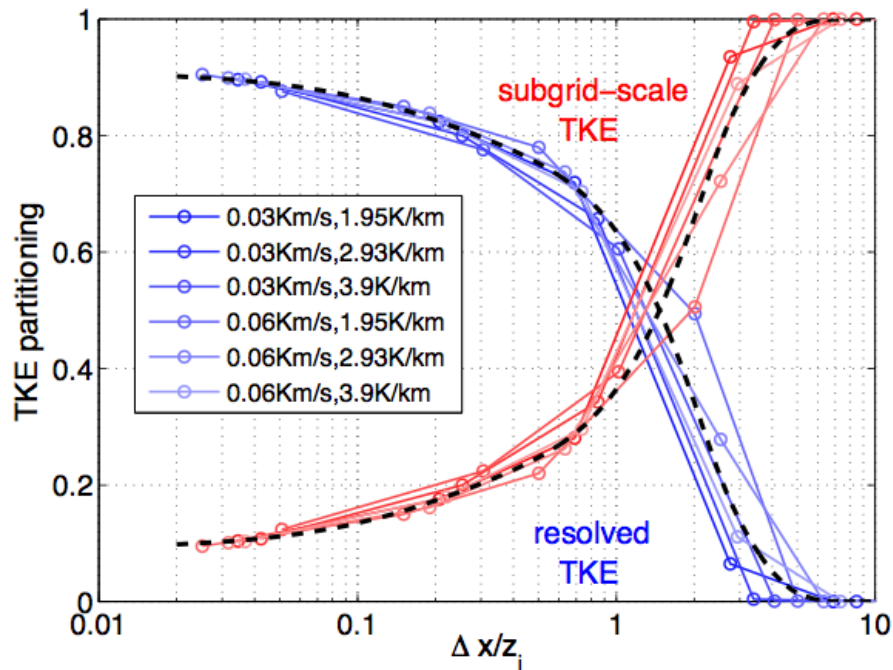
6 cases:

- 3 different stratifications
- 2 different surface heat flux values



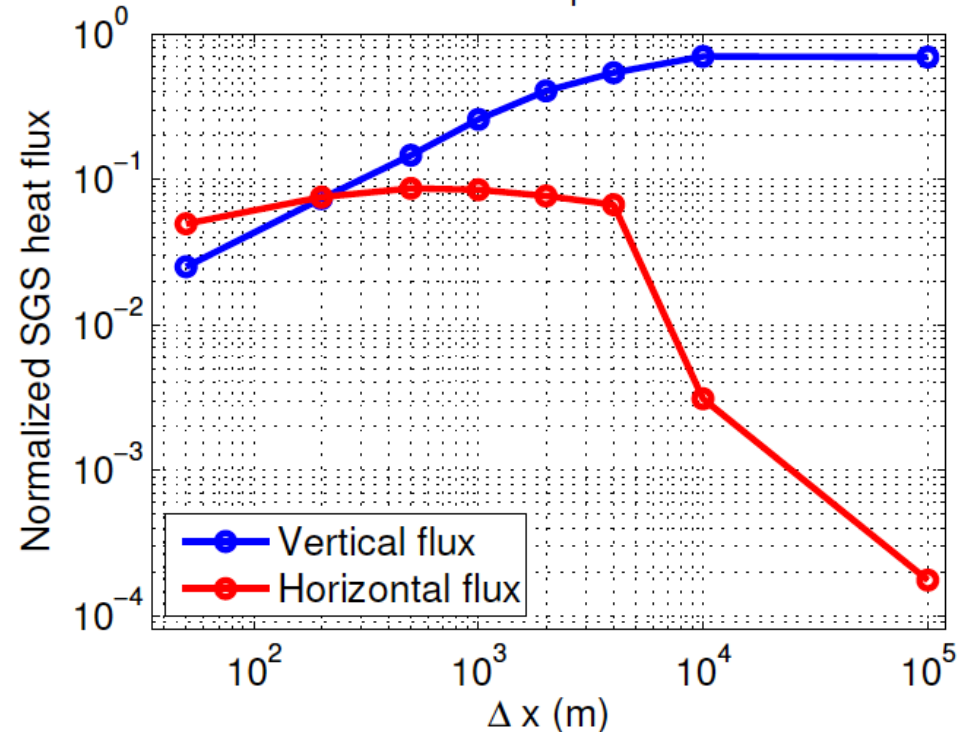
# Gradual transition from resolved to parameterized turbulence

## Turbulent Kinetic Energy



Partitioning between resolved and SGS TKE

## $0.1 < z/z_i < 0.5$



SGS vertical vs horizontal turbulent fluxes:  
Horizontal fluxes decrease significantly from 1 to 100 km

# SUMMARY

- **Unified** EDMF parameterization can represent boundary layer turbulence, shallow and deep convection (EDMF versions implemented into ECMWF, NAVGEM, NCEP)
- **Multiple Plumes**: New EDMF version using multiple plumes represents well shallow and deep convection
- Simple **scale-adaptive** approach leads to gradual transition from LES (50 m) to climate model resolutions (100 km)
- **Key Challenges**: Scale-adaptive plume models; Plume-plume interaction; Prognostic plumes; Coupling to microphysics; Stable boundary layer.